



THE UNIVERSITY *of* EDINBURGH

Edinburgh Research Explorer

Wave resource assessment for Scottish waters using a large scale North Atlantic spectral wave model

Citation for published version:

Venugopal, V & Nimaladinne, R 2015, 'Wave resource assessment for Scottish waters using a large scale North Atlantic spectral wave model', *Renewable Energy*, vol. 76, pp. 503-525.
<https://doi.org/10.1016/j.renene.2014.11.056>

Digital Object Identifier (DOI):

[10.1016/j.renene.2014.11.056](https://doi.org/10.1016/j.renene.2014.11.056)

Link:

[Link to publication record in Edinburgh Research Explorer](#)

Document Version:

Peer reviewed version

Published In:

Renewable Energy

General rights

Copyright for the publications made accessible via the Edinburgh Research Explorer is retained by the author(s) and / or other copyright owners and it is a condition of accessing these publications that users recognise and abide by the legal requirements associated with these rights.

Take down policy

The University of Edinburgh has made every reasonable effort to ensure that Edinburgh Research Explorer content complies with UK legislation. If you believe that the public display of this file breaches copyright please contact openaccess@ed.ac.uk providing details, and we will remove access to the work immediately and investigate your claim.



1 **Wave Resource Assessment for Scottish Waters Using a Large Scale North Atlantic**
2 **Spectral Wave Model**

3
4 **Vengatesan Venugopal*** and **Reddy Nimalidinne**

5
6 Institute for Energy Systems, School of Engineering
7 The University of Edinburgh, Edinburgh, EH9 3JL
8 United Kingdom

9
10 **Abstract**

11 This paper reports the methodology established in the application of a numerical wave model for
12 hindcasting of wave conditions around the United Kingdom, in particular for Scottish waters, for
13 the purpose of wave energy resource assessment at potential device development sites. The phase
14 averaged MIKE21 Spectral wave model has been adopted for this study and applied to the North
15 Atlantic region bounded by latitudes 10° N - 70° N and longitudes 10° E-75° W. Spatial and
16 temporal wind speeds extracted from the European Centre for Medium Range Weather Forecast
17 (ECMWF) have been utilised to drive the wave model. A rigorous calibration and validation of
18 the model has been carried out by comparing model results with buoy measurements for different
19 time periods and locations around Scotland. Significant wave height, peak wave period and peak
20 wave direction obtained from the model correlated very well with measurements. Spatially
21 varying statistical mean and maximum values of the significant wave height and wave power
22 obtained based on a one-year wave hindcasting are in good agreement with the UK Marine Atlas
23 values. The wave model can be used with high level of confidence for wave hindcasting and even
24 forecasting of various wave parameters and wave power at any desired point locations or for
25 regions. The wave model could also be employed for generating boundary conditions to small
26 scale regional wave and tidal flow models.

27 **Keywords:** Wave modelling, Spectral wave model, Orkney and Pentland waters, hindcasting,
28 wave power, wave parameters.

29 * Corresponding Author, V.Venugopal@ed.ac.uk

30 **1. Introduction**

31 Electricity generation from ocean waves and tidal current is an active research worldwide and a
32 number of successful technologies are now being investigated in many parts of the globe. Several
33 of these wave/tide power converters, are either being installed and tested currently or already
34 connected to grids (reNews [1]). According to reNews, the total wave and tidal technologies
35 installed in Scotland alone until now sums to 6.365 MW, and the rest of the countries in the world
36 contributed to only 6.56 MW. The Pentland Firth (see Figure 1), which is the region between the
37 north-east tip of Scotland and the south of Orkney Islands, is considered to be one of the best
38 sites in the world for generating electricity from tidal stream. Figure 1 also indicates the strategic
39 potential sites, licensed by the Crown Estate [2], where wave and tidal energy devices will be
40 deployed by various developers.

41
42 In Scotland, the Aquamarine Power [3] installed its Oyster 800 wave power machine at the
43 European Marine Energy Centre (EMEC) facility in Orkney at a water depth of 13 m and
44 commenced operational testing in June 2012. The company claims it produced the first electrical
45 power to the grid in the same month. The company is likewise planning to deploy its next-
46 generation machine Oyster 801 side by side, thus creating a wave farm. In addition, the
47 Aquamarine Power now has been consented from the Scottish Government to develop a 40MW
48 wave farm off the north-west coast of Lewis, Scotland, which will include the deployment of 40
49 to 50 Oyster devices along the coast of Lewis.

50
51 Pelamis Wave Power [4], another wave device developer, has also deployed and tested its
52 Pelamis P2 machine at the EMEC facility in Orkney, the Billia Croo test site, for Scottish Power
53 Renewables. The Pelamis P2 was installed at EMEC for the first time in May 2012 at a water

54 depth of approximately 50 m. Pelamis wave power plans to install 66 Pelamis machines for a 50
55 MW production off the Marwick Head in Orkney, for which the company claims to have an
56 agreement for lease awarded by the Crown Estate. In addition to the above two, few other wave
57 and tidal power companies, eg., Alstom [5], Andritz Hydro Hammerfest [6], AW Energy
58 technologies [7] , Voith Hydro [8] and Wello Oy [9], have also tested their technologies at EMEC
59 sites. Further details may be found in [10] for tidal power and [11] for wave power technologies.

60
61 As demonstrated above, Scotland, in particular Orkney, Pentland Firth and Outer Hebrides,
62 indeed, have become potential regions where both wave and tidal energy technologies can be
63 successfully installed and operated. Scotland is geographically well placed on the globe where
64 large energetic waves from the North Atlantic Ocean provide high level of sustainable wave
65 power resources; however, harvesting these energy sources increase the number of challenges
66 associated with it. An accurate estimation of wave conditions is essential not only for the
67 evaluation of wave power, but also to estimate normal operational and extreme wave scenarios
68 for assessing the survivability and economic viability of the technology and predicting any
69 associated risks.

70
71 The UK target is to source 15% of its energy from renewables by 2020, with a commitment to
72 target an 80% reduction in CO2 emissions by 2050. The Scottish Government has committed to
73 the development of a successful marine renewable energy industry in Scotland and targeting to
74 achieve 20% of European Union's energy consumption from renewable sources by the year 2020
75 [12]. Scotland's target is to produce up to a 25% of Europe's tidal power and 10% of its wave
76 power from the seas around it.

77

78 To speed up these targets, several funding schemes have been developed and the UK's
79 Engineering and Physical Sciences Research Council (EPSRC), under its SUPERGEN Marine
80 Challenge - Accelerating the Deployment of Marine Energy (Wave and Tidal) scheme, has
81 funded several projects one of which is the 'TeraWatt: Large Scale Interactive coupled 3D
82 Modelling for Wave and Tidal Energy Resource and Environmental Impact' consortium. The
83 work reported in this paper is part of the research carried out for the TeraWatt project which
84 would concentrate on the questions: (i) what is the best way to assess the wave and tidal resource
85 and the effects of energy extraction, (ii) what are the physical consequences of wave and tidal
86 energy extraction and (iii) what are the ecological consequences of wave and tidal energy
87 extraction. In order to address the above questions, an accurate wave and or tidal resource
88 mapping must be produced for the regions where technology deployment activities are planned.

89
90 Although, there have been several wave modelling studies carried out in the past for North
91 Atlantic and the UK seas, the purpose of them were manifold. For example, Swail et al., [13] and
92 Swail et al., [14] investigated the longer term variation in ocean wave parameters for North
93 Atlantic using a discrete spectral type wave model called OWI 3-G driven by the NCEP/NCAR
94 global reanalysis wind data. Dodet et al., [15] studied the variability in the North-East Atlantic
95 Ocean using a 57-year hindcast (1953–2009), obtained with the wave model WAVEWATCH III
96 (Tolman, [16]), which was forced with 6-hours wind fields from the NCEP/NCAR Reanalysis
97 project. The spatial resolution of the wind input used for this work was 1.875° (longitude) by
98 1.905° (latitude) on a Gaussian grid. Their aim was to investigate changes in significant wave
99 height, mean wave direction and peak wave period. Galanis et al [17] explored the characteristics
100 of significant wave height by statistical approach for North Atlantic Ocean using satellite records
101 and simulated records using the WAM wave model (WAMDI Group [18]). They have produced

102 North Atlantic wide Weibull distribution's 'shape parameters' and 'scale parameters' which
103 would fit the significant wave height hindcast by WAM and also from the satellite records.

104
105 Numerical models potentially play several important roles in the assessment of marine energy
106 resources and they also serve to identify commercially exploitable sites. An UK wide wave
107 power resource Atlas has already been produced by ABPmer [19], however the limitations with
108 this Atlas, is that, this was produced based on the wave information made available from the UK
109 Met Office's UK Waters Wave model with a spatial resolution of 12 km and Global Wave Model
110 with a spatial resolution of 60 km. While sufficiently useful information can be obtainable from
111 this Atlas, for specific sites or regions such as Orkney and Pentland Firth where the sea
112 bathymetry is highly variable within a short horizontal space, and also considering future large
113 array scale developments which might require wave information on a spatial scale less than 12
114 km, it becomes obvious that development of a finer scale wave model capable of accurately
115 providing wave conditions in shallow, intermediate and deep water depths is highly essential,
116 which is what attempted in this work. While a large number of public domain numerical wave
117 models are available, based on the industry partners discretion within the TeraWatt consortium, it
118 has been recommended to use the commercially available MIKE 21 suite [20] for this research,
119 as the results produced could be adaptable by the industry partners for their use as MIKE 21 suite
120 appears to be a common popular and highly preferable tool among them. For the present work
121 the authors propose to use the commercial software MIKE 21 spectral wave model [20] with
122 wind input at 0.125 by 0.125 deg resolutions, for hindcasting wave conditions and wave power
123 for North Atlantic ocean, but focussing mainly on the potential wave energy development
124 locations around Scotland.

125 An overview of the wave model, bathymetry and mesh construction, methodology adopted in
126 selecting parameters which describe model physics, boundary conditions, calibration and
127 validation of the model to various locations and time periods, analysis of results and the
128 evaluation of performance indices etc, have been detailed in the sections below. It is anticipated
129 that this model would be very useful for hindcasting and forecasting wave conditions for seas
130 around the UK and Scotland, and serve as a tool for supplying boundary hydrodynamic
131 parameters for small scale regional wave and tidal models. Further, this wave model, when run
132 for longer time periods, would supply wave conditions to estimate site specific extreme wave
133 parameters for device designs and assessing survivability limits. Moreover, this model can
134 provide site specific wave parameters for assessing environmental impact (eg., sediment transport
135 change patterns) and ecological consequences of energy extraction.

136

137 **2. Wave model overview**

138 The spectral wave module from MIKE 21 suite [20] has been selected for the simulation of
139 waves and it is a widely used numerical tool by both the scientific community and industry
140 worldwide. The model simulates the growth, decay and transformation of wind-generated sea and
141 swells in offshore and coastal areas. This model accounts for the wave growth by the action of
142 wind, non-linear wave-wave interaction, dissipation of energy due to white-capping, bottom
143 friction and depth-induced wave breaking, refraction and shoaling, wave-current interaction and
144 the effect of time-varying water depth. A cell-centred finite volume method is applied in the
145 discretization of the governing equations in geographical and spectral space and a multi-sequence
146 explicit method is applied for the wave propagation with the time integration carried out using a
147 fractional step approach. This model produces phase averaged wave parameters as output for the
148 computational area.

149 The wind waves are expressed by the wave action density spectrum $N(\sigma, \theta)$, where σ is the
 150 relative (intrinsic) angular frequency and θ is the direction of wave propagation. The relative
 151 angular frequency can be related to the absolute angular frequency (ω) by the linear dispersion
 152 relationship as,

$$154 \quad \sigma = \sqrt{gk \tanh(kd)} = \omega - \bar{k} \cdot \bar{U} \quad (1)$$

155 where, g is gravity constant; k is wave number; d is water depth; \bar{U} is current velocity vector
 156 and \bar{k} is wave number vector with magnitude k and direction θ .

157
 158 MIKE 21 spectral wave model includes two methods of wave simulation namely, (i) the
 159 directional decoupled parametric formulation and (ii) the fully spectral formulation, both based
 160 on the wave action conservation equations in either Cartesian (for small scale applications) or
 161 spherical (for large scale applications) co-ordinate systems (Komen et al, [21], Young, [22]). The
 162 first formulation is based on a parameterisation of zeroth and first order moment of the wave
 163 action spectrum as dependent variables, whereas the second formulation involves the directional
 164 frequency wave action spectrum as the dependent variable.

165
 166 The wave action density spectrum $N(\sigma, \theta)$ can be related to the energy density $E(\sigma, \theta)$ by the
 167 relation

$$168 \quad N(\sigma, \theta) = \frac{E(\sigma, \theta)}{\sigma} \quad (2)$$

169 In the fully spectral formulation, the governing equation is the wave action balance equation. The
 170 conservation equations for wave action in Cartesian co-ordinates is given by

171
$$\frac{\partial N}{\partial t} + \nabla \cdot (\bar{v}N) = \frac{S}{\sigma} \quad (3)$$

172 where, $N(\bar{x}, \sigma, \theta, t)$ is the action density, t is the time, $\bar{v} = (c_x, c_y, c_\sigma, c_\theta)$ is the propagation velocity
 173 (as expressed in eqns 4-6) of a wave group in the four dimensional phase space, \bar{x}, σ, θ & t ,
 174 ∇ is the four-dimensional differential operator and S is the source term for energy balance
 175 equation. The wave group propagation velocities $c_x, c_y, c_\sigma, c_\theta$ in four-dimensional phase space are:

176
$$(c_x, c_y) = \frac{d\bar{x}}{dt} = \bar{c}_g + \bar{U} = \frac{1}{2} \left(1 + \frac{2kd}{\sinh(2kd)} \right) \frac{\sigma}{k} + \bar{U} \quad (4)$$

177
$$c_\sigma = \frac{d\sigma}{dt} = \frac{\partial \sigma}{\partial d} \left[\frac{\partial d}{\partial t} + \bar{U} \cdot \nabla_{\bar{x}} d \right] - c_g \bar{k} \cdot \frac{\partial \bar{U}}{\partial s} \quad (5)$$

178
$$c_\theta = \frac{d\theta}{dt} = -\frac{1}{k} \left[\frac{\partial \sigma}{\partial d} \frac{\partial d}{\partial m} + \bar{k} \cdot \frac{\partial \bar{U}}{\partial m} \right] \quad (6)$$

179
 180 where, $\nabla_{\bar{x}}$ is the two dimensional differential operator in the \bar{x} space, $\bar{x} = (x, y)$ is the Cartesian
 181 co-ordinates, s is the space co-ordinate in wave direction θ , and m is the co-ordinate
 182 perpendicular to s . The source function term S is given by

183
$$S = S_{in} + S_{nl} + S_{ds} + S_{bot} + S_{surf} \quad (7)$$

184 where, S_{in} is the momentum transfer of wind energy to the wave generation; S_{nl} is the energy
 185 transfer due to non-linear wave –wave interaction; S_{ds} is the energy dissipation of wave energy
 186 due to white-capping; S_{bot} is the energy dissipation due to bottom friction; S_{surf} is the energy
 187 dissipation due to depth-induced breaking.

188 The source functions S_{in} , S_{nl} and S_{ds} are similar to WAM Cycle 4 model (Komen et al., [21],
189 WAMDI Group [18] and the wind input is based on Janssen's [23-24] quasi-linear theory. Further
190 details can be found in [20].

191

192 **3. Model set-up**

193 ***3.1 Bathymetry and mesh generation***

194 An unstructured computational mesh (see Figure 2) was constructed using MIKE 21 mesh
195 generator and this covered the North Atlantic region $10^{\circ}\text{E} - 75^{\circ}\text{W}$ and $10^{\circ}\text{N}-70^{\circ}\text{N}$. It is well
196 known fact that swells generated in the Atlantic Ocean travels a long way to reach Scottish
197 waters, and they tend to carry higher energy which is highly beneficial to wave energy
198 community. As the objective was to capture long distance swells propagating towards the UK,
199 although it takes up high computing resources, this had been the reason for selecting such a large
200 computational domain. The sea water depth data compiled from the sources, GEBCO [25] and
201 Marine Scotland [26] have been used to generate the bathymetry within the computational
202 domain as can be seen in Figure 2. The grid resolution of the GEBCO bathymetry data was 30 arc
203 seconds, which was used for most of the model domain except for Orkney, Pentland Firth,
204 Shetland and north-west coast of Lewis regions which were covered by the Marine Scotland's
205 measured bathymetry data. The size of both GEBCO and Marine Scotland data sets were too
206 large for the MIKE 21 mesh generator to manage at one time, hence a data filter had been applied
207 with the purpose to reduce the data size but without losing data integrity, which resulted in the
208 size of the spatial grids reduced to 100 m (Easting) x 100 m (Northing) for the entire Orkney and
209 Pentland waters and device deployment locations around it, and also along the coast in the north.
210 For the rest of the UK, Ireland and English Channel, about 2 km x 2 km grid data resolution was

211 selected. In the Icelandic and Faroe Islands regions about 5 km x 5 km, in the North Sea 3 km x 6
212 km, and for the rest of the North Atlantic deep water locations 10 km x 10 km spacing were used.

