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Soil microbial community, dissolved organic matter and nutrient cycling interactions change along an elevation gradient in subtropical China

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29 **ABSTRACT**

30 To identify possible dominating processes involved in soil microbial community assembly,
31 dissolved organic matter (DOM) and multi-nutrient cycling (MNC) interactions and contribute to
32 understanding of climate change effects on these important cycles, we investigated the interaction
33 of soil chemistry, DOM components and microbial communities in five vegetation zones - ranging
34 from evergreen broad-leaved forest to alpine meadow - along an elevation gradient of 290 to 1960
35 m in the Wuyi Mountains, Fujian Province, China. Soil DOM composition and microbial
36 community assembly were characterized using Fourier transform ion cyclotron resonance mass
37 spectrometry (FT-ICR MS) and Illumina MiSeq high-throughput sequencing, respectively. Sloan's
38 neutral model and the modified stochasticity ratio were used to infer community assembly
39 processes. Key microbial drivers of the soil MNC index were identified from partial least squares
40 path models. Our results showed that soil DOM composition is closely related to the vegetation
41 types along an elevation gradient, the structure and composition of the microbial community, and
42 soil nutrient status. Overall, values of the double bond equivalent (DBE), modified aromaticity
43 index (AI_{mod}) increased, and H/C ratio and molecular lability boundary (MLBL) percentage
44 decreased with elevation. Lignins/CRAM-like structures compounds dominated soil DOM in each
45 vegetation type and its relative abundance decreased with elevation. Aliphatic/protein and lipids
46 components also decreased, but the relative abundance of aromatic structures and tannin increased
47 with elevation. The alpha diversity index of soil bacteria gradually decreased with elevation, with
48 deterministic processes dominating the microbial community assembly in the highest elevation
49 zone. Bacterial communities were conducive to the decomposition of labile degradable DOM
50 compounds ($H/C \geq 1.5$) at low elevation. In the cooler and wetter conditions at higher-elevation sites
51 the relative abundance of potentially resistant soil DOM components ($H/C < 1.5$) gradually increased.
52 Microbial community diversity and composition were important predictors of potential soil
53 nutrient cycling. Although higher elevation sites have higher nutrient cycling potential, soil DOM
54 was assessed to be a more stable carbon store, with apparent lower lability and bioavailability than
55 at lower elevation sites. Overall, this study increases understanding of the potential linkage
56 between soil microbial community, multiple nutrient cycling and DOM fate in subtropical
57 mountain ecosystems that can help predict the effect of climate change on soil carbon

58 sequestration and thus inform ecosystem management.

59

60 **Keywords:** dissolved organic matter (DOM), FT-ICR MS, labile components, microbial
61 community assembly, refractory components, soil multi-nutrient cycling

62

63 **1. Introduction**

64 Globally, the carbon (C) content of soil organic matter (SOM) is more than three times that of the
65 atmospheric C pool or the C storage of living terrestrial vegetation (Schmidt et al., 2011;
66 Gougoulias et al., 2014). Climate warming significantly affects the stock and stability of SOM and
67 increases the release of CO₂ from the soil carbon pool, thereby causing a positive feedback between
68 the terrestrial carbon cycle and climate change (Koven et al., 2011). Soil microorganisms play a
69 pivotal role in global C cycling (Abatenh et al., 2018), because a large proportion of SOM is derived
70 from storage and reprocessing within the microbial food web (Liang et al., 2019).

71 Although dissolved organic matter (DOM) accounts for less than 2% of SOM (Swenson et al.,
72 2015), it plays an important role in regulating nutrient cycles and soil micro-ecology as a dynamic
73 soil carbon pool (Fouché et al., 2020). Soil DOM compounds provide soluble organic substrates for
74 heterotrophic microbes and play a role in extracellular electron transfer across cell membranes,
75 changing the cellular redox state (Wang Y. H. et al., 2021). This alters the microbial niche and
76 thereby affects the composition and diversity of the functional microbial community (Li et al., 2019;
77 Mladenov et al., 2010; Wang Y. H. et al., 2021). The production and degradation of the different
78 components of DOM are inseparable from DOM-microbe interactions (Wang W. X. et al., 2021).
79 For example, low molecular mass plant-derived DOM molecules are consumed and transformed by
80 microorganisms, generating larger molecular mass DOM compounds with increasing soil depth
81 (Roth et al., 2019). Furthermore, changes in structure and composition of the microbial community
82 are tightly coupled to DOM turnover (Wu X. Q. et al., 2018). Analysis of changes in the temporal
83 and spatial distribution of DOM compounds and their chemodiversity is key for evaluating SOM
84 stabilization mechanisms. Fourier transform ion cyclotron resonance mass spectrometry (FT-ICR
85 MS) provides detailed information on various DOM compounds at the molecular level (e.g., Roth et
86 al., 2019; Wang W. X. et al., 2021; Wu X. Q. et al., 2018).

87 Artificial simulation warming experiments allow us to explore the response of SOM to
88 short-term temperature increase. For example, in a 4 °C soil warming experiment conducted for 4.5
89 years, it was reported that soil nutrient availability changed, which affected the composition and
90 metabolism processes of the microbial community (Dove et al., 2021). However, the applicability of
91 these results to understanding SOM processing under complex environmental conditions and over
92 long-time scales is questionable. Moreover, these experiments did not fully consider the impact of
93 substrate availability on microbial activities. An alternative approach is to use elevation gradients as
94 natural experiments to test the ecological and evolutionary responses of biological groups to climate
95 change (Körner, 2007). Climate, vegetation types, and soil heterogeneity vary greatly over short
96 spatial distances in mountain ecosystems (Tang et al., 2020). These differences have been exploited
97 in “temporal and spatial substitution” studies across geographic gradients (Rustad, 2008) in which
98 soil transplant experiments have been conducted along elevation gradients to explore the response
99 of surface SOM components to rising temperature (Cao et al., 2020; Luan et al., 2014).
100 Bioavailability of DOM, the most bioavailable fraction of SOM, is controlled by its intrinsic
101 characteristics, soil properties and external factors such as temperature and precipitation patterns
102 (Marschner and Kalbitz, 2003). The sharply heterogeneous environment along an elevation
103 gradient, such as precipitation, temperature, rainfall and vegetation type, may lead to variation in
104 water-extractable SOM (Huang et al., 2015). Hence, soil heterogeneity and environmental
105 variability have a great impact on the DOM inventory along elevation gradients. For example, Dai
106 et al. (2021) reported a positive correlation between the total genetic diversity of microorganisms
107 related to C cycling and elevation.

108 Niche and neutral processes are complementary in regulating the assembly of microbial
109 communities simultaneously (Chen Q. L. et al., 2021; Zhou and Ning, 2017). Niche-based theory
110 holds that the assembly of microbial community structure is controlled by deterministic processes
111 due to the different habitat preferences and fitness among microorganisms, including abiotic factors
112 (environmental filtering) and biotic interactions (such as facilitation and predation) (Chen et al.,
113 2019; Zhou and Ning, 2017). Thus, environmental heterogeneity is the dominant driver of microbial
114 assembly. Conversely, neutral theory asserts that microbial assemblage is generated by stochastic
115 processes, including probabilistic dispersal, random speciation and extinction, and ecological drift

116 (Zhou and Ning, 2017). A recent study showed that stochastic processes mainly determined the
117 assembly of *phoD*-harboring bacteria (an indicator of organic phosphorus transformation) at high
118 elevations (> 1500 m) with less environmental constraint, compared to the bacteria assembly at low
119 elevation (< 1500 m), in the pristine Shennongjia forest, north-western Hubei Province, China (Wan
120 et al., 2021). Nevertheless, the relationship between soil DOM compositions and its molecular
121 properties and the assembly patterns of microbial community along subtropical elevation gradients
122 remain largely unexplored.

123 Microbial community assembly determines the presence, abundance, and composition of
124 microbes, thereby impacting multiple ecosystem functions (Wan et al., 2021; Xun et al., 2019). One
125 of the methods for assessing ecosystem multi-functionality is to standardize and then average the
126 values of multiple functions to a single index (Byrnes et al., 2014). The multiple nutrient cycling
127 (MNC) index provides a quantifiable measure of an ecosystem to sustain multiple functions
128 simultaneously (Delgado-Baquerizo et al., 2016; Jiao et al., 2018; Jiao et al., 2021).
129 Microorganisms drive the soil nutrient cycling in terrestrial ecosystems (Zhou et al., 2023). Thus,
130 analyzing microbial community assembly processes and DOM compounds, and investigating how
131 microbial community diversity and composition drive soil multiple nutrient cycling, are needed for
132 predicting the effect of climate change on soil microbial community and the drivers of soil C
133 cycling.