213
214 For triangulation of unstructured mesh the ‘Natural neighbour’ interpolation method [20], which
215 is a geometric estimation technique that uses natural neighbourhood regions generated around
216 each point in the data set and suitable for dealing with a variety of spatial data themes with
217 clustered or highly linear distributions, has been selected. In total 71,793 elements with various
218 mesh resolutions have been produced for the entire computational domain as shown in Figure 2.
219 Highly finer mesh resolutions with a mesh area of 0.0005 square degrees for Orkney and
220 Pentland waters, 0.001 square degrees for the Hebrides and North West regions of Scotland (see
221 Figure 3) and 0.75 square degrees for North Atlantic Ocean were used. Such a high resolution in
222 the mesh was necessary for describing the shallow water hydrodynamics within the model; and
223 also for providing input boundary conditions if another small scale model is involved. One such
224 exercise for a 3-dimensional combined wave and tidal flow model, which was separately
225 constructed for the Terawatt project, can be seen in Venugopal and Nimalidinne [27].

226

227 ***3.2 Model forcing and physical processes***

228 The model was forced with 10 m level U- and V- wind speed data obtained from the operational
229 products of the European Centre for Medium-Range Weather Forecasts (ECMWF [28]) at 6 hrs
230 interval with a spatial resolution of $0.125^\circ \times 0.125^\circ$. The model was run in ‘fully spectral’ mode
231 as briefed in section 2 above, with ‘Instationary formulation’. The ‘coupled’ type of air-sea
232 interaction has been chosen for the wind boundary input with a Charnock parameter of 0.01.
233 According to the coupled model, the sea roughness, z_o is given by [20]

234

235
$$z_o = \frac{z_{charnock} u_*^2}{g} \left(1 - \frac{\tau_w}{\rho_{air} u_*^2} \right)^{-1/2} \quad (8)$$

236 where, $z_{charnock}$ is the Charnock parameter, u_* is the friction velocity which was calculated by
 237 Janssen [24] by assuming a logarithmic profile for wind speed, the τ_w is the wave induced stress,
 238 ρ_{air} is the density of air and g is the gravity constant.

239
 240 The fully spectral formulation is a computationally expensive technique, however when the model
 241 is forced with wind input, it ensures fetch unlimited wave growth, decay and transformation of
 242 wind sea and swells. The number of frequencies used for the model were 25 with $f_{min}=0.04$ Hz.
 243 The frequency factor was 1.1 and a logarithmic distribution of frequencies was generated. The
 244 directional discretisation had 24 directional bins, each with 15° resolution, with 360 degree wave
 245 coverage.

246
 247 No current, ice coverage and diffraction were included into the model as this would further
 248 increase computational efforts and also as it can be seen later that without including these
 249 additional inputs, a successful calibration was achieved. Dissipation due to whitecapping, bottom
 250 friction and depth-induced wave breaking were considered in the simulations and the energy
 251 transfer was activated. A quadruplet-wave interaction has been applied. A low order fast
 252 algorithm has been chosen as the solution technique with the ‘maximum number of levels in
 253 transport calculation’ as 32. The source function describing the dissipation due to white-capping
 254 was based on the theory of Hasselmann [29] and Komen et al., [21]. The values applied for $Cdis$
 255 and $DELTA_{dis}$ (δ) were 2 and 0.8 respectively, which were also incidentally found to be close to
 256 the values suggested by Bidlot et al., [30], who revised the whitecapping formulation proposed
 257 by Komen et al., [21], for combined wind sea and swell generation conditions. Bottom friction

258 was considered according to *Nikuradse roughness* (Weber, [31]) and the value applied was 0.04
259 m. The formulation for wave breaking was based on breaking model specified gamma (Nelson,
260 [32-33], Ruessink et al., [34]). The *gamma* and *alpha* values were applied as 0.8 and 1
261 respectively. For detailed description of the above source terms refer to MIKE21 SW manual
262 [20]. The process of selection of model parameters values is further discussed in section 5.1
263 below.

264
265 The integral wave parameters such as significant wave height, peak wave period, mean wave
266 period, energy period, peak wave direction, mean wave direction, directional standard deviation,
267 and wave power have been resolved for every 30 minutes as point series and for every 6 hours as
268 area (contour) series.

269 270 **4. Wave data sources for model calibration and validation**

271 Measured wave data from wave buoys deployed around Scotland have been utilised for model's
272 calibration and validation. Details of their names, locations and duration of the data are listed in
273 Table 1. The locations of the buoys are shown in Figure 4. The Cefas, Blackstone, Moray Firth
274 and Firth of Forth buoys data are in public domain from the WaveNet [35]. The Bragar buoy has
275 been deployed for the Hebridean Marine Energy Futures project (Vogler and Venugopal, [36])
276 and the data are not in public domain yet; for this reason though this wave data were used for
277 model calibration and validation, the results discussed in section 5 will not include Bragar data.
278 The time series of significant wave height (H_{m0}), peak wave period (T_P) and peak wave direction
279 (Dir_P) were only accessible from the public domain buoy data. While it is possible to resolve
280 various wave parameters including wave power from MIKE21 model, only the above wave
281 parameters available from wave buoys have been selected for calibration and validation.

282

283 **5. Results and Discussion**

284 **5.1 Calibration of the wave model**

285 The wave conditions hindcast for May 2012 for Cefas, Blackstone, Moray Firth and Firth of
286 Forth are shown in Figures 5-8 respectively. All four wave measurement's summary statistics
287 have been stored at 30 minutes interval by Wavenet [35] and hence the model output parameters
288 have also been extracted at corresponding time stamps. Note that the wave buoy denoted as
289 Orkney-E in Figure 4 is also privately owned and the data was not accessible at the calibration
290 stage. Usually at the calibration stage, the primary task is to match the model output parameters
291 with measured wave parameters, by tuning those model input parameters that account for source
292 functions given in Eqn. (7). The input parameters which may need tuning include bottom
293 friction, wave breaking parameters, whitecapping, wind, current and water level data, mesh
294 resolution and input boundary. For the current work, during the model calibration stage, initially
295 it was decided to carry out the hindcasting with model's default values describing whitecapping,
296 bottom friction and wave breaking, and this has produced values significantly different from the
297 measurements. As there is no single methodology exists in selecting a set of optimised values of
298 the tuneable input parameters to describe the relevant physical processes, it was decided to
299 attempt a trial and error approach. This involved running the model for a number of cases with
300 various combinations of parameter values, yet keeping the values within the range recommended
301 in the literature as provided in [20]. As an example, when the whitecapping coefficient, C_{dis} ,
302 which control the rate of white-cap dissipation, was changed to 3.0 from its default value of 4.5
303 and another control parameter δ was kept at its default value of 0.5, and at the same time
304 keeping the wave breaking parameter $\gamma = 0.8$ and $\alpha = 1.0$ and bottom friction (represented by the
305 Nikuradse roughness), $kn = 0.04$ m, has produced significantly larger wave heights than

306 measurements. The same combination but with $Cdis = 1.5$ has produced small wave heights than
307 measurements, and at the same time, hindcast peak wave periods were slightly higher than
308 measurements. Finally, the values mentioned above in section 3.2 were found to be producing an
309 overall good agreement with measured H_{m0} , T_P and Dir_P which were finally adopted for further
310 calibration and validation.

311

312 As illustrated in Figures 5-8, the comparison of hindcast significant wave height from model has
313 resulted in an excellent agreement with measured significant wave heights for all four sites. Also,
314 with the exception of few time periods, in general, a very good comparison was found for peak
315 wave periods and peak wave directions for all four sites. This is a known fact that trying to
316 correlate peak wave periods or peak wave directions from model with measurements may
317 produce significant discrepancies in contrast to mean wave periods and mean wave directions.
318 However, these latter two parameters were not available from none of the buoy measurements to
319 compare with. Also it is worth noting that there were significant differences between the model
320 and measurements for up to the first 3 days for all sites, and this was due to the model initiating
321 from a cold start where a fully developed sea condition might not have yet been reached.
322 Nevertheless, referring to Table 2, the quality parameters calculated for four sites [eg., for
323 significant wave height: Bias is in the range -0.09 to +0.11 m, Root Mean Square Error, RMSE is
324 in the range 0.23 – 0.40 m, Scatter Index in the range 0.21 – 0.28), for peak wave period (Bias: -
325 0.38 to +0.29 s, RMSE: 1.62 – 2.76 s, Scatter Index: 0.17 – 0.39)] clearly illustrates that the
326 hindcast model performed well. The definitions for quality indices are given through Eqns (9)-
327 (14). For peak wave direction the quality indices are relatively poor, yet they are considered
328 satisfactory. One must also bear in mind that for the above quality indices calculations, the time
329 series for the whole month was used without avoiding the model initial ramp up period of 3 or 4

330 days, which could have influenced the statistics as well. The Pearson's correlation coefficients, R
 331 - values calculated, for example for wave heights, ranged from 0.88 to 0.97 which indicates the
 332 calibration for significant wave height was highly accurate.

$$333 \quad Bias = \frac{1}{N} \sum_{i=1}^N (x_{o_i} - x_{m_i}) \quad (9)$$

$$334 \quad RMSE = \sqrt{\frac{1}{N} \sum_{i=1}^N (x_{m_i} - x_{o_i})^2} \quad (10)$$

$$335 \quad SI = \frac{RMSE}{\bar{x}_o} \quad (11)$$

$$336 \quad \bar{x}_o = \frac{1}{N} \sum_{i=1}^N (x_{o_i}) \quad (12)$$

$$337 \quad \bar{x}_m = \frac{1}{N} \sum_{i=1}^N (x_{m_i}) \quad (13)$$

$$338 \quad R = \frac{\sum_{i=1}^N (x_{o_i} - \bar{x}_o)(x_{m_i} - \bar{x}_m)}{\sqrt{\sum_{i=1}^N (x_{o_i} - \bar{x}_o)^2 (x_{m_i} - \bar{x}_m)^2}} \quad (14)$$

339 where, x_o is the observed (field) data and x_m is the model data.

340

341 **5.2 Validation of the wave model**

342 With sufficient confidence built up on the calibration of the model as described above, as next
 343 step, an attempt to validate the model for different time periods have been undertaken. Although
 344 the model validation has been performed for different time periods (i.e., Oct-2011, Jan-2012,
 345 March-2012, see Table 1.), considering space limitations the time series for October 2011 are
 346 only shown in Figures 9-12 for four sites. The performance indices (or quality parameters)
 347 calculated for Cefas, Blackstone, Moray Firth and Firth of Forth for Oct-2011, Jan-2012 and

348 March-2012 are listed in Tables 3 – 5 respectively. As indicated by the quality parameters, an
349 excellent agreement has been noticed for significant wave height, however for Moray Firth, some
350 discrepancies in peak wave period and direction are seen. Also for Moray Firth and Firth of
351 Forth, sudden changes in wave directions are distinct and while the model captured this better for
352 Firth of Forth, resolving the same parameter for Moray Firth has not been perfect.

353

354 **5.3 Wave Hindcasting**

355 As previously discussed, the key objective of setting up this model is to assess wave power
356 resources for potential device deployment locations around Orkney and to provide boundary
357 conditions for running regional/small scale wave and tidal flow models within the TeraWatt
358 project. The time duration considered for calibration and validation processes in the above
359 sections were relatively small as only monthly hindcasting have been undertaken. In the
360 assessment of the impact of energy extraction (this could be accomplished by including an
361 individual or array of energy extraction devices into the wave model directly or can be inferred
362 from another hydrodynamic software or CFD methods and eventually be linked with MIKE21)
363 on morphological, ecological and other environmental changes, a longer term data would usually
364 be required. It would then create an interest to learn about how well the wave model is able to
365 hindcast longer term data. With this in mind, the model simulation was carried out for the year
366 2010 for which wave measurement data for Blackstone, Cefas, Orkney-E, Moray Firth and Firth
367 of Forth were available and the results are presented in Figures 13 to 17 respectively. These
368 results generally indicate that the hindcasting of wave parameters, particularly the significant
369 wave height, for all five sites agreed well with the measurements for most of the time periods. As
370 the hindcasting data covered both summer and winter months, it is evident that the model was
371 able to resolve wave conditions for different seasons of a year.

372
373 Inspection of individual locations reveals some interesting features; for Blackstone (Figure 13),
374 the peaks in the significant wave height time series are mostly captured by the model, however, it
375 appears though the highest peak that occurred in November has been under-predicted. A good
376 agreement in peak periods and peak wave directions is encouraging. A similar observation for
377 Cefas location was noticed as seen in Figure 14, however, for a considerable period from mid of
378 September to mid of November, there were no data recorded by the Cefas buoy, yet the model
379 appear to provide the missing wave height, period and direction data and fills this gap in the
380 measurement. The authors believe that this ‘filling data’ could be as accurate as the real measured
381 data as the model data joins and fits in well with measured data for time periods where the buoy
382 data was missing. This is further confirmed by a similar trend in the variation of wave
383 parameters recorded at the Blackstone site in Figure 13, as the two measurement locations are
384 close by and it could be possible that the wave growth and propagation could have related
385 patterns.

386
387 The results presented in Figure 15 are for the Orkney-E location where the wave buoy data for
388 the year 2010 has been provided by the European Marine Energy Centre. Note that the location
389 has not been previously used as a calibration site; nonetheless, the excellent comparison seen in
390 Figure 15 indicates that the model indeed performed well for un-calibrated regions as well.
391 Similar to Cefas location, Orkney-E buoy also had missing data and evidently the model data was
392 found to be filling this gap and linking well with the measured data wherever the data were
393 missing. Another observation in Figure 15 is that the buoy measurements appear to have
394 spurious data, eg., large significant wave height in December and large wave periods of
395 magnitude about 40 seconds in June and December, that are not seen in the model results, which

396 again confirms that the model data could be considered as a substitute for unreliable or erratic
397 measurements.

398
399 The model predictions are compared with measurements for Moray Firth and Firth of Forth in
400 Figures 16 and 17 respectively. Note that these two locations are in the North Sea. For both
401 locations the significant wave height produced a very good correlation with measured data. While
402 the comparison of peak wave period and peak wave direction for Firth of Forth agreed better, for
403 Moray Firth site, some considerable difference in T_p with model data was observed, in
404 comparison to Blackstone, Cefas and Orkney-E sites. It appears that for both these North Sea
405 sites, the significant wave height for majority of the time period in the year 2010 is less than 2 m
406 and yet the model was able to predict this accurately. Looking at the peak periods, although the
407 hindcast values are within the bound of measurements, they are often quite variable. The wave
408 heights appear to be about less than 1m for most of the time for the months from April to August,
409 however the corresponding wave periods varies from about 3 to 12 s indicating that some of the
410 long period waves carried far less energy from North Sea.

411
412 Moreover, a glance at the peak wave direction shows that most of the time the waves have
413 travelled from North East (about 30 deg from due North) to East (90 deg), and as Shetland lies in
414 its path, the island would have acted as a barrier altering the wave propagation, however it is
415 difficult to confirm this without further study. Further, in the peak direction plot, many single
416 vertical lines in the measured data can be seen and these are corresponding to only one single
417 point deviating abruptly from the 'expected' peak wave direction, i.e., a sudden change of wave
418 direction, say from 90 deg to 270 deg and back to 90 deg within 30 minutes is not normal. As this
419 behaviour occurred many times in the measured data, this may make the measured wave direction

420 not credible. The another reason for the discrepancy could be that the model was run without
421 accounting for shallow water triad-wave energy transfer and currents, particularly, the strong tidal
422 currents (see Venugopal and Nimalidinne [27]) that occurs in the Pentland Firth, plus any
423 likelihood occurrence of wave diffraction around the north east tip of Scottish mainland, which
424 would have had some impact on the wave propagation and modification; however realising the
425 good correlation with significant wave height, it is difficult to pinpoint the sources responsible at
426 this stage. While a nearly similar observation is noticed for Firth of Forth in Figure 17, however
427 the model agreed relatively better for this location than Moray Firth. Perhaps, if one trusts that
428 the Shetland indeed obstructed the wave, it may then be reasonable to believe that the Firth of
429 Forth was less influenced by this island, as this is not directly in the downstream when waves
430 propagate in the 30-90 deg sector.

431
432 The another way of inspecting the model results is to represent the data as scatter plots as shown
433 in Figures 18-20 for three locations. The results are also listed in Table 5 as quality index
434 parameters for all five sites under investigation. For all three sites (Figures 18-20), the data for
435 significant wave height from model and measurement are found to be in close proximity to or on
436 the equality line illustrating that the significant wave heights were highly accurately resolved,
437 which is also indicated by values of low Bias (-0.10 to +0.27 m), low RMSE (0.25 to 0.45m), low
438 Scatter Index (0.19 to 0.3) and very high correlation coefficient above 0.94 in Table 5. For peak
439 wave period, except the Moray Firth site, the correlation coefficient is found to be above 0.64 and
440 low values of Bias, RMSE and Scatter index are obtained. For peak wave direction, the
441 correlation is lower than wave height and wave period, however their correlation coefficients,
442 except Moray Firth, are found to be above 0.57. This low value can be explained by re-visiting
443 correlation plots for peak wave direction, in that, the large scatter is attributed to the way in

444 which the wave direction is represented; for example, a measured direction of 0.0 deg and model
445 produced direction of 360 deg are literally the same, however when this is represented as a scatter
446 plot they would produce 'zero' correlation and the same applicable to any other values close to
447 0.0 deg or 360 deg, thus yielding a large scatter and incorrect correlation values.

448
449 In addition to the above sites, three relatively shallow water locations have also been considered
450 for inspection of wave conditions produced by the MIKE21 model, and the results are shown in
451 Figures 21 and 22, for significant wave heights and mean wave directions respectively for the
452 year 2010. Noting that these shallow water locations have been randomly selected at which no
453 measured data were available to compare with the model outputs, it was decided to use another
454 numerical wave model's results for verification. In order to perform this, wave data have been
455 downloaded from the ECMWF wave model archives which were produced using the WAM
456 model [18]. The downside of it is, the type of access the authors have with the ECMWF wave
457 data, allows only to download gridded data stored at a minimum grid spacing of 0.125 deg x
458 0.125 deg resolution, which pose a problem when matching with a chosen shallow water location
459 to the ECMWF grids; however, the authors have made an attempt and selected three locations
460 where the water depth was shallow and the ECMWF data was available. These are denoted as
461 Isle of Lewis (58.375°N, 6.625°W, water depth, $d = 16.6$ m), Westray (59.25°N, 3.0°W, $d = 13.75$
462 m) and Dornoch (57.875°N, 3.875°W, $d = 11.0$ m) in Figure 4.

463
464 Note that in Figures 21 and 22, it was not possible to compare wave periods, as the wave period
465 stored in the present wave model (i.e., mean wave period) was different from the wave period
466 (i.e., energy wave period) available for download from the ECMWF. Considering the time it will
467 consume to re-run the present model to produce energy wave period, this idea was not pursued.

468
469 It is evident that for Isle of Lewis (correlation coefficient $R = 0.95$, $SI = 0.19$ and $Bias = 0.13$)
470 and Dornoch ($R = 0.9$, $SI = 0.45$, $Bias = 0.17$) sites, the significant wave height obtained from the
471 present model matched very well with ECWMF model, however, for the Westray site some
472 significant differences ($R = 0.53$, $SI = 0.57$ and $Bias = -0.27$) were noticed. In addition, for the
473 Westray site, the value of the significant wave height obtained from the ECMWF model in
474 December is over 9 m, which makes one to wonder the likelihood of such a large magnitude
475 wave events to occur in a 13.75 m water depth!

476
477 In the case of mean wave direction (Figure 22), it appears that the MIKE21 model provides a
478 consistent (refer to Figure 15 for Orkney site), less scattered values throughout the time period
479 considered, whereas the WAM model's wave direction often rapidly changes its course.
480 Considering the fact that these results are both from numerical models, it would be difficult to
481 take side as to which model is accurate for these shallow water locations; nevertheless based on
482 the results obtained, the MIKE21 model, can be applied for resource assessment in shallow water
483 conditions.

484

485 **5.4 Comparison of significant wave height and wave power with UK Marine Atlas**

486 In the above sections, wave model calibration, validation and hindcasting have been presented for
487 single point locations, and in this section, in particular, the significant wave height and wave
488 power are presented as contour maps for a region that comprise the boundaries roughly
489 representing the Scottish waters. Figures 23 and 24 show the contour maps of the statistical mean
490 and maximum significant wave height derived for the whole year 2010 and these plots illustrate
491 the spatial variation of significant wave height for different locations. Referring to Figure 1, for

492 the locations west of Orkney mainland where wave device deployment activities are planned
493 (noted by yellow colour rectangular boxes), the annual mean significant wave height (refer
494 Figure 23) is observed to be about 1.75 to 2.0 m and the maximum significant wave height is
495 found to be about 7 to 8 m. Further, it is encouraging to note that the annual mean significant
496 wave height reported from the ‘enhanced model’ (ABPMer Report, [37], see Figure 25) in the
497 Atlas of UK Marine Renewable Energy Resources (ABPMer, [19]) for the same location is found
498 to be 2.01 to 2.25 m which is very close to the one year average found from the present work.
499 While the UK Marine Atlas was calculated based on a hindcasting of the average of 7 years data
500 (1 June 2000 to 31 May 2007), the enhanced model, which mainly covered Orkney and Pentland
501 Firth (region marked by thick black line in Figure 25), results were based on a 20 year (1 Jan
502 1990 to 31 Dec 2009) hindcast. Also the Figure 25 itself an obvious explanation of the need for
503 setting up another refined wave model, such as the present work, as in the UK Marine Atlas, the
504 values appeared to be based on a course mesh (for regions other than the enhanced model), in
505 which the variation of wave heights are represented by rectangles of constant values for a large
506 area which may not be realistic.

507
508 The wave energy flux or wave power (P) in a sea state transported at any water depth can be
509 calculated as

$$510 \quad P = \rho g \int_0^{2\pi} \int_0^{\infty} S(f, \theta) \bar{C}_g(f, \theta) df d\theta \quad (15)$$

511 where, $S(f, \theta)$ is the directional energy spectral density at frequency f and wave propagation
512 angle θ , and $\bar{C}_g(f, \theta)$ is the resultant wave group velocity ([20]).

513

514 The wave power calculated using Eqn (15) at point locations corresponding to Cefas, Blackstone,
515 Orkney, Moray Firth and Firth of Forth are shown in Figures 26(a) and(b) for the year 2010. In
516 Figure 26(b), the power scale axis has been limited at 50 kW/m for Cefas, Blackstone, Orkney
517 sites and 20 kW/m for Moray Firth and Firth of Forth sites, for better visualisation at these
518 limited levels. This figure demonstrates, as expected, that the wave power during the winter
519 months is very high, reaching over 600 kW/m for Cefas and Blackstone sites. For the Orkney-E
520 site, a value of about 300 kW/m is obtained for Dec 2010. Lower values of wave power are
521 observed for Firth of Forth and Moray Firth, making these sites not a candidate for wave power
522 developments.

523
524 The statistical mean and maximum wave power calculated for the same region as in Figures 23
525 and 24, are shown in Figures 27 and 28 as contour plots. From [37], the contour lines (not shown
526 here) representing the annual mean wave power for Orkney wave power strategic regions,
527 indicate values from 30-40 kW/m which was based on 20 year hindcast by the enhanced model as
528 mentioned above. This value is however, comparable to the one obtained from the present model
529 which has produced a value in the range about 25 to 35 kW/m, thus increasing the confidence in
530 using the present model for wave power calculations.

531
532 Additionally, the wave power rose diagram plotted in Figure 29, depicts the proportion of the
533 wave power with respect to wave propagation direction for the year 2010. The data have been
534 worked out using wave power computed for every 30 min blocks for the whole year. Each
535 division of the x and y axis represents a fraction of 5% level. The circle marked with white colour
536 indicates the fraction of the wave power resource which is less than 5 kW/m. It is clear from this
537 plot that for Moray Firth and Firth of Forth about 70 to 73% of wave power is less than 5 kW/m,

538 whereas, for the Orkney-E site only 22.5% are found to be less than 5 kW/m and the most
539 probable wave direction appears to be from due West. For the Blackstone and Cefas sites, wave
540 power values over 105 kW/m have been hindcast and the majority of the waves appear to be
541 propagating from South-West direction, and only about 3 to 9 % of wave power are less than 5
542 kW/m.

543
544 It becomes clear that the model hindcast needs to be carried out for longer time period as done in
545 [37], if the aim is to estimate statistically consistent wave resources. Considering the limited
546 computational resources, it was not possible to execute the model for such a long period of
547 hindcasting for the present work, however, having built up the confidence, the future work will
548 include extended periods with the inclusion of tidal currents and other relevant hydrodynamic
549 processes. Despite the wave model results are based on one year hindcast, it is evident from the
550 plots, tables and arguments presented above that the model performed well and could be adopted
551 for reliable hindcasting and even forecasting of wave conditions and wave power for regions in
552 question.

553

554 **6. Conclusions**

555 A large scale wave model, comprising North Atlantic Ocean bounded by latitudes 10° N - 70° N
556 and longitudes 10° E-75° W, has been developed using the state-of-art MIKE21 suite for
557 hindcasting of wave parameters and wave power. The model included finer scale bathymetry and
558 grid resolutions around Scotland, specifically to the Orkney and Pentland Waters, where wave
559 and tidal energy device deployment activities are consented by the Crown Estate. The model was
560 forced by the wind data obtained from European Centre for Medium Range Weather Forecasts
561 (EWCMMWF) at 0.125 deg resolution. The methodology behind the processing of bathymetry,

562 mesh construction and selection of model input parameters which account for source terms and
563 energy transfer have been described. A comprehensive model calibration and validation has been
564 conducted for four sites, Cefas and Blackstone in the North Atlantic Ocean, and, Moray Firth and
565 Firth of Forth in the North Sea. In addition, a one-year hindcasting has been undertaken.

566
567 The wave hindcasting for the year 2010 has successfully reproduced significant wave heights for
568 Cefas, Blackstone, Orkney-E, Moray Firth and Firth of Forth sites with correlation coefficients
569 higher than 0.96. The peak wave periods for Cefas, Blackstone, Orkney-E sites were found to be
570 well in agreement with buoy measurements with correlation coefficients above 0.69, however, for
571 Moray Firth and Firth of Forth significant differences between model and measured values noted
572 by less marked correlation coefficients of 0.39 and 0.64 respectively. The impact of tidal
573 currents, wave diffraction and triad wave interactions have not been considered in the present
574 model, doing so may have improved the results for Moray Firth and Firth of Forth, which
575 however needs further work. The annual mean significant wave height and wave power obtained
576 for Orkney strategic wave power deployment sites based on one-year wave hindcast were found
577 to be close to the values reported in the Atlas of UK Marine Renewable Energy Resources.

578
579 The results of the study illustrated that the wave model could be employed with high level of
580 confidence for wave hindcasting and even forecasting of various wave parameters and wave
581 power, in particular, for Orkney and Pentland Firth waters and Outer Hebrides.

582 **Acknowledgments**

583 The authors are grateful for the financial support of the UK Engineering and Physical Sciences
584 Research Council (EPSRC) through the Terawatt-Large scale Interactive coupled 3D modelling
585 for wave and tidal energy resource and environmental impact consortium. The authors are also

586 grateful to Cefas (UK) for wave buoys data, European Centre for Medium-Range Weather
587 Forecasts (ECMWF) for providing wind data, European Marine Energy Centre (EMEC) for
588 providing wave buoy data for Orkney and for Marine Scotland Sciences for providing
589 bathymetry data.

590

591 **References**

- 592 [1] renews. <http://renews.biz/special-reports/>, 2014; accessed on 22/07/14.
- 593 [2] The Crown Estate. UK wave and tidal key resource areas project. Technical report,
594 [http://www.thecrownestate.co.uk/energy-infrastructure/wave-and-tidal/pentland-firth-and-](http://www.thecrownestate.co.uk/energy-infrastructure/wave-and-tidal/pentland-firth-and-orkney-waters/enabling-actions/projects-and-publications/)
595 [orkney-waters/enabling-actions/projects-and-publications/](http://www.thecrownestate.co.uk/energy-infrastructure/wave-and-tidal/pentland-firth-and-orkney-waters/enabling-actions/projects-and-publications/), Oct 2012; accessed on
596 28/05/14.
- 597 [3] Aquamarine Power. <http://www.aquamarinepower.com/projects>, accessed on 28/05/14.
- 598 [4] Pelamis wave power. [http://www.pelamiswave.com/our-projects/project/2/ScottishPower-](http://www.pelamiswave.com/our-projects/project/2/ScottishPower-Renewables-at-EMEC)
599 [Renewables-at-EMEC](http://www.pelamiswave.com/our-projects/project/2/ScottishPower-Renewables-at-EMEC), accessed on 28/05/14.
- 600 [5] Alstom. <http://www.alstom.com/power/renewables/ocean-energy/>, accessed on 28/05/14.
- 601 [6] Andritz Hydro Hammerfest. <http://www.hammerfeststrom.com/>, accessed on 28/05/14.
- 602 [7] AW Energy. <http://aw-energy.com/>, accessed on 28/05/14.
- 603 [8] Voith Hydro. <http://voith.com/en/products-services/hydro-power/ocean-energies-587.html>,
604 accessed on 28/05/14.
- 605 [9] Wello Oy. <http://www.wello.eu/penguin.php>, accessed on 28/05/14.
- 606 [10] <http://www.emec.org.uk/about-us/our-tidal-clients/>, accessed on 28/05/14.
- 607 [11] <http://www.emec.org.uk/about-us/wave-clients/>, accessed on 28/05/14.
- 608 [12] <http://www.scotland.gov.uk/Topics/marine/marineenergy>, accessed on 28/05/14.
- 609 [13] Swail VR, Cardone VJ, Cox AT. A long term North Atlantic wave hindcast. In: 5th
610 International Workshop on Wave Hindcasting and Forecasting, January 26-30, Melbourne,
611 Florida, 1998.
- 612 [14] Swail VR, Ceccacci EA, Cox AT. The AES40 North Atlantic wave reanalysis: validation
613 and climate assessment. In: 6th International Workshop On Wave Hindcasting and
614 Forecasting, November 6-10, Monterey, California, USA, 2000.

- 615 [15] Dodet G, Bertin X, Taborda, R. Wave climate variability in the North-East Atlantic Ocean
616 over the last six decades. *Ocean Modelling* 2010; 31: 120–131.
- 617 [16] Tolman HL. 2009. User manual and system documentation of WAVEWATCH III version
618 3.14. NOAA/NWS/NCEP/MMAB Technical Note 276, 2009; 194 p.
- 619 [17] Galanis G, Chu PC, Kallos G, Kuo Y-H, Dodson CTJ. Wave height characteristics in the
620 north Atlantic ocean: a new approach based on statistical and geometrical techniques.
621 *Stochastic Environmental Research and Risk Assessment* 2012; 26(1): 83-103.
- 622 [18] WAMDI Group. The WAM model—A third generation ocean wave prediction model. *J.*
623 *Phys. Oceanogr.*, 1988; 18: 1775–1810.
- 624 [19] ABPmer. Atlas of UK Marine Renewable Energy Resources. [http://www.renewables-](http://www.renewables-atlas.info)
625 [atlas.info](http://www.renewables-atlas.info); 2008. Date of access 07 July 2014.
- 626 [20] MIKE 21. Wave Modelling User guide, Scientific Document, Danish Hydraulic Institute,
627 Denmark; 2012.
- 628 [21] Komen GJ, Cavaleri L, Donelan M, Hasselmann K, Hasselmann S, Janssen PAEM.
629 *Dynamics and Modelling of Ocean Waves*. Cambridge Univ. Press, New York; 1994, 532.
- 630 [22] Young IR. Wind generated ocean waves. Elsevier Ocean Engineering Book Series, Vol 2,
631 Eds. Bhattacharyya, R and McCormick, M.E, ISBN: 978-0-08-043317-2; 1999.
- 632 [23] Janssen PAEM. Wave induced stress and the drag of airflow over sea waves, *J. Phys.*
633 *Oceanogr.* 1989; 19: 745-754.
- 634 [24] Janssen PAEM. Quasi-linear theory of wind wave generation applied to wave forecasting, *J.*
635 *Phys. Oceanogr.* 1991; 21: 1631-1642.
- 636 [25] GEBCO. <http://www.gebco.net/> ; accessed on 28/05/14.
- 637 [26] Marine Scotland. [http://www.scotland.gov.uk/Topics/marine/science/](http://www.scotland.gov.uk/Topics/marine/science/MSInteractive/datatype/Bathymetry/data)
638 [MSInteractive/datatype/Bathymetry/data](http://www.scotland.gov.uk/Topics/marine/science/MSInteractive/datatype/Bathymetry/data). accessed on 28/05/14.
- 639 [27] Venugopal V, Nimalidinne R. Marine energy resource assessment for Orkney and Pentland
640 waters with a coupled wave and tidal flow. In: *Proceedings of the 33th International*
641 *Conference on Ocean, Offshore and Arctic Engineering OMAE2014*, June 8-13, 2014, San
642 Francisco, AC, USA, paper no OMAE2014-24027, 2014.
- 643 [28] ECMWF. <http://www.ecmwf.int/>; accessed on 28/05/14.
- 644 [29] Hasselmann K. On the spectral dissipation of ocean waves due to whitecapping, *Bound.*
645 *Layer Meteor.* 1974; 6: 107-127.