134 The Wuyi Mountains have a distinct vertical zonation of vegetation types and strong climatic
135 variation along an elevation gradient in subtropical south-eastern China, providing a unique
136 experimental platform to explore soil-plant-microbial interactions (Wang et al., 2009; Wu et al.,
137 2016). To identify the soil microbial community assembly process, diversity patterns, and soil DOM
138 components, we applied FT ICR-MS to characterize soil DOM components, combined with
139 Illumina MiSeq sequencing, to soil samples along an elevation gradient in the Wuyi Mountains,
140 subtropical China. Specifically, we hypothesized that: i) the chemodiversity of DOM molecules and
141 the content of refractory components increases with elevation due to the shift to specialized
142 microbes with greater C degradation ability; ii) deterministic processes control soil bacterial and
143 fungal community assembly at high elevation; iii) changes in soil microbial diversity result in
144 variation in multi-nutrient cycling along the elevation gradient.

145

146 **2. Material and methods**

147 **2.1. Site description, soil sample collection and preparation, and soil chemical analysis**

148 Our study was conducted in the Wuyi Mountain National Nature Reserve (27°33'-27°54' N,
149 117°27'-117°51' E) in northern Fujian Province, south-eastern China. Five sites with different
150 vegetation zones along an elevation gradient were selected, which are long-term ecological study
151 sites described in Bu et al. (2011). The vegetation zones and elevations of the five sites were:
152 evergreen broad-leaved forest (EB), coniferous and broad-leaved mixed forest (CB), coniferous
153 forest (CF), subalpine dwarf forest (DF), and alpine meadow (AM), located at 290, 1070, 1400,
154 1800 and 1960 m above sea level, respectively. Additional information about soil type, climate
155 (mean annual air temperature (MAT), and mean annual precipitation (MAP)) and the main plant
156 species in each vegetation zone are in Table S1. Winter was selected as the optimal time for
157 sampling to explore soil microbial effects on DOM as previous research in the same area reported
158 higher soil microbial biomass in winter (He et al., 2009).

159 We randomly demarcated three plots (20 m × 20 m) in each vegetation zone. Eight cores (3 cm
160 diameter) were taken in the upper 20 cm of soil in an S-shape in each plot on 19 January 2019 (in
161 winter) at mid-slope positions. For each plot, all eight soil samples were mixed thoroughly to form
162 a single composite sample (~1 kg) and then placed in an airtight bag. Thus 15 soil samples were
163 generated, consisting of one composite sample from each of the three plots for each of the five
164 vegetation zones. The samples were stored at 4 °C for transport to the laboratory within 48 hours
165 (Roth et al., 2019). In the laboratory, fresh soil samples were sieved (2 mm) to remove gravel and
166 plant debris.

167 Part of the sieved soil was air-dried at room temperature and used to determine soil pH,
168 available phosphorus (AP) and available potassium (AK). Some of the 2-mm sieved air-dried soil
169 was ground by hand and passed through a 0.149 mm nylon sieve and used for determination of soil
170 total carbon (TC) and total nitrogen (TN) contents. The remaining fresh soil samples (< 2 mm) were
171 cryopreserved at -80 °C and then thawed as required to represent “fresh” soil for determination of
172 ammoniacal-nitrogen (NH₄⁺-N), nitrate-nitrogen (NO₃⁻-N), dissolved organic carbon (DOC), and
173 the characterization of DOM molecular composition and the soil microbial community. The

174 analysis methods for soil chemical properties and the molecular composition of soil DOM are
175 detailed in Supplementary Material Sections S1 and S2, respectively. We mixed equally the three
176 replicate soil samples at each site into one sample (Ye et al., 2020) for DOM extraction.

177

178 **2.2. Soil microbial community analysis**

179 Bacterial and fungal DNA were extracted from each fresh soil sample using a Power Soil DNA
180 Isolation Kit (MoBio Laboratories, Inc., CA). We subjected soil samples to high-throughput
181 sequencing using the Illumina Miseq PE300 sequencing platform at Beijing Allwegene Technology
182 Co., Ltd. (Beijing, China). DNA quality was detected by 1% agarose gel electrophoresis. The
183 primers 338F (5'-ACTCCTACGGGAGGCAGCAG-3') and 806R
184 (5'-GGACTACHVGGGTWTCTAAT-3') were used to amplify the V3–V4 region of the bacterial
185 16S rRNA. For fungal communities, the ribosomal ITS1-ITS2 region was targeted using primers
186 ITS1 (5'-CTTGGTCATTTAGAGGAAGTAA-3') and ITS2 (5'-TGC GTTCTTCATCGATGC-3').
187 The PCR reaction mixtures and PCR amplification procedures used for soil bacteria and fungi were
188 the same as in Zhang et al. (2022), and are detailed in Supplementary Material Section S3. Reads
189 were demultiplexed, quality-filtered, and processed to obtain valid sequences. Effective sequences
190 with a similarity of 97% were clustered into the same operational taxonomic unit (OTU).
191 Trimmomatic (v 0.36) and PEAR (v 0.9.6) were used to control the quality of raw sequence reads,
192 and the paired-end sequences were merged using Flash (v1.20) and PEAR (v0.9.6). Chimeric
193 sequences were removed with the Vsearch (v2.7.1) (uchime method) (Zhang et al., 2022). The
194 functional annotation of prokaryotic taxa (FAPROTAX) database was used for annotation of the
195 bacterial community (Sansupa et al., 2021) using the online tools
196 (<http://cloud.biomicroclass.com/CloudPlatform/SoftPage/FAP>).

197

198 **2.3. Soil multi-nutrient cycling analysis and DOM molecular properties**

199 We determined the soil MNC index from the seven nutrient variables measured: TC, TN, NH₄⁺-N,
200 NO₃⁻-N, AP, AK, and DOC. The MNC value was calculated for the average value of the
201 standardized score of each soil nutrient variable, normalized for individual nutrient concentrations
202 on a common scale ranging from 0 to 1 (Jiao et al., 2021; Zhang et al., 2022) (Eqs. 1 and 2).

203
$$STD_i = \frac{X_i - X_{min}}{X_{max} - X_{min}} \quad (\text{Eq. 1})$$

204
$$MNC = \frac{\sum_{i=1}^n STD_i}{n} \quad (\text{Eq. 2})$$

205 where STD_i is the standardized individual soil nutrient variable i and X_i , X_{min} , and X_{max} are the
206 individual soil nutrient concentrations and their minimum and maximum values across all soil
207 samples, respectively. n represents the number of all nutrient variables. Classification and
208 calculation of soil DOM molecular properties (DBE, AI_{mod} , average nominal oxidation state of
209 carbon (NOSC), molecular lability boundary (MLBL) percentage, and Shannon index) are detailed
210 in Supplementary Material Section S4 and Table S2.

211

212 **2.4. Data and statistical analysis**

213 Apart from where stated, all statistical analyses were conducted using R v4.1.3 (R Core Team, 2018)
214 and the significance level used was $P < 0.05$. Differences in soil chemical properties and the relative
215 abundance and alpha diversity indexes of microbes were tested using one-way analysis of variance
216 (ANOVA), with multiple comparisons among vegetation zones conducted using least significant
217 difference (LSD) tests (SPSSv19, SPSS Inc., Chicago, IL, USA). The Shapiro-Wilk test and
218 Levene's test were used to check that data fulfilled the normality and homogeneity of variance
219 assumptions for ANOVA, respectively. If these assumptions were not met, we applied the
220 non-parametric Kruskal-Wallis test with Bonferroni-adjusted P values using the "agricolae"
221 (v1.3-5) package (de Mendiburu, 2021).

222 We calculated the alpha diversity indices (Chao1 index, Shannon index, and Faith's
223 phylogenetic diversity whole tree index (Faith's PD whole tree index)) of microbial communities in
224 each of the 15 soil samples using Qiime (v.1.8.0) (Taketani et al., 2017). Differences in microbial
225 communities between elevations and the degree of separation between- and within- elevations were
226 evaluated using the NMDS and analysis of similarities (ANOSIM) (Chen et al., 2019).

227 We calculated Spearman's rank correlation coefficients between the soil bacterial and the
228 fungal OTUs with the top 300 relative abundances using the "Hmisc" and "igraph" packages in R.
229 We used the False Discovery Rate (FDR) method to adjust the P values. Only correlations with
230 absolute correlation coefficient values > 0.8 correlations and adjusted P values < 0.01 were

231 considered. The "igraph" package was also used to calculate the topological characteristics of the
232 co-occurring network compared with 10,000 random networks generated according to the Erdős–
233 Rényi model whose edges were assigned to any node with the same probability (Jiao et al., 2016).
234 The degree distributions of the co-occurring networks were plotted to assess their form (Fig. S1).
235 Gephi (0.9.3) (<https://gephi.org/>) was used to visualize the co-occurrence networks.