- 646 [30] Bidlot J, Janssen P, Abdalla S. A revised formulation of ocean wave dissipation and its
647 model impact. Technical Report Memorandum 509, ECMWF, Reading, UK; 2007.
- 648 [31] Weber SL. Bottom friction for wind sea and swell in extreme depth-limited situations, J.
649 Phys. Oceanogr. 1991; 21: 149-172.
- 650 [32] Nelson RC. Design wave heights on very mild slopes: An experimental study, Civil. Eng.
651 Trans., Inst. Eng. 1987; 29: 157-161.
- 652 [33] Nelson RC. Depth limited wave heights in very flat regions. Coastal Eng. 1994; 23: 43-59.
- 653 [34] Ruessink BG, Walstra DJR, Southgate H.N. Calibration and verification of a parametric
654 wave model on barred beaches. Coastal Eng. 2003; 48: 139-149.
- 655 [35] WaveNet. [http://www.cefas.defra.gov.uk/our-science/observing-and-modelling/monitoring-](http://www.cefas.defra.gov.uk/our-science/observing-and-modelling/monitoring-programmes/wavenet.aspx)
656 [programmes/wavenet.aspx](http://www.cefas.defra.gov.uk/our-science/observing-and-modelling/monitoring-programmes/wavenet.aspx); accessed on 28/05/14.
- 657 [36] Vogler A, Venugopal V. Hebridean marine energy resources: wave power characterisation
658 using a buoy network. In: Proceedings of 31th International Conference on Offshore
659 Mechanics and Arctic Engineering, OMAE2012, Rio de Janeiro, Brazil, OMAE2012-
660 83658, 2012.
- 661 [37] APBMer. Report R.1936, Unpublished, 2012.

662
663
664
665
666

Tables

Table 1. Details of buoy data used for model calibration and validation

Process	Buoys/site name	Latitude	Longitude	Water depth (m)	Time period
Calibration	Cefas	57.292333° N	7.914333° W	100	May 2012
	Blackstone	56.062000° N	7.056833° W	97	May 2012
	Moray Firth	57.966333° N	3.333167° W	54	May 2012
	Firth of Forth	56.188167° N	2.503833° W	65	May 2012
Validation	Cefas	57.292333° N	7.914333° W	100	Oct 2011, Jan 2012,Mar 2012
	Blackstone	56.062000° N	7.056833° W	97	Oct 2011, Jan 2012,Mar 2012
	Orkney E	58.970200° N	3.390900° W	53	-
	Moray Firth	57.966333° N	3.333167° W	54	Oct 2011, Jan 2012,Mar 2012
	Firth of Forth	56.188167° N	2.503833° W	65	Oct 2011, Jan 2012,Mar 2012

Table 2. Quality Indices-May 2012

Site	Wave parameters	Mean	Bias	RMSE	Bias/Mean	SI	R
Cefas	Hm0 (m)	1.88	0.11	0.40	0.06	0.21	0.96
	Tp (s)	9.30	0.01	1.62	0.00	0.17	0.69
	Dirp (deg)	216.77	-33.78	148.58	-0.16	0.69	0.45
Blackstone	Hm0 (m)	1.50	0.05	0.31	0.03	0.21	0.97
	Tp (s)	8.57	0.29	2.30	0.03	0.27	0.60
	Dirp (deg)	291.00	-20.40	60.35	-0.07	0.21	0.27
Moray Forth	Hm0 (m)	0.91	-0.15	0.26	-0.17	0.28	0.88
	Tp (s)	7.08	-0.33	2.76	-0.05	0.39	0.45
	Dirp (deg)	79.86	-15.24	64.46	-0.19	0.81	0.36
Firth of Forth	Hm0 (m)	0.97	-0.09	0.23	-0.09	0.24	0.88
	Tp (s)	6.97	-0.38	2.08	-0.05	0.30	0.40
	Dirp (deg)	71.26	-4.57	33.37	-0.06	0.47	0.78

Table 3. Quality indices for October 2011

Site	Wave parameters	Mean	Bias	RMSE	Bias/Mean	SI	R
Cefas	Hm0 (m)	3.25	0.44	0.79	0.13	0.24	0.90
	Tp (s)	11.32	0.01	1.83	0.00	0.16	0.74
	Dirp (deg)	272.50	-14.06	31.01	-0.05	0.11	0.68
Blackstone	Hm0 (m)	3.25	-0.06	0.46	-0.22	0.14	0.94
	Tp (s)	10.90	0.24	1.88	0.02	0.17	0.72
	Dirp (deg)	266.50	-3.35	24.14	-0.01	0.09	0.73
Moray Forth	Hm0 (m)	1.26	-0.38	0.51	-0.30	0.41	0.91
	Tp (s)	7.59	2.55	5.82	0.34	0.77	0.23
	Dirp (deg)	141.91	-45.58	95.19	-0.32	0.67	0.42
Firth of Forth	Hm0 (m)	1.16	-0.15	0.26	-0.13	0.22	0.97
	Tp (s)	7.12	0.80	3.18	0.11	0.45	0.52
	Dirp (deg)	128.72	-13.06	75.44	-0.10	0.59	0.55

Table 4. Quality Indices-Jan 2012

Site	Wave parameters	Mean	Bias	RMSE	Bias/Mean	SI	R
Cefas	Hm0 (m)	4.90	0.45	0.75	0.09	0.15	0.96
	Tp (s)	12.89	0.28	1.37	0.02	0.11	0.81
	Dirp (deg)	271.66	-8.50	25.73	-0.03	0.09	0.53
Blackstone	Hm0 (m)	4.48	0.06	0.54	0.01	0.12	0.96
	Tp (s)	12.94	0.16	1.36	0.01	0.1	0.77
	Dirp (deg)	277.59	-11.88	24.45	-0.04	0.09	0.47
Moray Forth	Hm0 (m)	1.31	-0.37	0.55	-0.28	0.42	0.67
	Tp (s)	7.49	2.24	5.68	0.30	0.76	0.16
	Dirp (deg)	160.13	-70.40	117.84	-0.44	0.74	0.32
Firth of Forth	Hm0 (m)	1.07	-0.17	0.31	-0.16	0.29	0.86
	Tp (s)	6.50	1.27	3.42	0.20	0.53	0.35
	Dirp (deg)	156.37	61.88	113.81	0.40	0.73	0.27

Table 5. Quality Indices-March 2012

Site	Wave parameters	Mean	Bias	RMSE	Bias/Mean	SI	R
Cefas	Hm0 (m)	3.79	0.36	0.59	0.10	0.16	0.97
	Tp (s)	12.85	0.17	1.42	0.01	0.11	0.82
	Dirp (deg)	269.44	-8.32	19.67	-0.03	0.07	0.71
Blackstone	Hm0 (m)	3.25	0.02	0.47	0.01	0.14	0.95
	Tp (s)	12.75	0.18	1.53	0.01	0.12	0.78
	Dirp (deg)	280.41	-16.62	22.74	-0.06	0.08	0.52
Moray Forth	Hm0 (m)	0.87	-0.38	0.47	-0.44	0.54	0.80
	Tp (s)	8.58	2.45	5.84	0.29	0.68	0.33
	Dirp (deg)	136.92	-56.45	111.21	-0.41	0.81	0.21
Firth of Forth	Hm0 (m)	0.62	-0.16	0.23	-0.25	0.37	0.88
	Tp (s)	6.58	1.66	5.21	0.25	0.79	0.08
	Dirp (deg)	128.08	-27.09	102.51	-0.21	0.80	0.31

Table 6. Quality indices for Jan-Dec 2010

Site	Wave parameters	Mean	Bias	RMSE	Bias/Mean	SI	R
Cefas	Hm0 (m)	2.28	0.27	0.44	0.12	0.19	0.95
	Tp (s)	10.44	0.42	1.71	0.04	0.16	0.72
	Dirp (deg)	283.5	-16.94	40.93	-0.06	0.14	0.59
Blackstone	Hm0 (m)	2.03	0.23	0.45	0.11	0.22	0.94
	Tp (s)	10.15	0.61	2.02	0.06	0.20	0.71
	Dirp(deg)	268.09	-6.85	48.44	-0.03	0.18	0.61
Moray Forth	Hm0 (m)	1.13	-0.16	0.34	-0.14	0.30	0.94
	Tp (s)	7.58	0.11	3.01	0.01	0.40	0.39
	Dirp(deg)	100.96	-36.90	84.78	-0.37	0.84	0.27
Firth of Forth	Hm0 (m)	1.15	-0.10	0.25	-0.09	0.22	0.96
	Tp (s)	7.40	-0.13	1.85	-0.02	0.25	0.64
	Dirp(deg)	86.50	-17.14	55.97	-0.20	0.65	0.57
Orkney	Hm0 (m)	1.67	0.03	0.31	0.02	0.19	0.95
	Tp (s)	10.21	0.45	1.96	0.04	0.19	0.69
	Dirp(deg)	304.41	-6.13	22.23	-0.02	0.07	0.75

Figures

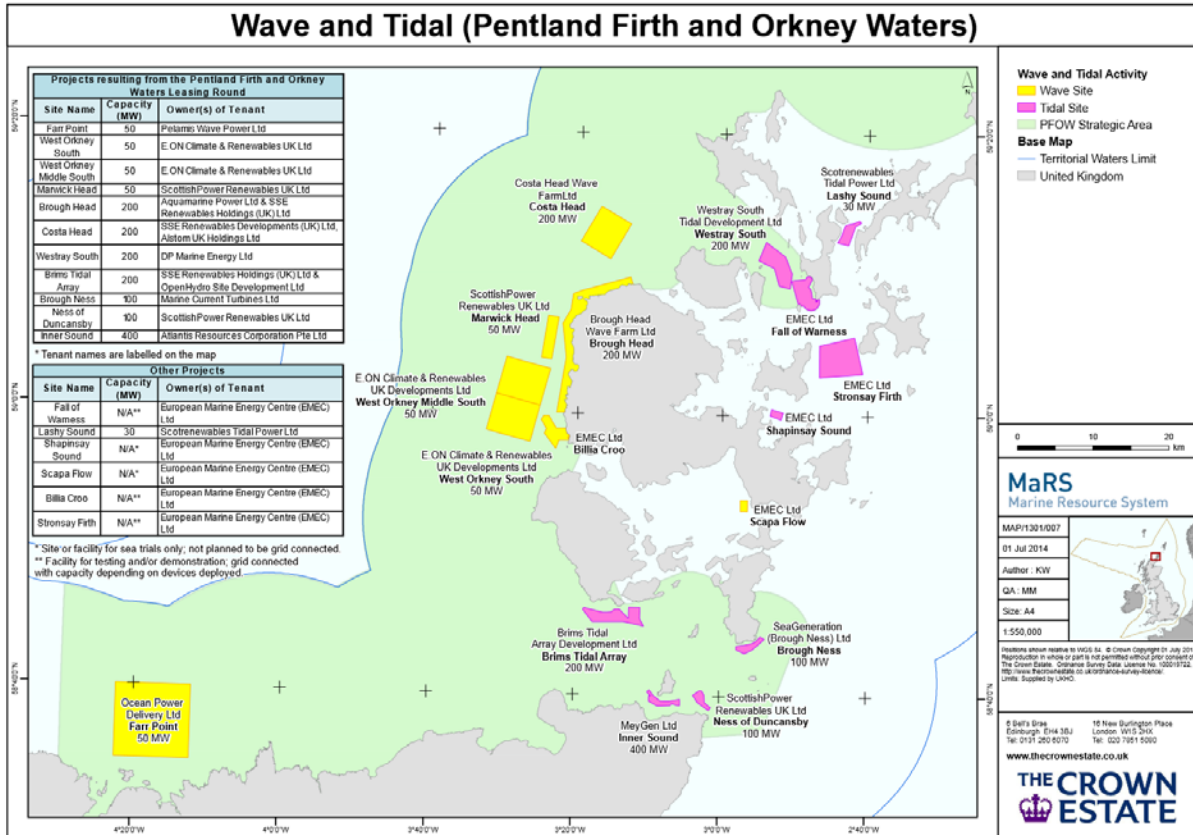


Figure 1. Location of Pentland Firth showing wave and tidal energy leasing sites [2].

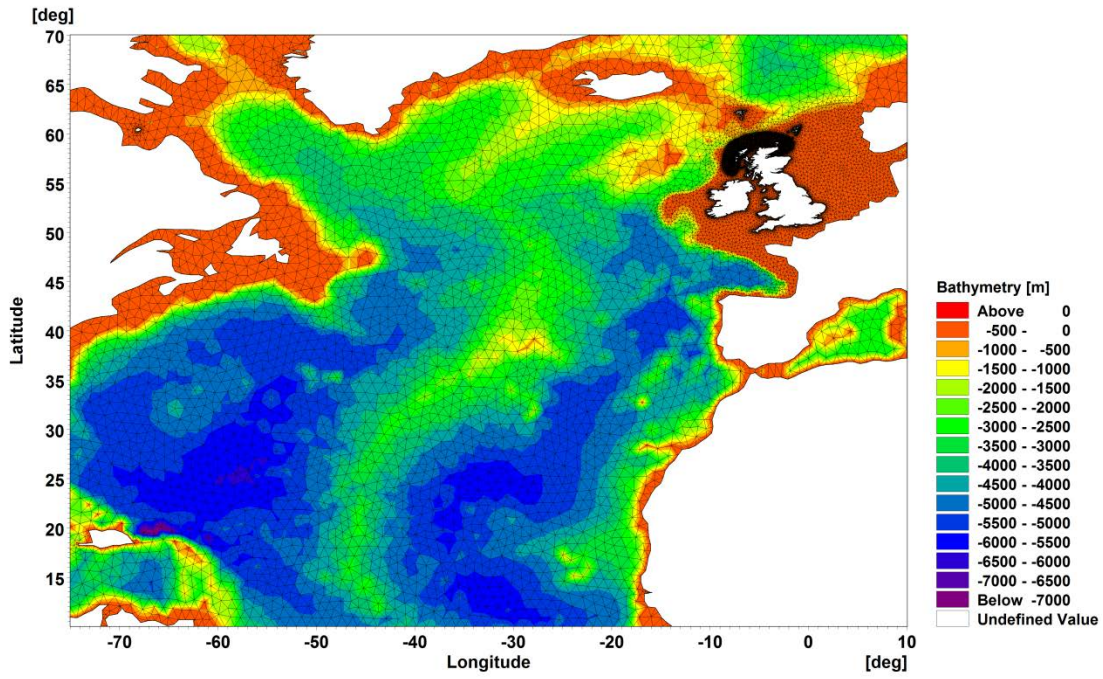


Figure 2. Computational domain for North Atlantic wave model

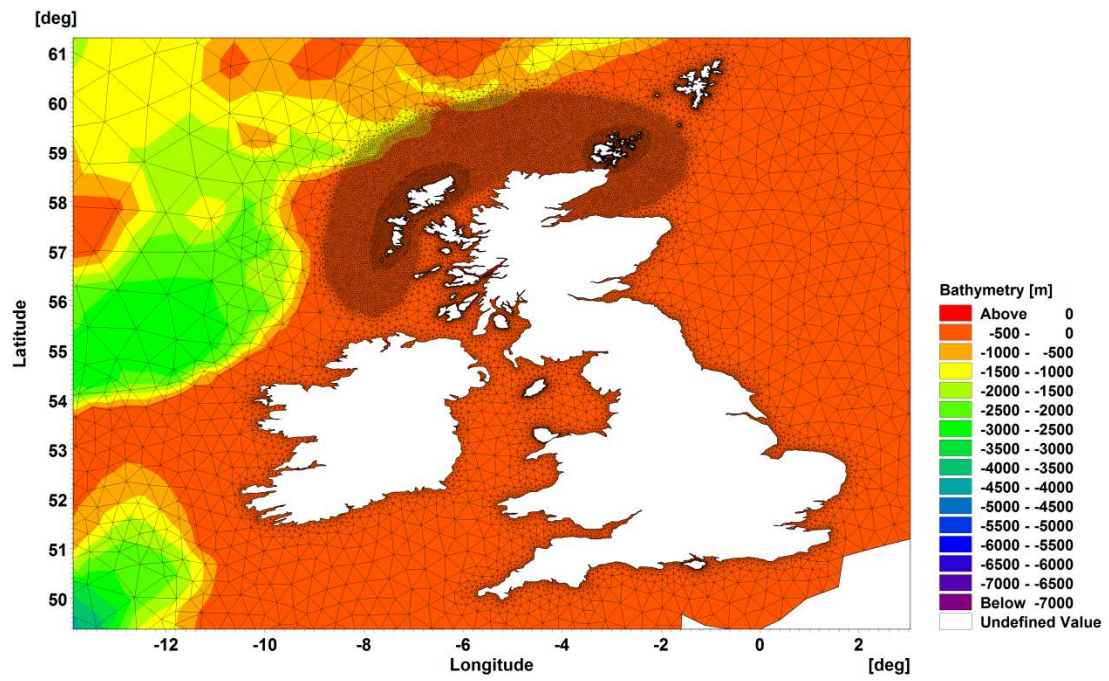


Figure 3. Enlarged view of the computational mesh for UK/Scotland

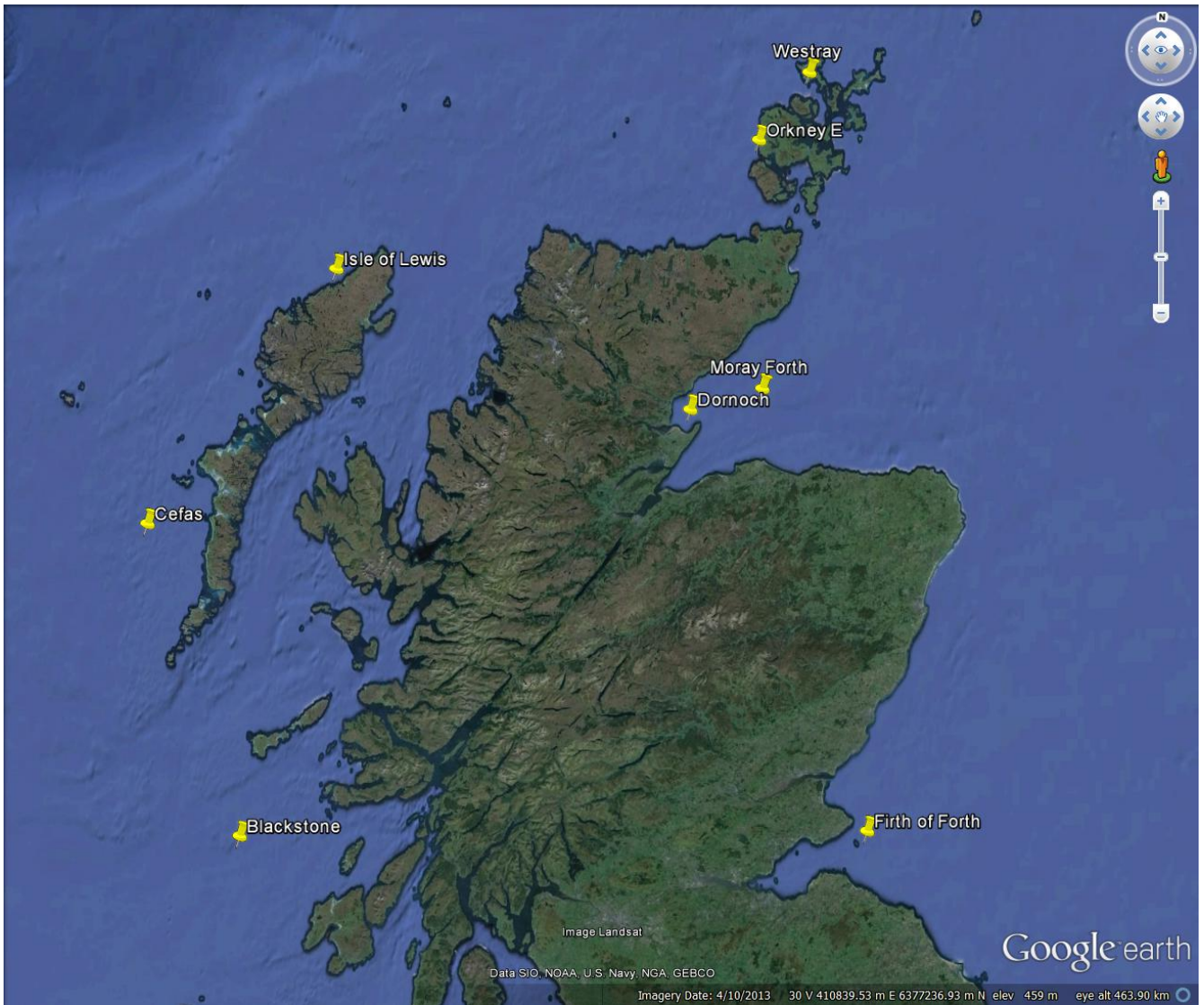


Figure 4. Location of wave buoys in google earth.

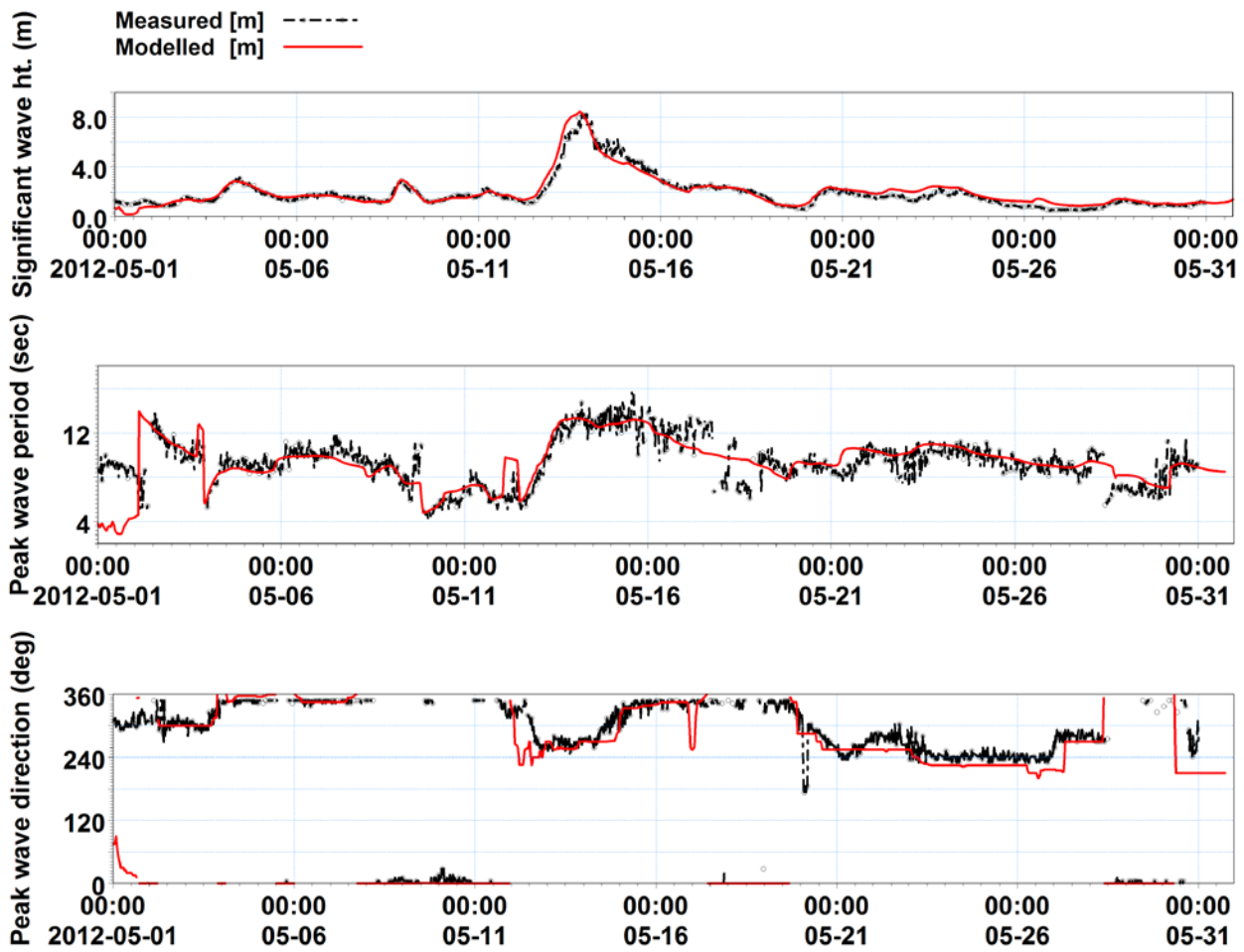


Figure 5. Comparison of significant wave height, peak wave period and peak wave direction between model and Cefas wave buoy, for May 2012; Model calibration phase.

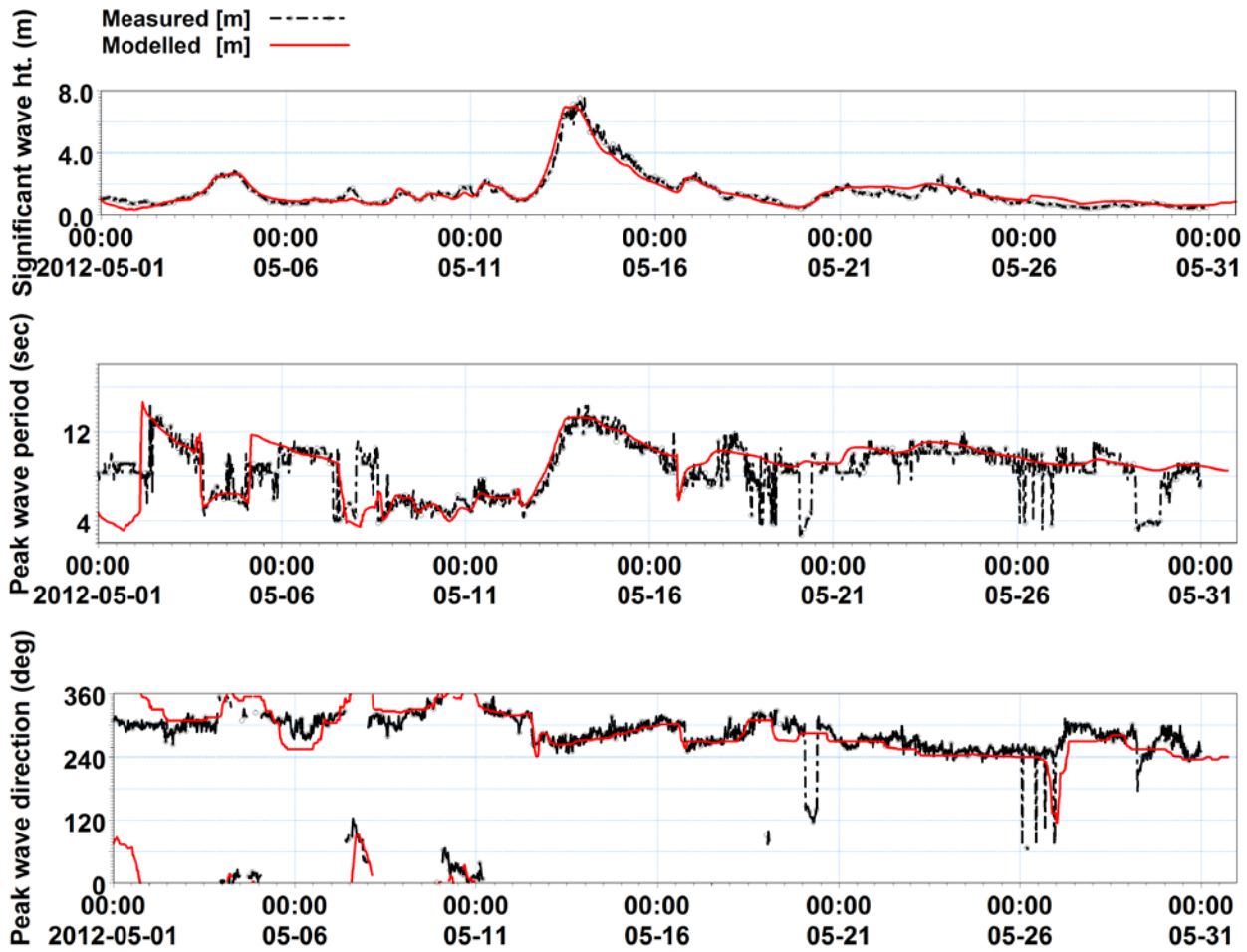


Figure 6. Comparison of significant wave height, peak wave period and peak wave direction between model and Blackstone wave buoy, for May 2012; Model calibration phase.

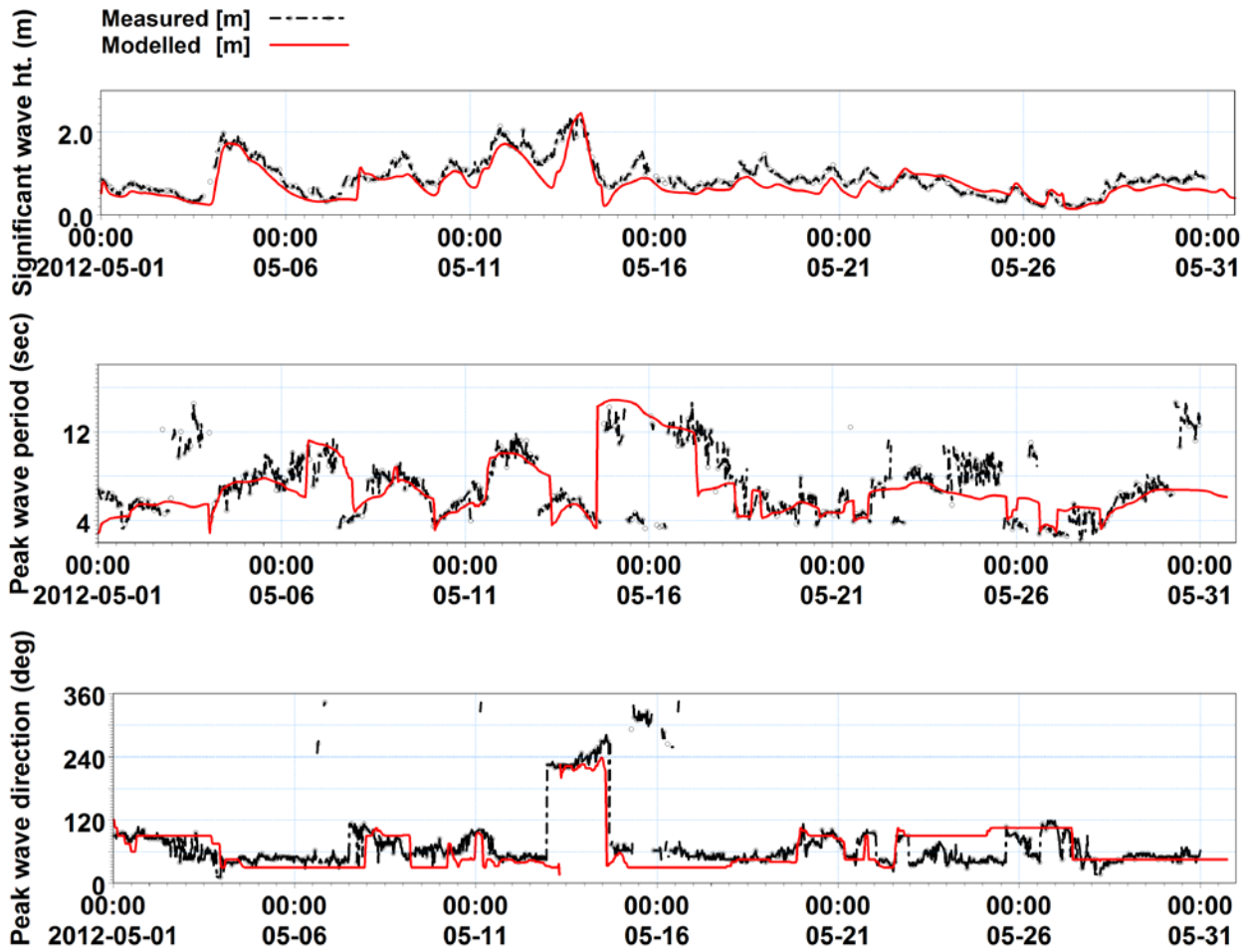


Figure 7. Comparison of significant wave height, peak wave period and peak wave direction between model and Moray Firth wave buoy, for May 2012; Model calibration phase.