236 To assess the contribution of neutral processes to microbial community assembly, we fitted
237 Sloan’s neutral community model (NCM) to estimate the relationship between the occurrence
238 frequency and mean relative abundance variations of OTUs, as described in detail in Jiao et al.
239 (2021). The modified stochasticity ratio (MST) to infer microbial community assembly processes
240 (Qiu et al., 2020) and Levins’ niche breadth index were calculated for the soil microbial community
241 using the “NST” package and the “niche.width” function in the “spaa” package (Chen Q. L. et al.,
242 2021), respectively. The "EcolUtils" package was used to randomly arrange the occurrences of
243 OTUs 1000 times to simulate the OTU occurrence frequency. OTUs with observed occurrence
244 exceeding the upper 95% confidence interval were considered generalists, with those below the
245 lower 95% confidence interval considered specialists, and the remainder considered neutral taxa
246 (Zhang et al., 2018).

247 Associations between soil DOM composition and microbial OTUs were assessed by Pearson
248 product-moment correlation analysis and visualized as co-occurrence network diagrams using
249 Cytoscape 3.5.1. The “random-forest” package was used to identify the main microbial predictors
250 of soil MNC values along the elevation gradient. The most important predictors were assessed as
251 those with higher percentage increases in the MSE (mean squared error) of the variable. The "A3"
252 package was used to assess the significance of the model with 5000 permutations of the response
253 variables (Gao et al., 2021; Jiao et al., 2018), whilst the "rfPermute" package was used to estimate
254 the significance of each predictor on the MNC index with 5000 trees. We used partial least squares
255 path models (PLS-PM) through the “plspm” package to investigate the direct and indirect effects of
256 microbial community characteristics on the MNC index. From initial modeling, variables with
257 loading values < 0.7 and variance inflation factor (VIF) > 10 were removed, with the final PLS-PM
258 constructed using the remaining variables (Yun et al., 2022). DOM data used in this manuscript can
259 be found at <https://www.zenodo.org/record/8188815>.

260

261 3. Results

262 3.1. Soil chemical properties and DOM composition along the elevation gradient

263 Soils were acidic and pH values were very similar among sites, ranging from mean pH of 4.55 in the
264 subalpine dwarf forest (DF) vegetation zone to 4.73 in the alpine meadow (AM) zone (Table 1). Soil
265 $\text{NH}_4^+\text{-N}$ concentration generally increased with elevation, whilst $\text{NO}_3^-\text{-N}$ and AP did not vary
266 significantly between zones. DOC concentration increased with elevation, reaching a mean of 109.4
267 mg kg^{-1} at AM, double that in the EB and coniferous and broad-leaved mixed forest (CB) vegetation
268 zones. Soil TN and TC concentrations also increased significantly with elevation, although the C/N
269 ratio remained relatively constant. In general, the MNC index increased with elevation (Table 1).

270 An UpSet plot were used to visualize the common-shared, partial-shared, and unique DOM
271 molecules for the five vegetation zones along the elevation gradients in the Wuyi Mountains (Fig.
272 S1). The number of detected molecular formulae of soil DOM at the five different vegetation zones
273 ascending with altitude from evergreen broad-leaved forest (EB) to AM was 1168, 3106, 1206, 3392,
274 and 2325, respectively. The common shared molecules (present in all five samples) accounted for
275 the largest proportion from the intersection of different vegetation zones. Inter sample rankings
276 analysis was performed to reveal possible differences in quality in the 703 common DOM
277 components present in the five vegetation zones (Fig. 1). DOM qualities differed considerably
278 across the vegetation zones at different elevations. At the EB site, DOM molecular formulae
279 ranked 1 (relatively high peak intensities) exhibited $\text{H/C} > 0.9$ and $\text{O/C} \leq 0.5$, whilst those ranked
280 4 and 5 mainly had $\text{O/C} > 0.5$. In contrast, at sites DF and AM, DOM formulae ranked 4 and 5
281 mainly exhibited $\text{H/C} > 1$ and $\text{O/C} \leq 0.55$. The DOM molecular formulae ranked 1 at site AM
282 were dominated by $\text{H/C} < 1$, whilst formulae ranked 1 and 2 at DF displayed high m/z values ($>$
283 400) with $\text{O/C} > 0.5$. Unique DOM molecular formulae at DF and AM were dominated by tannin-
284 and aromatic structures- like compounds, respectively (Fig. S2).

285 Lignins/CRAM-like (carboxylic rich alicyclic molecule) structures dominated soil DOM
286 composition in all vegetation zones, and its abundance decreased with increasing elevation (Fig.
287 S3). The relative abundance of aliphatic/proteins and lipids also decreased with elevation, whilst
288 those of tannin and aromatic structures increased with elevation (Fig. S3d). The weighted mean

289 O/C ratio, AI_{mod} , and NOSC values of DOM molecules showed increasing trends with elevation,
290 whilst the H/C ratio and molecular lability boundary (MLB_L) percentage decreased (Table 2). The
291 Shannon index values of DOM molecules in the coniferous forest (CF) were slightly lower than in
292 the other vegetation zones (Table 2). This is consistent with the narrower ranges of the distribution
293 curves for the number of C and O atoms and M/Z values of DOM compositions in CF compared
294 to the other vegetation zones (Fig. S3).

295

296 **3.2. Soil microbial diversity, community structure and assembly processes along the elevation** 297 **gradient**

298 Within the soil bacterial community, *Acidobacteria* (36.2%–42.6%), *Proteobacteria* (24.5%–
299 32.3%), *Chloroflexi* (7.51%–14.2%), *Actinobacteria* (5.97%–8.4%) and *Planctomycetes* (2.14%–
300 5.77%) were the five most abundant phyla in all five vegetation types (numbers in parentheses are
301 minimum and maximum values of mean % abundance amongst the different sites, Fig. S4a). They
302 displayed no clear trend in abundance with elevation, although the relative abundance of
303 *Verrucomicrobia* (0.99%–2.69%) decreased with elevation (Fig. S3b). In the fungal community,
304 *Basidiomycota* (53.7%–82.5%) dominated at all sites. Its relative abundance was highest at the
305 lowest elevation in the evergreen broad-leaved forest (EB) zone, and then decreased before
306 increasing with elevation. The relative abundance of *Mortierellomycota* phylum also generally
307 increased with elevation (Fig. S4c).

308 The Chao1 index and Faith's PD index of soil bacteria decreased with increasing elevation (Fig.
309 S5a and Fig. S6a), whilst the indices for fungi were not significantly different between elevations
310 (Fig. S5d and Fig. S6b). The Shannon indices for both the bacterial and fungal communities
311 increased from EB to the coniferous and broad-leaved mixed forest (CB) zone and then showed a
312 decreasing trend with elevation (Fig. S5b, e). Non-metric multidimensional scaling (NMDS)
313 analysis showed that the soil bacterial and fungal communities at the different sites formed distinct
314 clusters, which are completely separate in the ordination space (Anosim test, Global $R = 1$, $P <$
315 0.001) (Fig. S5c, f). Beta diversity, estimated among the microbial communities in all 15 soil
316 samples based on Bray-Curtis dissimilarities, was significantly lower among bacteria than fungi
317 (Fig. S7).

318 The potential linkages among soil bacteria and fungi OTUs were revealed by the co-occurring
319 network analysis (Fig. 2a, b). Compared with the Erdős-Rényi random network, a higher clustering
320 coefficient and average path length were observed in the co-occurrence networks of bacterial and
321 fungal communities (Fig. 2a, b). This displayed that the generated networks were more clustered
322 than the identically sized random networks and that their degree distributions are non-random. The
323 modularity and average path length of the fungal community were greater than for the bacterial
324 network. In the fungal community most connections were positive (93%), in contrast to the bacterial
325 community where only 55% of connections were positive.

326 Neutral processes dominated assembly of the bacterial and fungal communities overall along
327 the elevation gradient, with 72.9% and 56.6% of the community variation explained by the neutral
328 community model (NCM), respectively (Fig. 2c, d). The values of m (migration rate) were 0.561
329 and 0.123 for the bacterial and fungal community, respectively (Fig. 2c, d). To detect the relative
330 importance of stochastic and deterministic processes in the assembly of the bacterial and fungal
331 communities at different elevations, we calculated the modified stochasticity ratio (MST) in each
332 vegetation zone (Fig. 2e). The MST value for the fungal community in CB was significantly higher
333 than in other vegetation zones. In addition, MST values were < 0.5 for both bacteria and fungi in the
334 AM zone.

335

336 **3.3. Correlation between soil chemical properties, DOM components, and microbial** 337 **community characteristics**

338 Correlation analysis across all elevations revealed that the Chao 1 index for the bacterial community
339 was significantly negatively correlated with tannins, aromatic structures, and the AI_{mod} value of
340 DOM compounds (Fig. 3a). Conversely, lignins/CRAM-like structures and aliphatic/proteins
341 components were significantly positively correlated with the bacterial Chao1 index (Fig. 3a). The
342 relative abundance of lignins/CRAM-like structures and aliphatic/proteins compounds were
343 significantly negatively correlated with soil TC, NH_4^+ -N, and TN contents ($P < 0.05$). However,
344 the relative abundance of tannins was significantly positively correlated with soil TN and TC
345 contents (Fig. S8).

346 Highly significant correlations ($P < 0.01$) between the 100 most abundant DOM molecular

347 formulae and bacteria and fungi OTUs at the phylum level were selected to construct network
348 diagrams. For bacteria (Fig. 3b), formulae of lignins/CRAM-like structures were positively
349 correlated with most microbes, whilst *Chloroflexi* was negatively correlated with some tannin
350 molecules. *Proteobacteria* was only positively correlated with individual DOM molecules, whilst
351 *Chloroflexi* only had negative correlations. The number of C atoms in DOM formulae correlated
352 with *Proteobacteria* had a small range of 15–21, while those correlated with *Acidobacteria* had a
353 relatively large range of 13–23. For fungi (Fig. 3c), there were a smaller number of correlations
354 between phyla and individual DOM molecules, and all correlations were positive and with
355 lignins/CRAM-like structures, apart from one formula. More DOM molecular formulae were
356 correlated with *Ascomycota* than with *Basidiomycota*.

357

358 **3.4. Microbial drivers of soil multi-nutrient cycling**

359 We applied random forest (RF) analysis to identify the major potential microbial drivers of soil
360 multiple nutrient cycling across all sites. Site elevation, alpha diversity (PD whole tree and Chao1
361 indices), and the relative abundance of *Verrucomicrobia* and *Firmicutes* were the most important
362 bacterial community indicators of the soil multi-nutrient cycling (MNC) index (Fig. 4a). The
363 significant fungal community predictors were site elevation, relative abundance of three fungal
364 phyla, and the alpha diversity (PD whole tree and Shannon indices). Partial least squares path
365 models (PLS-PM) analyzed the effects of microbial community characteristics on the soil MNC
366 index, which had best fit goodness-of-fit (GOF) values of > 0.70 for microbial communities (Fig.
367 4c). The PLS-PM also highlighted the strong direct effects of elevation on soil microbial
368 community diversity, although the effects are in opposite directions: high elevation has a negative
369 effect on the bacterial alpha diversity index (path coefficient = -0.91, $P < 0.001$) but a positive
370 effect on the fungal alpha diversity index (path coefficient = 0.59, $P < 0.05$). The Shannon
371 diversity index of fungal community exhibited direct negative effects on MNC (path coefficient =
372 -0.56, $P < 0.05$).

373

374 **4. Discussion**

375 **4.1. Refractory components of soil DOM molecules gradually increase along the elevation**

376 **gradient**

377 In our study elevation transect in the Wuyi Mountains, the MLB_L (%) decreased and AI_{mod} and
378 NOSC values of DOM molecules increased with elevation. Since higher AI_{mod} and NOSC values
379 indicate greater recalcitrance (low bioavailability) of DOM compounds (Cai and Jiao, 2023; Hu et
380 al., 2022), this indicates that soil DOM resistance increased with elevation. The DBE and O/C ratio
381 of soil DOM molecules increased along our study elevation gradient (Table 2), which is consistent
382 with the elevational variation in DOM components in restored areas of the Loess Plateau (Hu et al.,
383 2021), but contrasts with results in the higher altitude Sygera Mountains on the Tibetan Plateau
384 (Zhang et al., 2022). The DBE value reflects the degree of unsaturation of the DOM molecules
385 (Melendez-Perez et al., 2016), with higher values indicating greater unsaturation and lower
386 bioavailability (Cai and Jiao, 2023). Ligand exchange between the carboxyl / hydroxyl functional
387 groups of the DOM compounds and the surface of Fe oxides is the main mechanism of adsorption of
388 soil DOM molecules (Ding, 2020). An experimental study demonstrated that adsorption of the more
389 oxidized DOM components increased at low DOM concentrations, whilst at high DOM
390 concentrations selective adsorption of the less oxidized DOM components occurred (Avneri-Katz et
391 al., 2017). These processes appear to occur along our study elevation transect, with more oxidized
392 DOM components remaining in the soil liquid phase at the high-elevation sites. In contrast, at the
393 lower elevation sites, the oxidized DOM components were more readily adsorbed at soil solid phase
394 surfaces (and not extracted for analysis in our ultrapure water extraction (Bahureksa et al., 2021)),
395 thus decreasing the C accessibility for microbial degradation. Soil iron-containing minerals retain
396 great quantities of SOM (Lv et al., 2017). Highly oxidized and more aromatic DOM molecules are
397 preferentially adsorbed on hematite surfaces (Lv et al., 2017), and preferential adsorption of high
398 molar mass organic solutes occurs on goethite surfaces (Liu et al., 2014). The soil in the Wuyi
399 Mountains is highly leached, with mineral composition dominated by hematite in low-elevation red
400 earth soils and goethite at higher elevations in yellow earth soils (Lin, 2010). Soil mineralogy,
401 therefore, helps explain the lower molar mass (M/Z) and O/C values of the DOC molecules
402 extracted at the lowest elevation EB site.

403 The abundances of DOM molecules with lipids- and aliphatic/proteins-like compositions
404 decreased with increasing elevation on our study transect. Conversely, aromatic- and tannin- like

405 compounds increased, indicating DOM compounds derived from microbe to plants along the
406 elevation gradient (Shen et al., 2023). Tannins are secondary metabolites of polyphenols produced
407 by higher plants, which can leach from litter into soil, and are toxic to microbial metabolism and
408 can inhibit enzyme activity, thereby affecting nutrient cycling (Triebwasser et al., 2012; Kraus et
409 al., 2003). Feng et al. (2021) showed that soil microbial metabolic efficiency decreased with
410 elevation, which was attributed partly to the microbial community increasing investment in nutrient
411 acquisition via enzymes. Thereby, the decrease with elevation of microbial-derived DOM along
412 our elevation transect is attributed to the expected inhibition of microbial activity. We also found
413 that the relative abundance of lignin/CRAM-like structure compounds decreased with increasing
414 altitude. Lignins in soil are mainly generated from above-ground and subsurface (root) litter of
415 higher plants, though biotic, aerobic and co-metabolic degradation processes (Thevenot et al.,
416 2010). For example, white rot fungi of *Basidiomycetes* secrete ligninolytic enzymes, which can
417 convert and mineralize refractory DOM into CO₂ and water by co-metabolizing aliphatic and
418 aromatic substances (Agrawal et al., 2021). The change of lignin input source during the transition
419 from trees to graminaceous plants and the strong degradation of saprophytic fungi at higher
420 elevation may explain the change of lignin-like compounds to some extent along the study
421 transect.

422 Interestingly, the DOM molecules in soil from the coniferous forest (CF) site had a narrower
423 range of number of C and O atoms and M/Z values (Fig. S3), which is consistent with the results of
424 Li et al. (2021) for soil DOM of monoculture Chinese fir
425 (*Cunninghamia lanceolata* (Lamb.) Hook.) plantations. Soil DOM provides carbon-containing
426 substrates for the soil microbial community, and the latter affects DOM composition by supplying
427 metabolites and residues (Wu M. et al., 2021). Plant litter is an important main source of soil DOC
428 (Don and Kalbitz, 2005). Therefore, differences in litter-fall and the metabolites and residues of
429 microorganisms in vegetation zones along elevation gradients may be one important cause of
430 changes in DOM molecular composition. For instance, the low input of soil nutrients from
431 litter-fall (Ma et al., 2007) may partly explain the soil DOM variation in Chinese fir plantations of
432 different ages.

433 The van Krevelen diagram is widely used for visualization of DOM molecular types and

434 characteristics, and identification of possible chemical reactions (Kim et al., 2003). Nevertheless,
435 it is important to note that the molecular composition (molecular formulae) only of DOM is
436 obtained based on FT-ICR MS but not the isomer features (Hu et al., 2022). It is therefore
437 necessary in future studies to identify the structural features of DOM molecules to reveal their full
438 role in soil ecosystems (Qi et al., 2022).