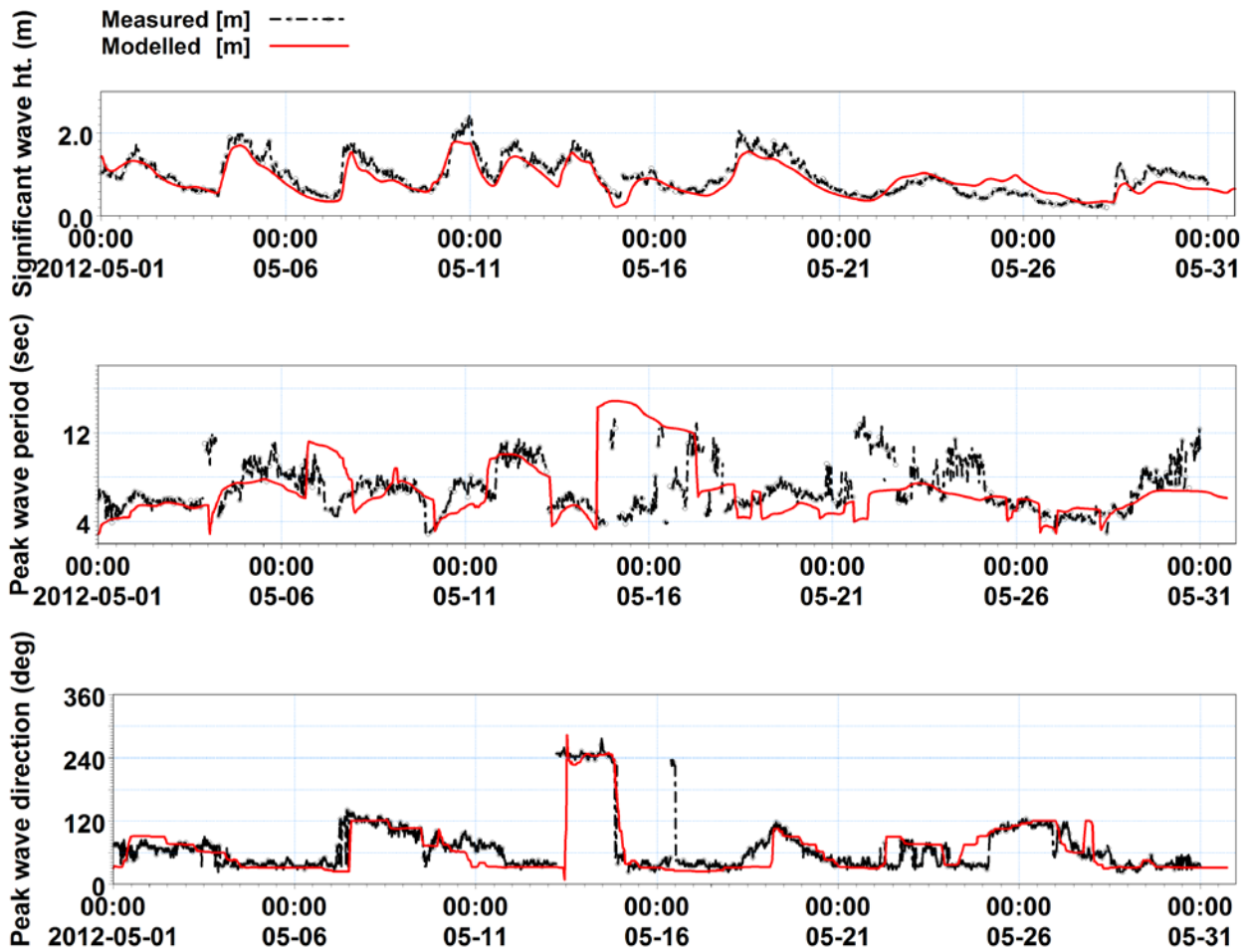


Figure 8. Comparison of significant wave height, peak wave period and peak wave direction between model and Firth of Forth wave buoy, for May 2012; Model calibration phase.

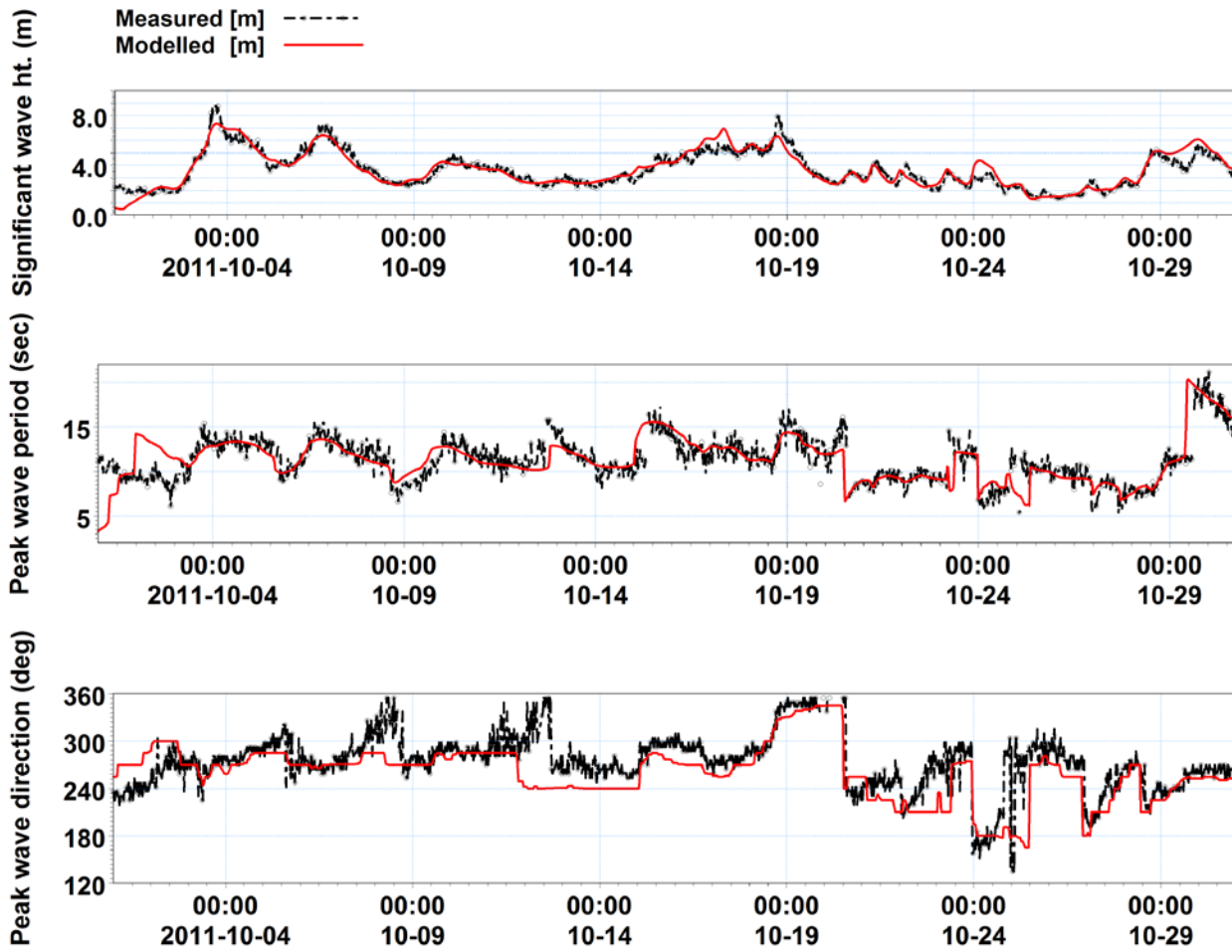


Figure 9. Comparison of significant wave height, peak wave period and peak wave direction between model and Cefas wave buoy, for October 2011; Model validation phase.

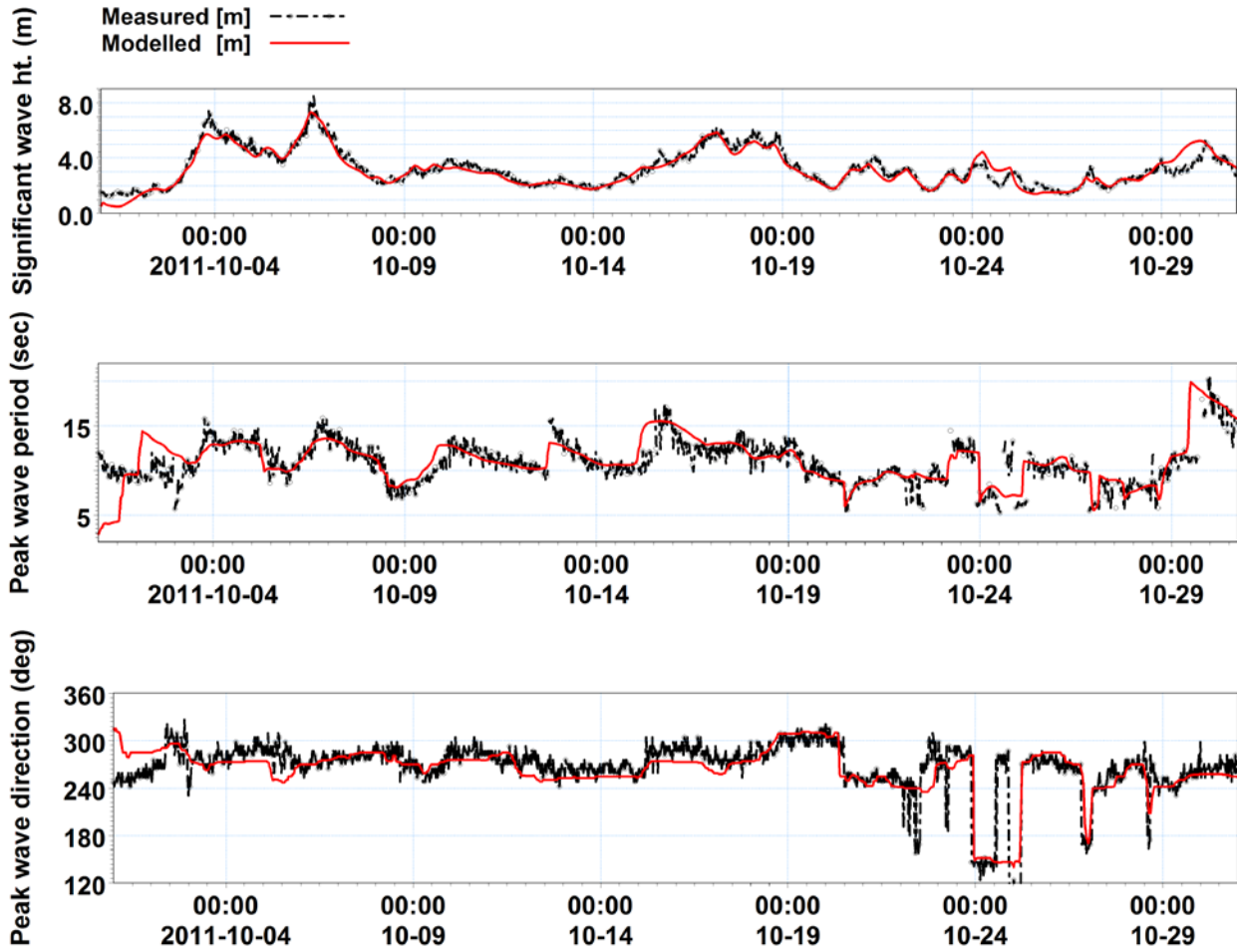


Figure 10. Comparison of significant wave height, peak wave period and peak wave direction between model and Blackstone wave buoy, for October 2011; Model validation phase.

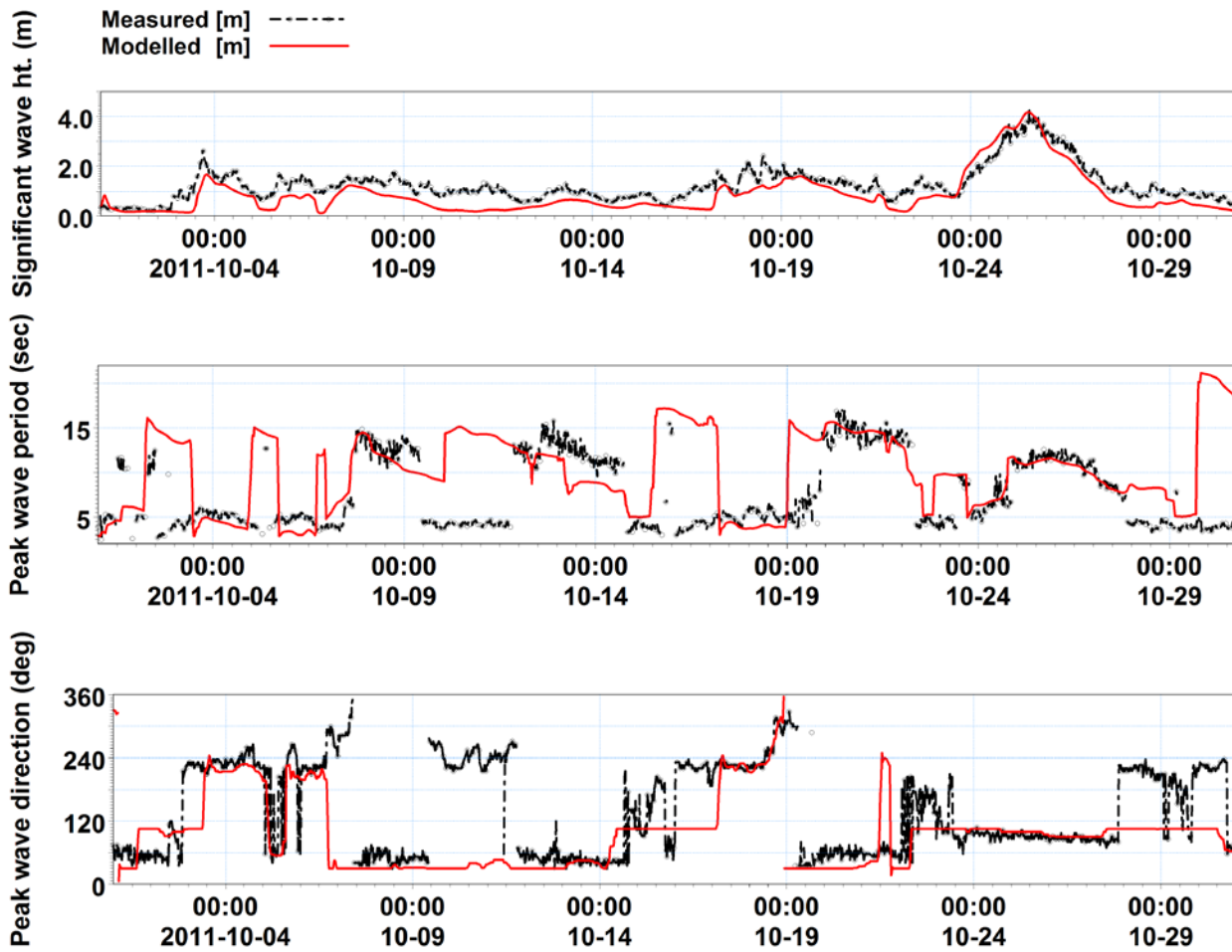


Figure 11. Comparison of significant wave height, peak wave period and peak wave direction between model and Moray Firth wave buoy, for October 2011; Model validation phase.