439

440 **4.2. Deterministic processes dominate the low-diversity microbial assembly at high elevation**

441 The MST values of the soil bacteria and fungi communities were below the 0.5 threshold in the
442 alpine meadow zone, indicating that deterministic processes dominate microbial assembly at the
443 highest elevation site (Fig. 2e). Furthermore, the microbial community at this site had low Shannon
444 diversity indices compared to many of the lower elevation vegetation zones (Fig. S5b, e).
445 Environmental filters, including temperature, precipitation, soil available nutrient contents and
446 physical properties, and vegetation types can select for microbial species with specific traits, such as
447 stress tolerances and nutrient acquisition (Anthony et al., 2020). Such effects have been reported in
448 a number of studies. For instance, microbial community diversity and enzyme activity were
449 affected more by changing climatic factors, such as mean annual temperature and precipitation,
450 than soil characteristics along an elevation gradient on Taibai Mountain in eastern China (Ren et
451 al., 2020). Mo et al. (2021) reported that deterministic processes influenced the microeukaryotic
452 plankton community assembly with narrow habitat niche breadths under high salinity conditions.
453 Furthermore, Xun et al. (2019) indicated that environmental selection increased in low-diversity
454 bacterial communities, which contributed to the dominance of deterministic processes. Reduced soil
455 temperature and increased soil moisture were observed with elevation in the Wuyi Mountains
456 related to lower air temperature and elevated precipitation (Table S1), which results in slower
457 decomposition rates of SOM at high elevation (Bu et al., 2012). Collectively, increasing
458 environmental and physiological stress on microorganisms along the elevation gradient (Ren et al.,
459 2020), including the reduced abundance of labile DOM ($H/C \geq 1.5$), high C/N ratio substrates
460 (Anthony et al., 2020) and increasing abundance of tannins, are suggested to lead to the dominance
461 of deterministic assembly processes for the microbial community at the highest elevation site in the
462 Wuyi Mountains.

463 Co-occurring network analysis is a useful method for identifying potential microbial
464 interactions and functional distribution (Chen W. J. et al., 2021). The results of co-occurring
465 network analysis showed that positive associations were more dominant among the fungal
466 community than the bacterial community at the study sites (Fig. 2a, b), suggesting possible niche
467 overlap and mutualistic or facilitative interactions between the fungal species (Zhu et al., 2021). The
468 lower average path length and higher average degree of the bacterial co-occurring network
469 indicates closer relationships among the OTUs and rapid distribution of any perturbation (Li et al.,
470 2020; Yuan et al., 2021).

471 The emergence and maintenance of soil microbial diversity are underpinned by soil
472 heterogeneity (Nunan et al., 2020). In this study, soil fungal communities had higher beta diversity
473 than bacterial communities (Fig. S7) which is consistent with the results of Jiao et al. (2018). The
474 bacterial community across all sites had lower path length, lower dispersal limitation (lower *m*
475 value), and wider habitat niche breadths than fungal taxa (Fig. 2a, c, d and Fig. S9), which have a
476 smaller body size and more flexible metabolic plasticity, and therefore are less environmentally
477 filtered (Wu W. X. et al., 2018). The interaction between disturbance and dispersal can homogenize
478 species composition, which was observed to cause decreased beta diversity (Catano et al., 2017).
479 The relative proportions of soil heterotrophic microorganisms are determined by the local
480 geographic environment, vegetation types, and soil factors (Zhuang et al., 1997). For example, Li et
481 al. (2021) found that the soil fungal Faith's PD index declined with elevation in the cold temperate
482 zone at high latitudes in China. Microbial diversity positively regulates the service functions of
483 terrestrial ecosystems by changing the nutrient supply and distribution of terrestrial ecosystems
484 (Delgado-Baquerizo et al., 2016).

485 Overall, in this study, the Chao 1 and Faith's PD diversity indexes of bacterial communities
486 decreased along a subtropical elevation gradient (Fig. S5a and Fig. S6a). Ma et al. (2022) also
487 reported that the bacteria community richness index and Faith's PD index decreased with elevation
488 in subtropical forests, with soil temperature largely explaining the changes in soil bacterial and
489 fungal diversity. Bacterial community diversity largely drives the utilization efficiency of C
490 (Domeignoz-Horta et al., 2020). Reduced abundance of labile DOM ($H/C \geq 1.5$) and increasing
491 abundance of tannins may reduce soil enzyme activity, thereby reducing bacterial community

492 diversity.

493 Microbial life history strategies also influence soil DOM composition. For example, Zeng et al.
494 (2022) found that microbial assemblies shifted from using labile to recalcitrant carbon with
495 increasing elevation. In our study, *Chloroflexi* had higher relative abundance at the highest
496 elevation AM site (Fig. S4a) and also its abundance was negatively correlated with
497 lignins/CRAM-like structures and tannins in DOM molecules (Fig. 3b). Positive correlations
498 between microbial compositions and DOM molecules are indicative of the generation of new
499 molecules by microbial activity, whilst negative correlations are interpreted as the decomposition
500 of DOM molecules based on resource-consumer relationships (Hu et al., 2022). *Chloroflexi* can
501 decompose recalcitrant C-containing components and are considered a *K*-selected microbe
502 (Adamczyk et al., 2021). Combined with our study results, this suggests the selection of substrates
503 with low bioavailability by *Chloroflexi*.

504

505 **4.3. Bacteria drive the degradation of labile soil DOM components, and the fungal** 506 **community is the key control on soil nutrient cycling**

507 The soil bacterial community contained a slightly higher percentage of habitat generalists and had
508 a larger niche breadth compared to the fungal community, which had a higher percentage of
509 habitat specialists (Fig. S9), indicating that the soil bacterial community had more competitive
510 ability than the fungi. Highly labile substrates, for which bacterial communities have a high affinity
511 (Osterholz et al., 2016), can reach steady-state equilibrium concentrations below detection limits.
512 The significant correlations across the study sites between the Chao1 index for bacteria and the H/C
513 and O/C ratios of DOM components (positive and negative, respectively) (Fig. 3a), indicate that the
514 bacterial community plays an important role in the degradation of labile components. The Chao1
515 index of the bacterial community was also significantly positively correlated with soil DOM
516 aliphatic/protein compounds, concurring with other studies reporting positive associations between
517 bacterial abundance and activity and aliphatic DOM components (Kamjunke et al., 2019). From the
518 results of the FAPROTAX database functional annotation, we found that bacterial OTUs with
519 functional photoheterotrophy increased with elevation, whilst those with functional nitrification
520 and aerobic nitrite oxidation decreased overall with altitude (Fig. S10). Furthermore, the soil MNC

521 index increased along the altitude gradient (Table 1). Thus, changes in microbial investment in
522 nutrient acquisition may lead to this interaction between soil nutrients and microbial communities.

523 Fungal extracellular polymeric substances and hyphae networks are conducive to the
524 transmission and aggregation of plant-derived litter derivatives to increase the decomposition
525 efficiency and persistence of SOM (De Beeck et al., 2021; Witzgall et al., 2021). Fungi are
526 regarded as oligotrophs, exhibiting effective degradation of refractory carbon substrates (Ho et al.,
527 2017). The soil fungal community plays a vital role in nutrient cycling and organic matter
528 transformation, such that fungal community structure and biomass provide an early indicator of soil
529 environmental changes (Wu et al., 2019). Specifically, the relative abundance of the fungal phylum
530 *Mortierellomycota* increased with elevation and was significantly positively correlated with soil AK
531 and DOC concentrations (Fig. S4c and S11a). *Mortierellomycota* are *r*-selected saprotrophic
532 microorganisms that have a low life expectancy and stronger growth in environments with high
533 available resources and low-stress exposure (Wu et al., 2021).

534 In contrast, the relative abundance of the bacterial phylum *Verrucomicrobia* was significantly
535 negatively correlated with soil $\text{NH}_4^+\text{-N}$, AK, and DOC concentrations (Fig. S11b, c), and with
536 tannin abundance (Fig. S12). This is interpreted as indicative of the oligotrophic lifestyle of
537 *Verrucomicrobia* (Navarrete et al., 2015). In our study, the relative abundance of tannins were
538 significantly positively correlated with soil TC and TN concentrations (Fig. S8), which showed
539 that tannins was beneficial to the sequestration of carbon and nitrogen. The dominant fungal
540 phylum *Basidiomycota* in our study had the highest relative abundance in the lowest elevation
541 vegetation zone EB, and then decreased, followed by an increasing trend with elevation from CB to
542 AM (Fig. S4c). *Basidiomycota* are representative of oligotrophic fungi, which can degrade
543 refractory compounds, and were found to be most abundant in warmer soil conditions along a
544 temperate forest transect characterized by a K-strategy fungal community (Li et al., 2021). Thus the
545 increase in the relative abundance of *Basidiomycota* with elevation in our study can partly explain
546 the shift to K strategy due to the refractory components ($\text{H/C} < 1.5$) of soil DOM with elevation.

547

548 **5. Conclusions**

549 It is well-known that the soil microbial community has a profound impact on DOM composition and

550 multi-nutrient cycling in ecosystems. Changes in the functional group of soil bacterial community
551 with elevation, related to investment in nutrient acquisition, may account for changes in soil
552 nutrient cycling. Neutral processes dominated overall the bacterial and fungal communities
553 assembly along the elevation transect, and bacteria have a higher niche breadth and lower dispersal
554 limitation than fungi. However, deterministic selection dominated the microbial community
555 assembly at the highest elevation site, which may be attributed to the reduced soil temperature and
556 increased soil moisture, and increased relative abundance of tannin and aromatic structures in soil
557 DOM. This increased abundance of potentially resistant DOM components could be interpreted as
558 enhanced soil carbon sequestration capacity with increasing elevation, but their stability is not
559 assured if vegetation zones at lower elevations with higher rates of DOM degradation move
560 upwards in response to climate warming. As a relatively active fraction of the soil C pool, DOM
561 molecules are greatly affected by microbial community, and anthropogenic-driven climate change
562 increases the possibility of DOM degradation. Since sampling was conducted in winter, care should
563 be taken in interpreting the interconnection between soil DOM components and microorganisms in
564 the Wuyi Mountains in other seasons.

565

566

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918 **Author Contributions**

919 All authors contributed to the study conception and design. **Shuzhen Wang**: Investigation,
920 Validation, Data Curation, Data analysis, Writing - first draft; **Kate V. Heal**: Supervision, Writing -
921 review & editing; **Qin Zhang**: Validation, Data analysis; **Yuanchun Yu**: Supervision; **Mulualem**
922 **Tigabu**: Reading and polishing the manuscript; **Shide Huang**: Soil sampling, **Chuifan Zhou**:
923 Conceptualization, Visualization, Resources, Supervision. All authors commented on previous
924 versions of the manuscript. All authors read and approved the final manuscript.

925

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929 University.

930

931 **Figure captions**

932 **Fig. 1.** Inter sample rankings analysis of common-shared DOM molecules in five vegetation zones
933 along the elevation gradient in the Wuyi Mountains, Fujian Province, China. Codes for the
934 different vegetation zones - EB: evergreen broad-leaved forest; CB: coniferous and broad-leaved
935 mixed forest; CF: coniferous forest; DF: subalpine dwarf forest; AM: alpine meadow.

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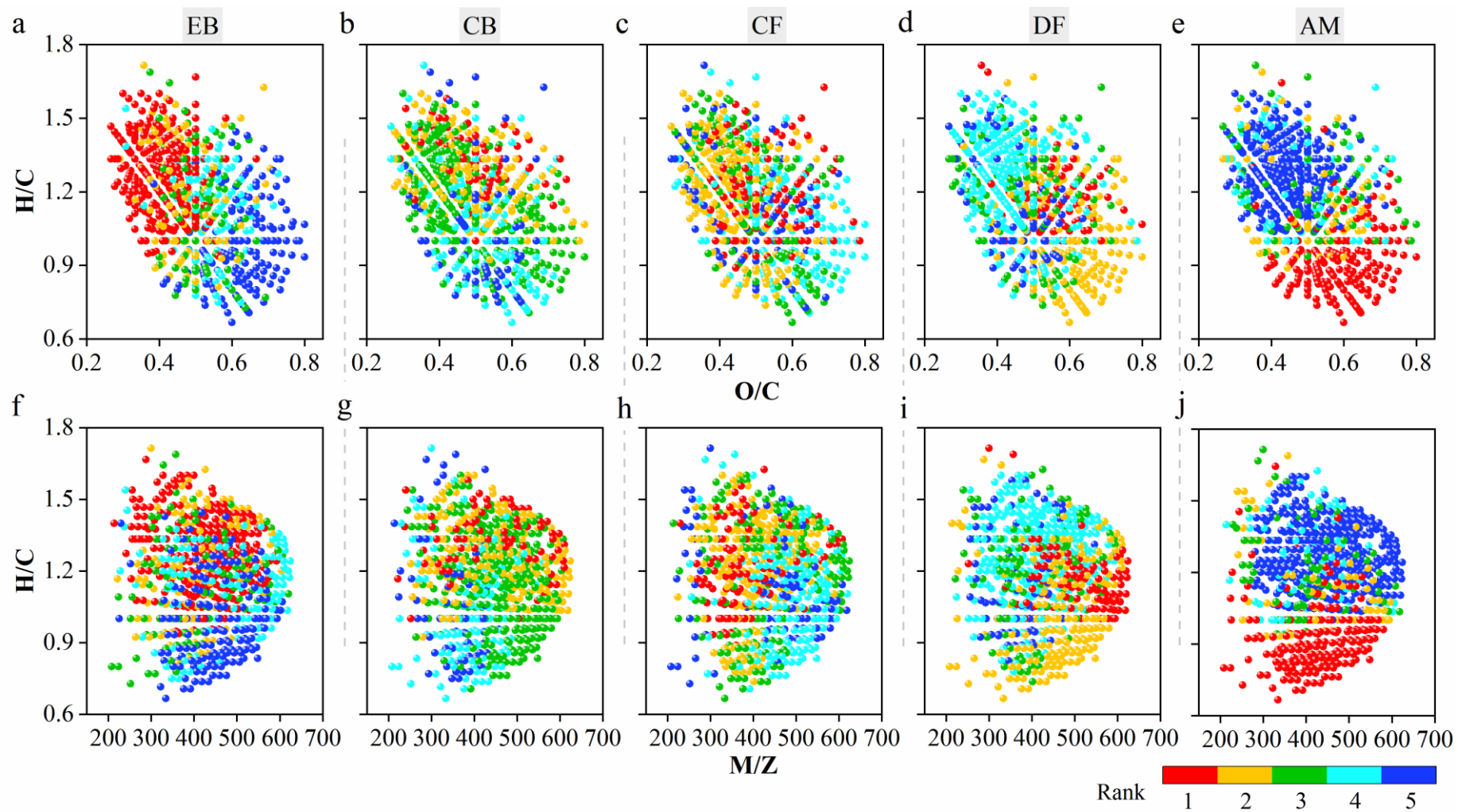
937 **Fig. 2.** Co-occurring network diagrams for soil bacteria (a) and fungi (b) (top 300 OTUs) across
938 all soil samples along the elevation gradient in the Wuyi Mountains, Fujian Province, China. The
939 box below each of (a) and (b) contains summary characteristics of the results of 10,000 random
940 networks generated for bacteria and fungi, respectively, according to the Erdős–Rényi model. Fit
941 of Sloan’s neutral community model (NCM) to the OTUs of the soil (c) bacterial and (d) fungal
942 community assemblies for all soil samples. The solid black line represents the best-fitting neutral
943 model and the dashed black lines represent the 95% confidence intervals around it. Blue points
944 between the dashed black lines represent OTUs that follow the neutral process. The green and
945 orange points indicate OTUs that occur more and less frequently than predicted by the NCM,
946 respectively. R^2 indicates the fit to the NCM, m is the estimated migration rate, and Nm is the
947 product of metacommunity size and m values. (e) Bar plots of the modified stochasticity ratio
948 (MST) values of the soil bacterial and fungal community assemblies in different vegetation zones
949 along the elevation gradient. The horizontal dashed purple line (MST = 0.5) was set as the
950 boundary between deterministic (MST < 0.5) and stochastic (MST > 0.5) assembly processes.
951 Values are means \pm standard error ($n = 3$). Different lowercase letters for bacteria/fungi indicate
952 significant differences between vegetation zones at different elevations according to LSD tests (P
953 < 0.05). See Fig. 1 for explanation of the codes for the different vegetation zones.