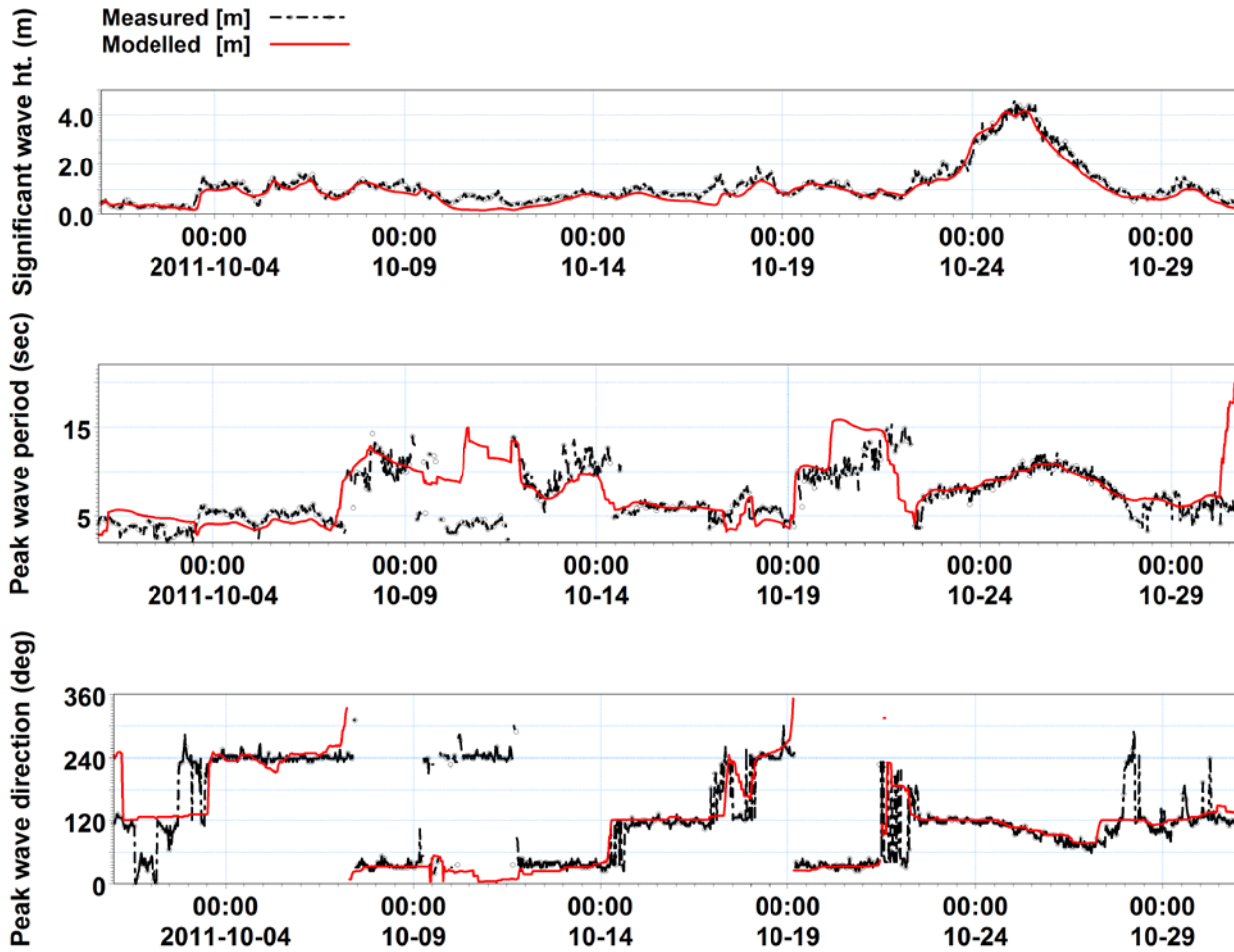


Figure 12. Comparison of significant wave height, peak wave period and peak wave direction between model and Firth of Forth wave buoy, for October 2011; Model validation phase.

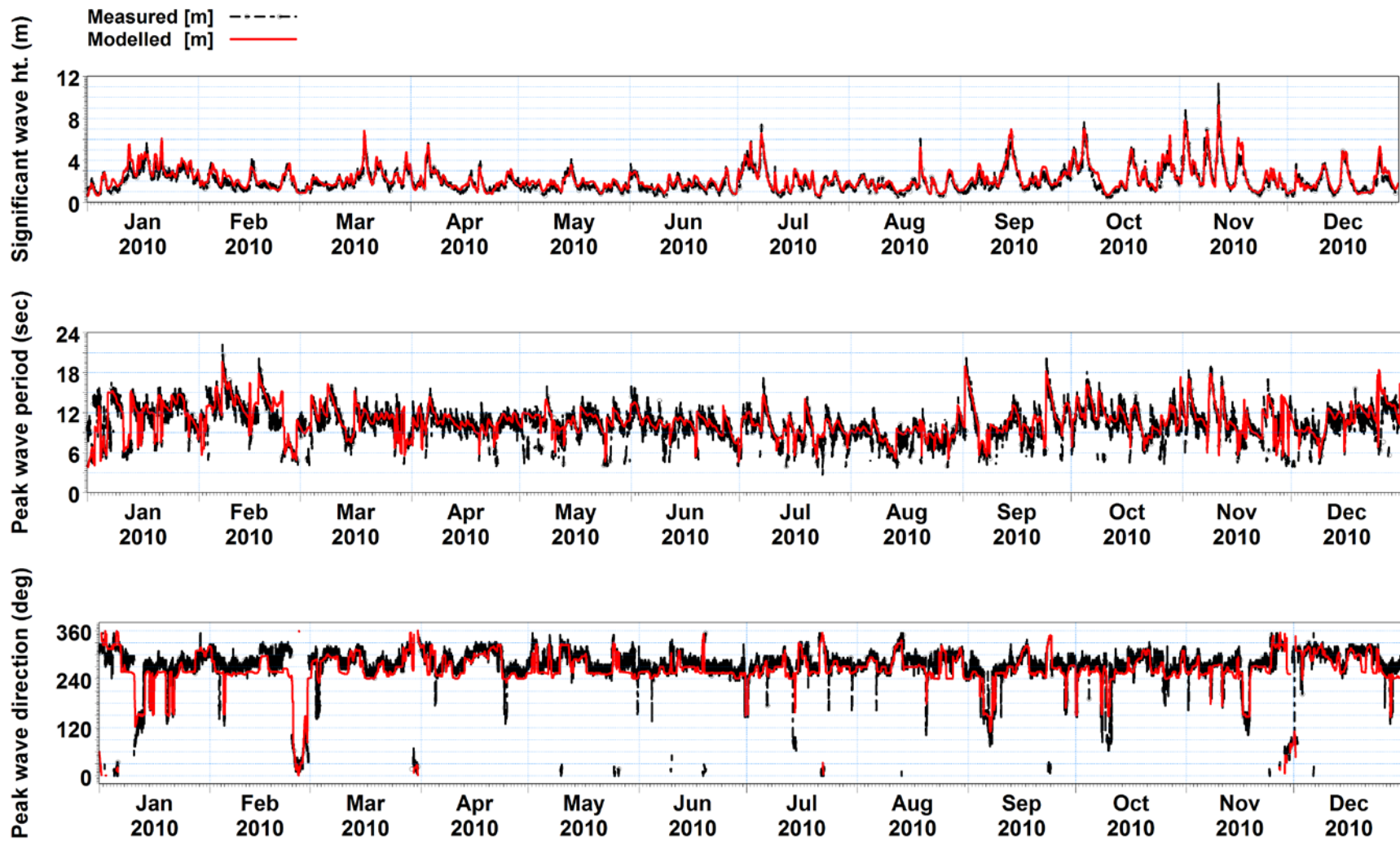


Figure 13. Comparison of significant wave height, peak wave period and peak wave direction between model and Blackstone wave buoy, for January-December 2010.

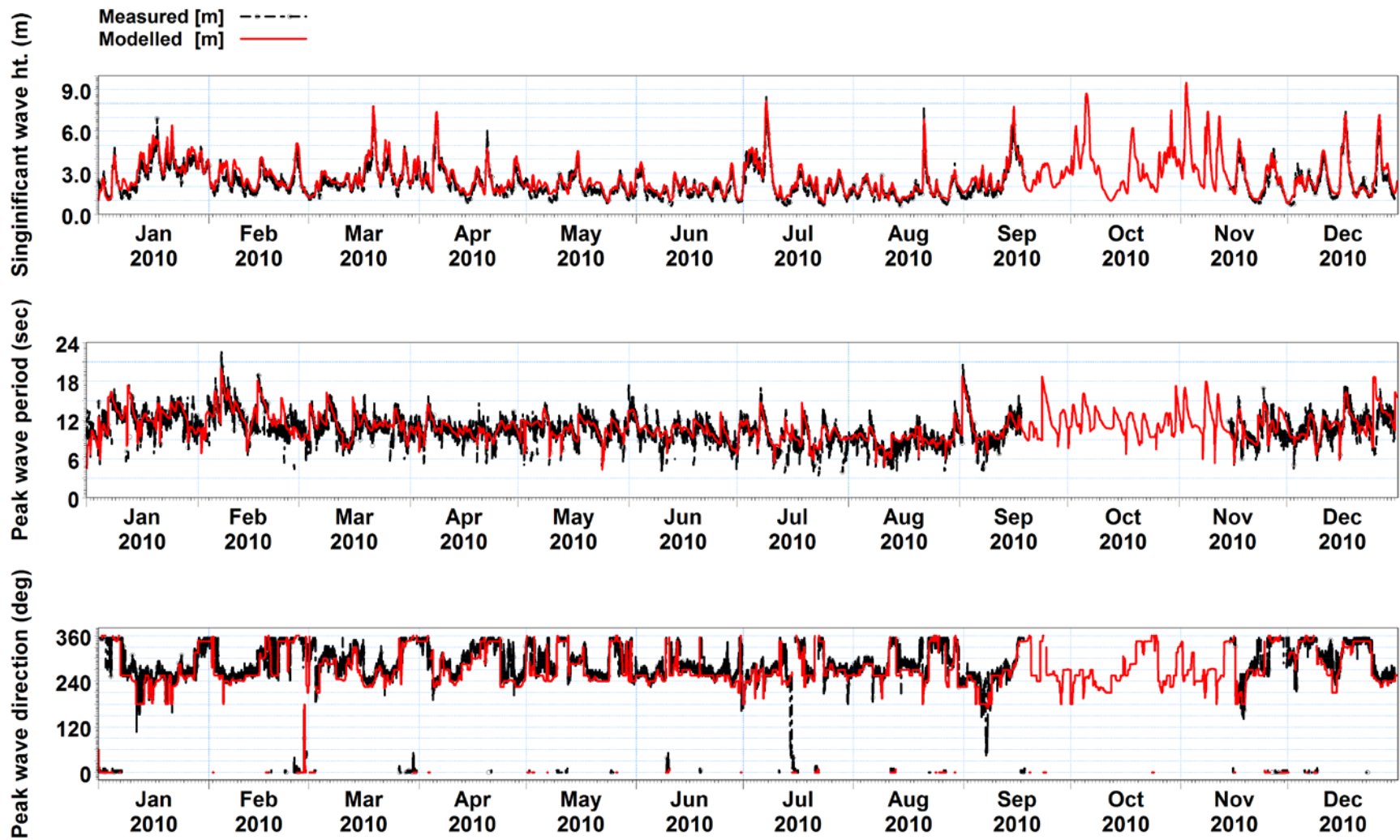


Figure 14. Comparison of significant wave height, peak wave period and peak wave direction between model and Cefas wave buoy, for January-December 2010.

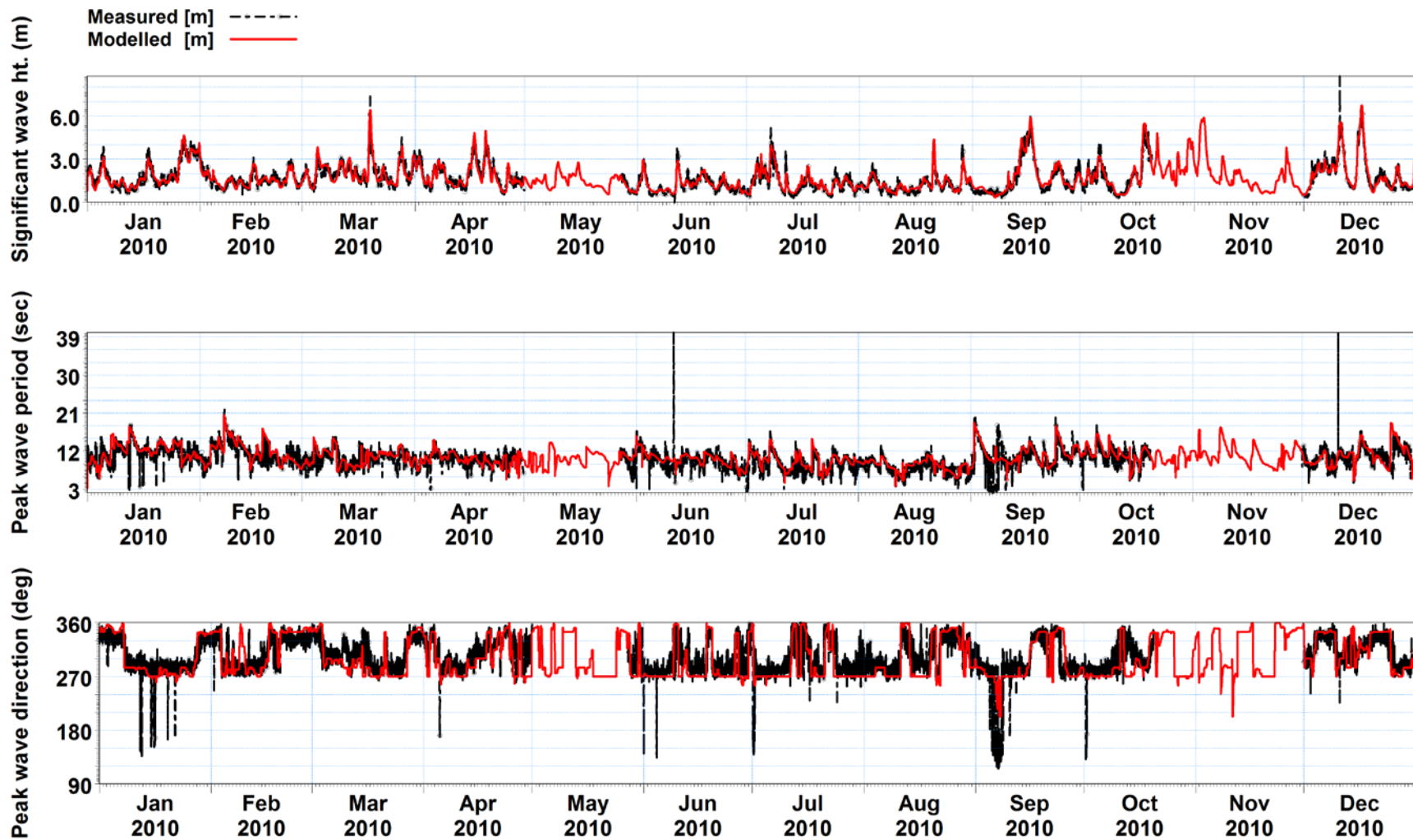


Figure 15. Comparison of significant wave height, peak wave period and peak wave direction between model and Orkney wave buoy, for January-December 2010.

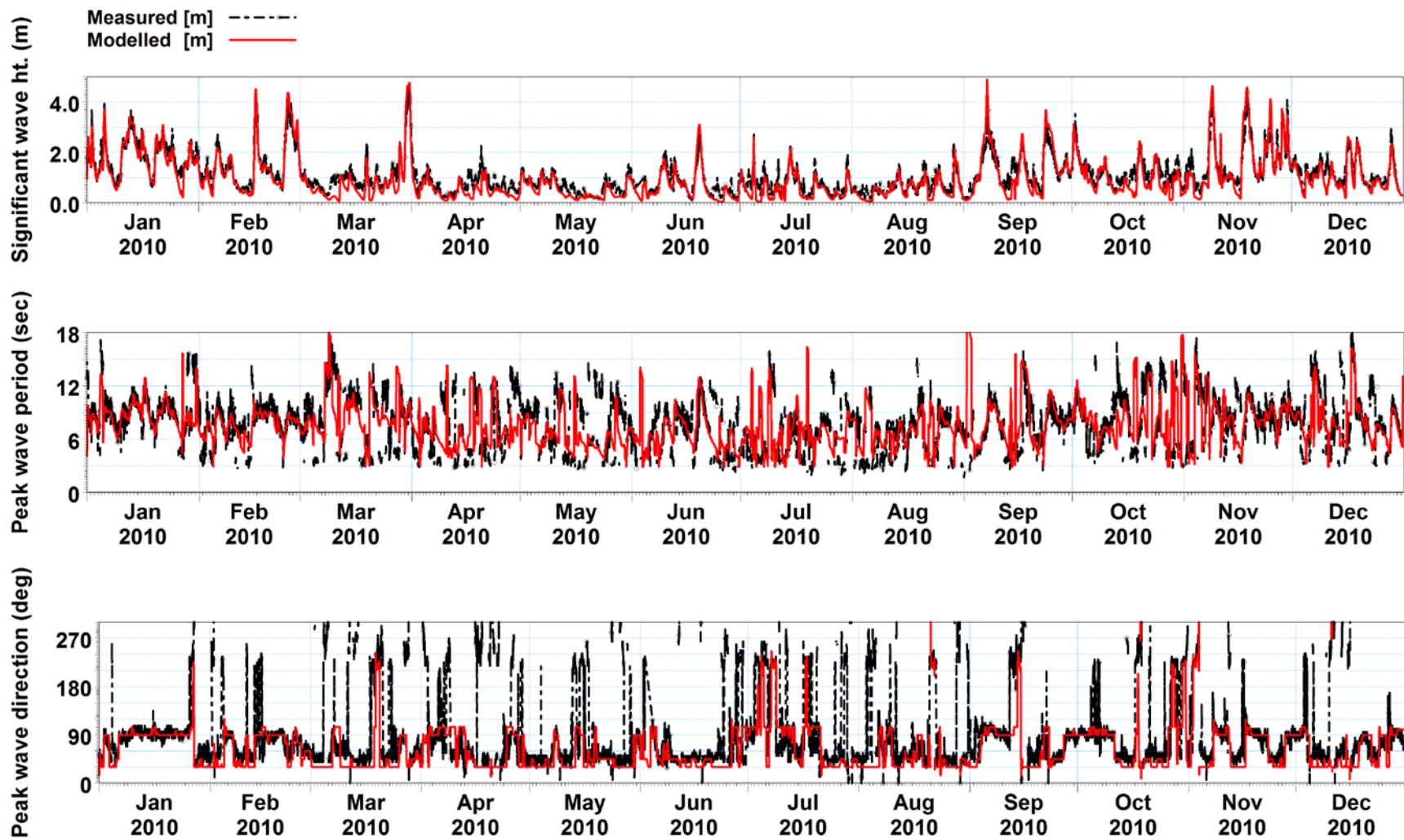


Figure 16. Comparison of significant wave height, peak wave period and peak wave direction between model and Moray Firth wave buoy, for January-December 2010.