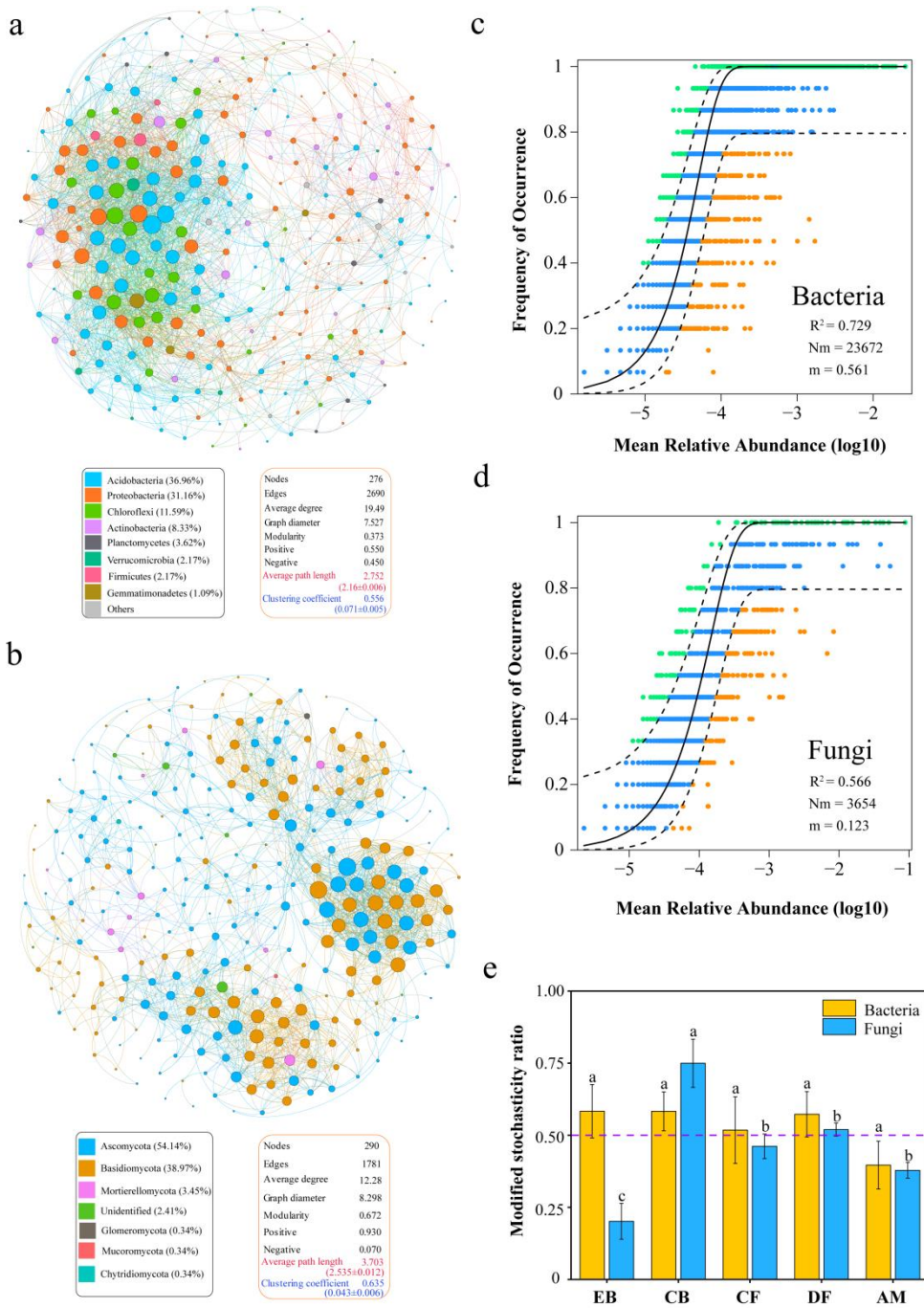
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955 **Fig. 3.** (a) Heatmap of correlation between microbial community characteristics (Chao1, Shannon
956 and Faith's PD indices, modified stochasticity ratio (MST)) and site elevation and DOM
957 components/characteristics (defined in Supplementary Material Section S1). *, $0.01 \leq P < 0.05$; **, $0.001 \leq P < 0.01$ indicate significant coefficients. Co-occurrence network diagrams of significant
958 ($P < 0.01$) correlations between the 100 most abundant soil DOM molecular formulae and the 100
959 highest relative abundance OTUs for bacteria (b) and fungi (c) in soil samples across the elevation
960 gradient of different vegetation zones. Blue circles represent the OTUs of bacterial and fungal
961 phyla. Green and red circles represent the lignin/CRAM-like structure and tannin molecules,
962 respectively. The sizes of the circles indicate the relative abundance of the OTUs and DOM
963 molecular formulae. The blue and red lines indicate significant positive and negative correlations,
964 respectively, between DOM formulae and microbial phyla.

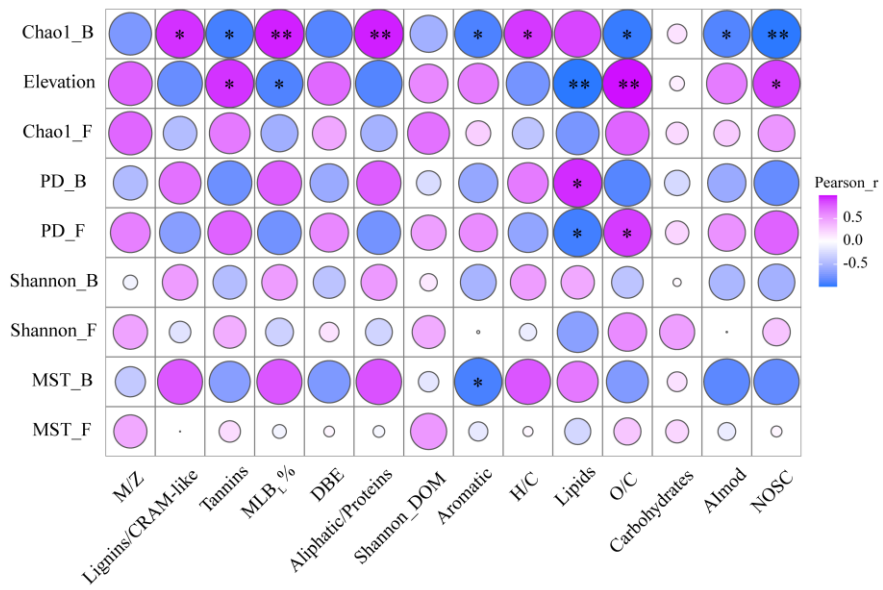
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967 **Fig. 4.** Random forest (RF) analysis to identify the importance of different microbial predictors of
968 the multiple nutrient cycling (MNC) index of the 15 soil samples along an elevation gradient in
969 the Wuyi Mountains, Fujian Province, China. (a) Bacterial community predictors: relative
970 abundance of the 10 most abundant bacterial phyla, alpha diversity indices (Shannon index, Chao1
971 index, Faith's PD index), the site pH, and elevation. (b) Fungal community predictors: relative
972 abundance of the 3 most abundant fungal phyla, alpha diversity indices (Shannon index, Chao1
973 index, Faith's PD index), the site pH, and elevation. Partial least squares path models (PLS-PM)
974 showing the effects of soil microbial communities characteristics on MNC (c). The line width is
975 proportional to the magnitude of the path coefficient. Numbers adjacent to arrows indicate the
976 direct effect size of the relationship. The black and blue lines indicate positive and negative path
977 coefficient, respectively. R^2 indicates the variance of the dependent variable explained by the
978 model. GOF is the goodness-of-fit index of the model. *, $0.01 \leq P < 0.05$; **, $0.001 \leq P < 0.01$;
979 ***, $P < 0.001$ indicate significant predictors or path coefficients.

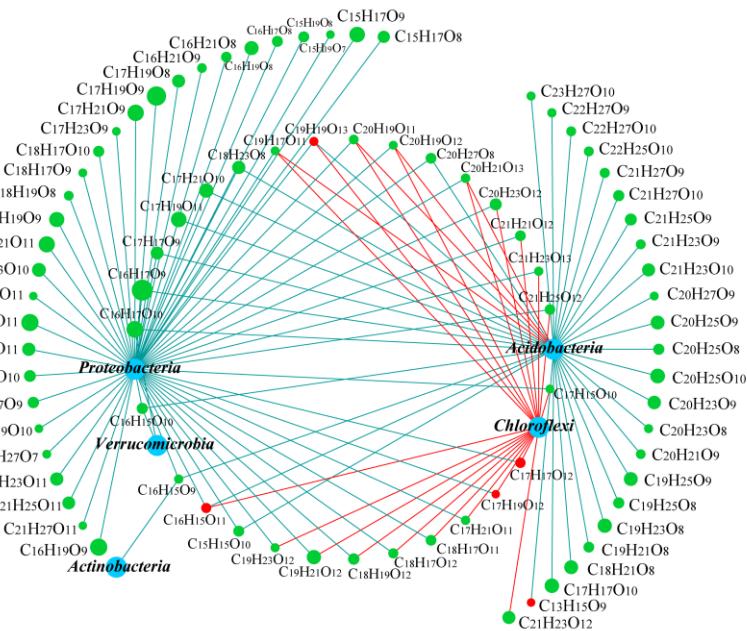




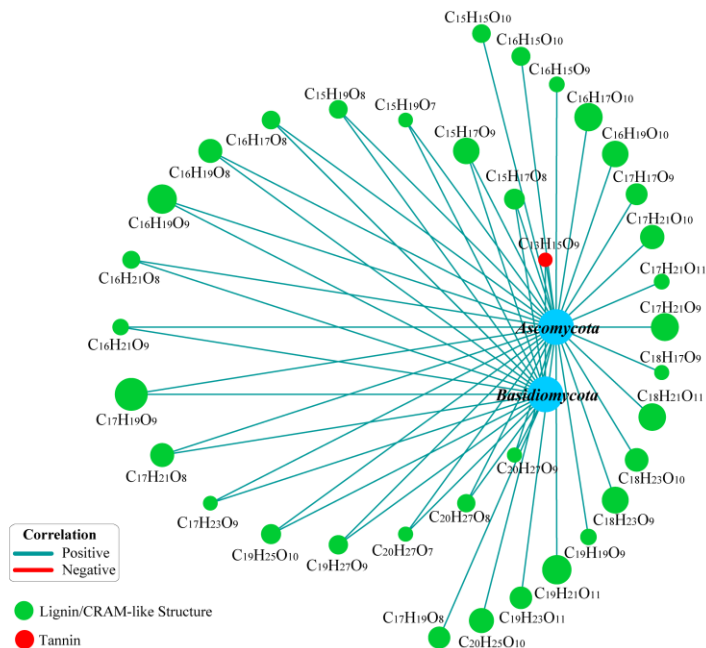
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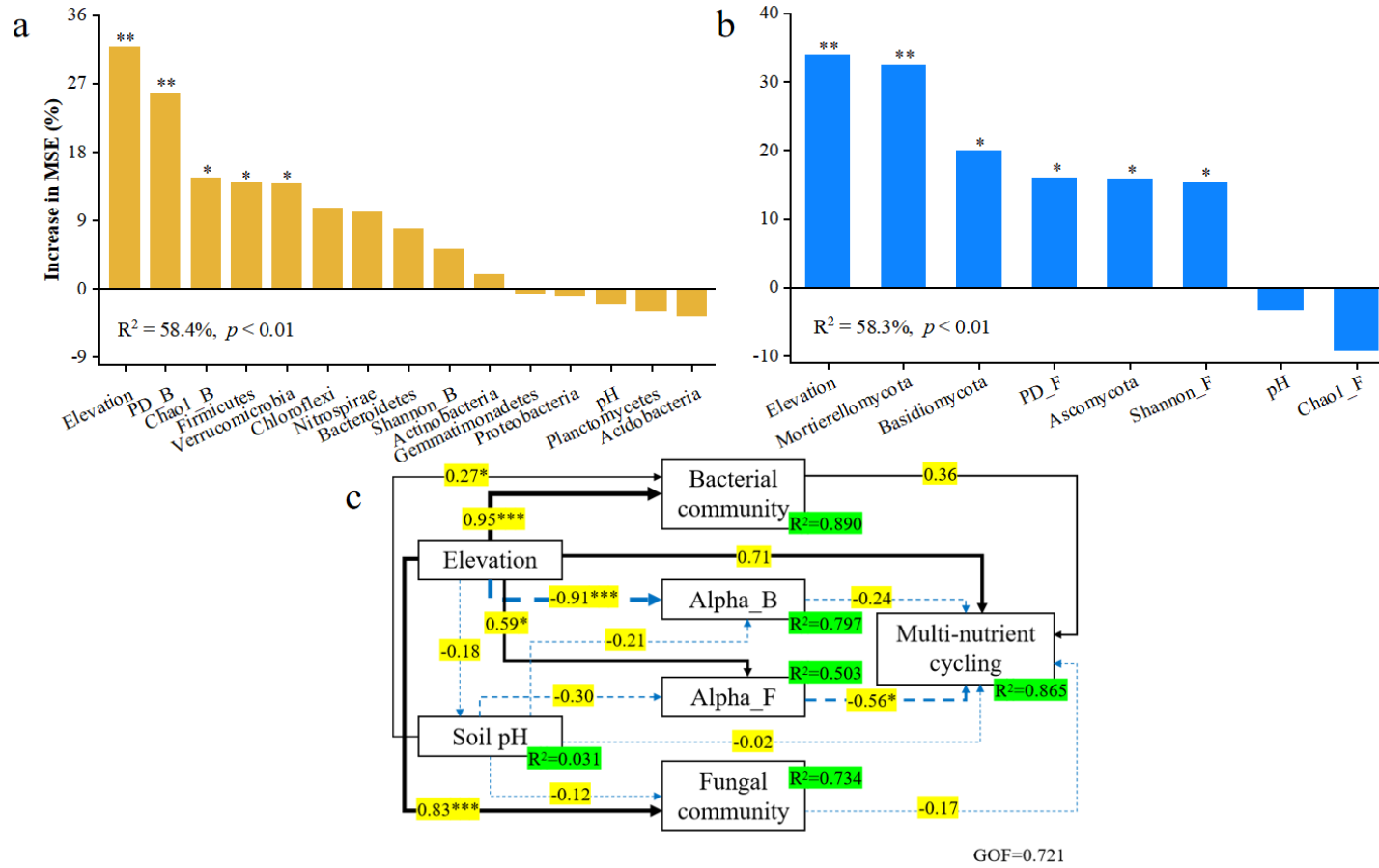
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985
 986 **Fig. 4**
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988

989 **Tables**

990 **Table 1.** Soil properties in different vegetation zones along an elevation gradient in the Wuyi Mountains, Fujian Province, China. All data are presented as means \pm
 991 standard error ($n = 3$). EB: evergreen broad-leaved forest; CB: coniferous and broad-leaved mixed forest; CF: coniferous forest; DF: subalpine dwarf forest; AM:
 992 Alpine meadow). Ammoniacal-nitrogen ($\text{NH}_4^+\text{-N}$); nitrate-nitrogen ($\text{NO}_3^-\text{-N}$); available phosphorus (AP); available potassium (AK); dissolved organic carbon (DOC);
 993 total nitrogen (TN); total carbon (TC); C/N ratio: TC/TN; MNC: soil multiple nutrient cycling index, calculated as explained in Section 2.5. $\text{NO}_3^-\text{-N}$, $\text{NH}_4^+\text{-N}$, and
 994 DOC reported as mg kg^{-1} fresh weight. Different lowercase letters in a row indicate significant differences between elevations ($P < 0.05$).

Soil property	Vegetation zone and elevation				
	EB (290 m)	CB (1070 m)	CF (1400 m)	DF (1800 m)	AM (1960 m)
pH	4.70 \pm 0.04ab	4.64 \pm 0.09ab	4.58 \pm 0.04ab	4.55 \pm 0.04b	4.73 \pm 0.03a
$\text{NH}_4^+\text{-N}$ (mg kg^{-1})	12.71 \pm 0.64a	15.23 \pm 1.70a	13.38 \pm 0.28a	14.23 \pm 0.62a	18.20 \pm 2.18a
$\text{NO}_3^-\text{-N}$ (mg kg^{-1})	1.89 \pm 0.60a	1.11 \pm 0.04a	1.99 \pm 0.10a	1.48 \pm 0.21a	1.37 \pm 0.11a
AP (mg kg^{-1})	5.38 \pm 1.09b	5.11 \pm 0.49b	11.64 \pm 4.89ab	7.95 \pm 0.41ab	11.67 \pm 0.93a
AK (mg kg^{-1})	156.96 \pm 15.91ab	113.33 \pm 16.53b	130.85 \pm 20.32b	200.07 \pm 7.48a	223.74 \pm 16.17a
DOC (mg kg^{-1})	54.28 \pm 4.06c	56.21 \pm 8.07c	93.93 \pm 9.80ab	80.08 \pm 4.71b	109.43 \pm 4.36a
TN (g kg^{-1})	1.47 \pm 0.26c	1.20 \pm 0.55c	3.42 \pm 0.21b	4.53 \pm 0.30b	6.96 \pm 0.67a
TC (g kg^{-1})	20.88 \pm 1.23cd	19.29 \pm 7.70d	45.53 \pm 1.54bc	60.29 \pm 3.12ab	98.15 \pm 14.76a
C/N	14.78 \pm 1.78a	16.80 \pm 0.89a	13.36 \pm 0.39a	13.36 \pm 0.50a	14.01 \pm 1.34a
MNC	0.22 \pm 0.03c	0.16 \pm 0.07c	0.41 \pm 0.03b	0.43 \pm 0.02b	0.68 \pm 0.08a

995 **Table 2.** Intensity weighted average (wa) values of the different characteristics of DOM molecules and relative abundance of different element combinations from
 996 soils along an elevation gradient, Wuyi Mountains, Fujian Province, China. $n = 1$ as a composite sample for the three plots at each elevation was analyzed by FT-ICR
 997 MS. DBE: double bond equivalent; NOSC: average nominal oxidation state of carbon; AI_{mod} : modified aromaticity index; $MLBL$: molecular lability boundary. The
 998 calculation methods are described in Supplementary Material Section S4. Vegetation zone codes are as given in Table 1.

Vegetation zone code and elevation									Characteristics of soil DOM molecules		
	DBE wa	M/Z wa	H/C wa	O/C wa	NOSC wa	AI_{mod} wa	$MLBL$ %	Shannon index	CHO %	CHON %	CHOS %
EB (290 m)	8.49	409.15	1.25	0.48	-0.26	0.25	14.47	6.78	88.39	8.12	3.48
CB (1070 m)	9.20	439.60	1.21	0.52	-0.16	0.26	11.63	7.42	90.43	5.54	4.03
CF (1400 m)	8.37	414.03	1.24	0.53	-0.13	0.24	11.44	6.70	80.17	9.46	10.37
DF (1800 m)	10.20	455.60	1.13	0.55	0.004	0.29	7.96	7.60	86.79	7.83	5.39
AM (1960 m)	11.59	453.73	0.98	0.57	0.19	0.39	3.20	7.45	88.53	10.91	0.56

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 1000
 1001
 1002