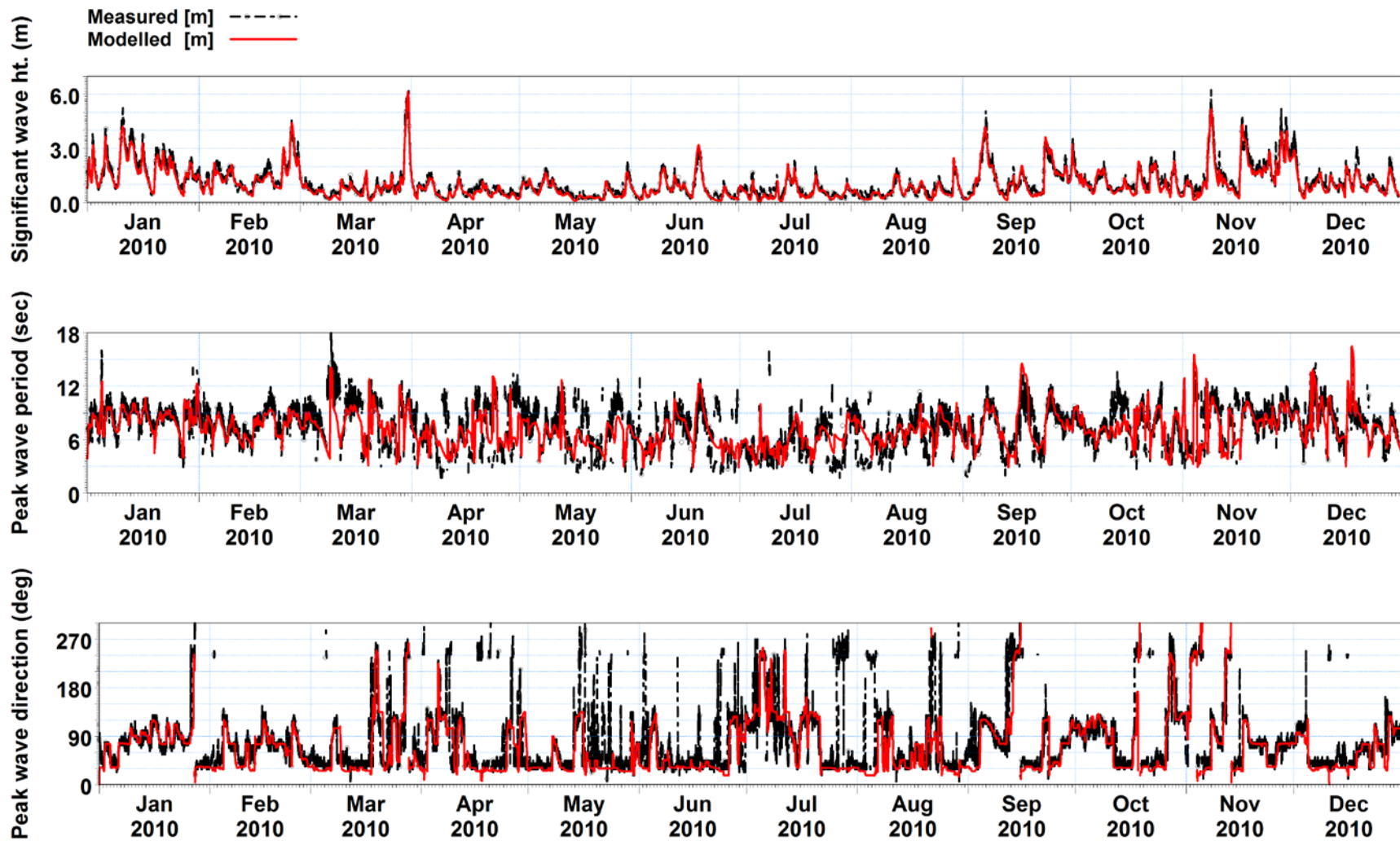


Figure 17. Comparison of significant wave height, peak wave period and peak wave direction between model and Firth of Forth wave buoy, for January-December 2010.

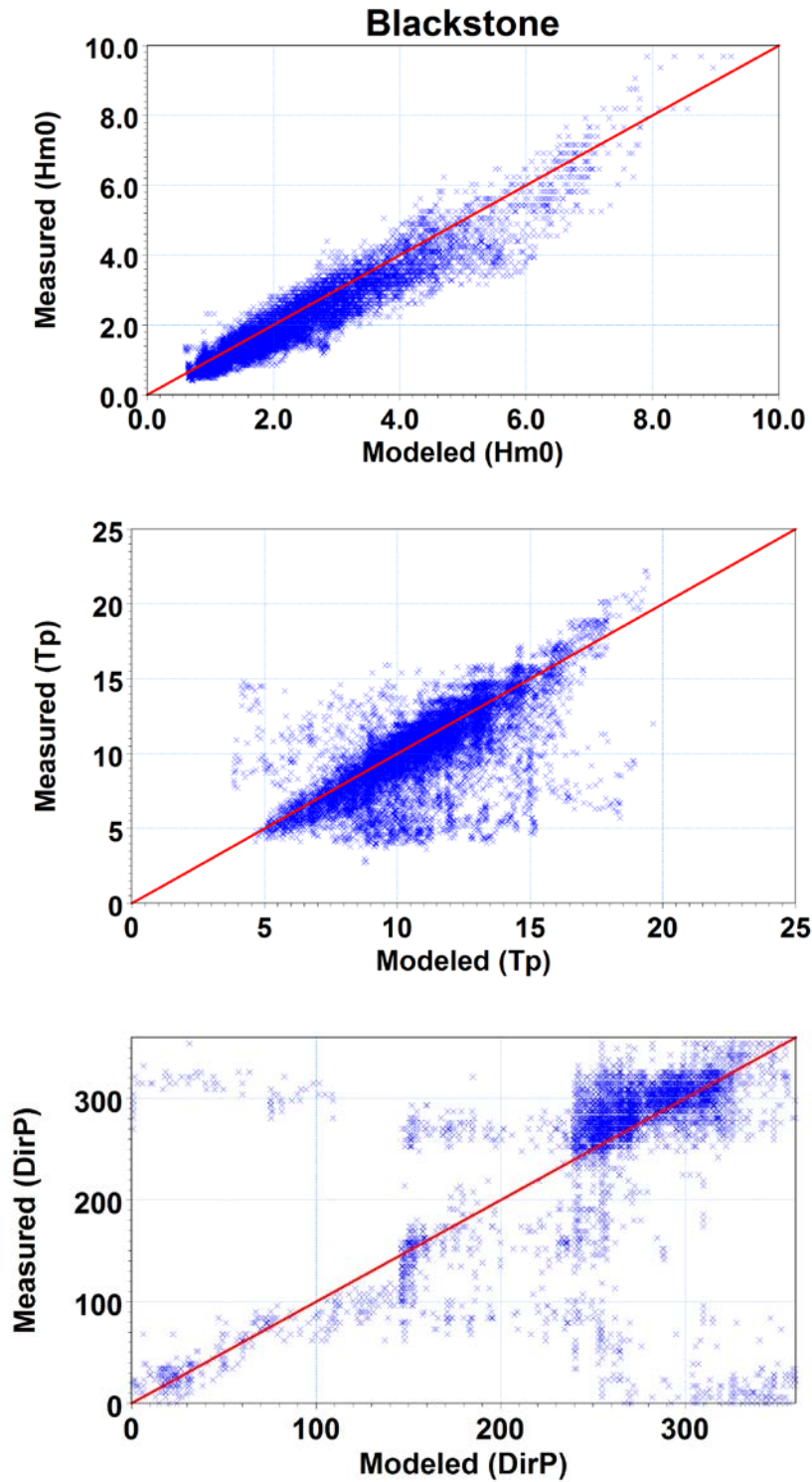


Figure 18. Correlation plots for significant wave height (H_{m0}), peak wave period (T_p) and peak wave direction (Dir_p) between model and measurements for Blackstone buoy, for January-December 2010.

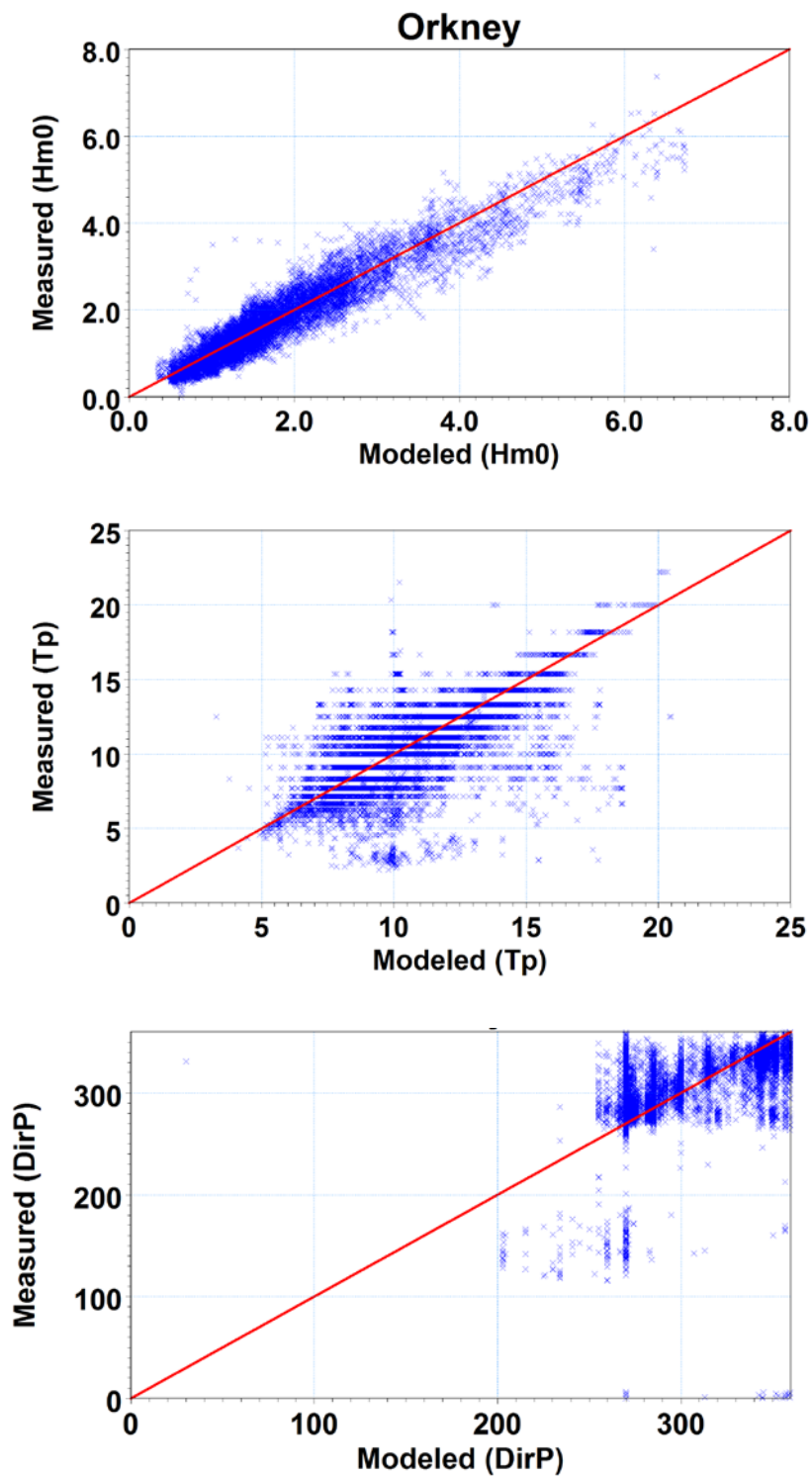


Figure 19. Correlation plots for significant wave height (H_{m0}), peak wave period (T_p) and peak wave direction (Dir_p) between model and measurements for Orkney buoy, for January-December 2010

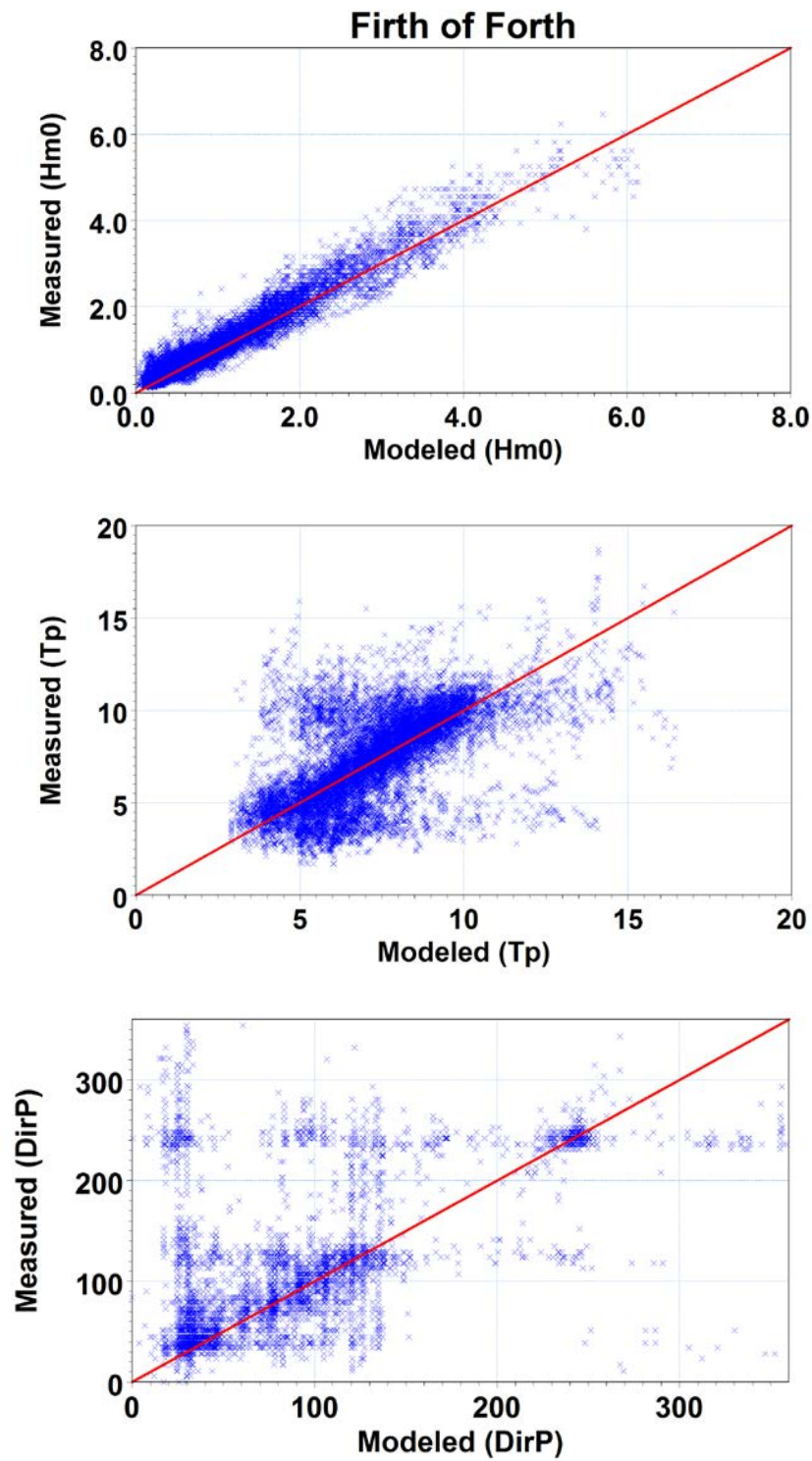


Figure 20. Correlation plots for significant wave height (H_{m0}), peak wave period (T_p) and peak wave direction (Dir_p) between model and measurements for Firth of Forth buoy, for January-December 2010

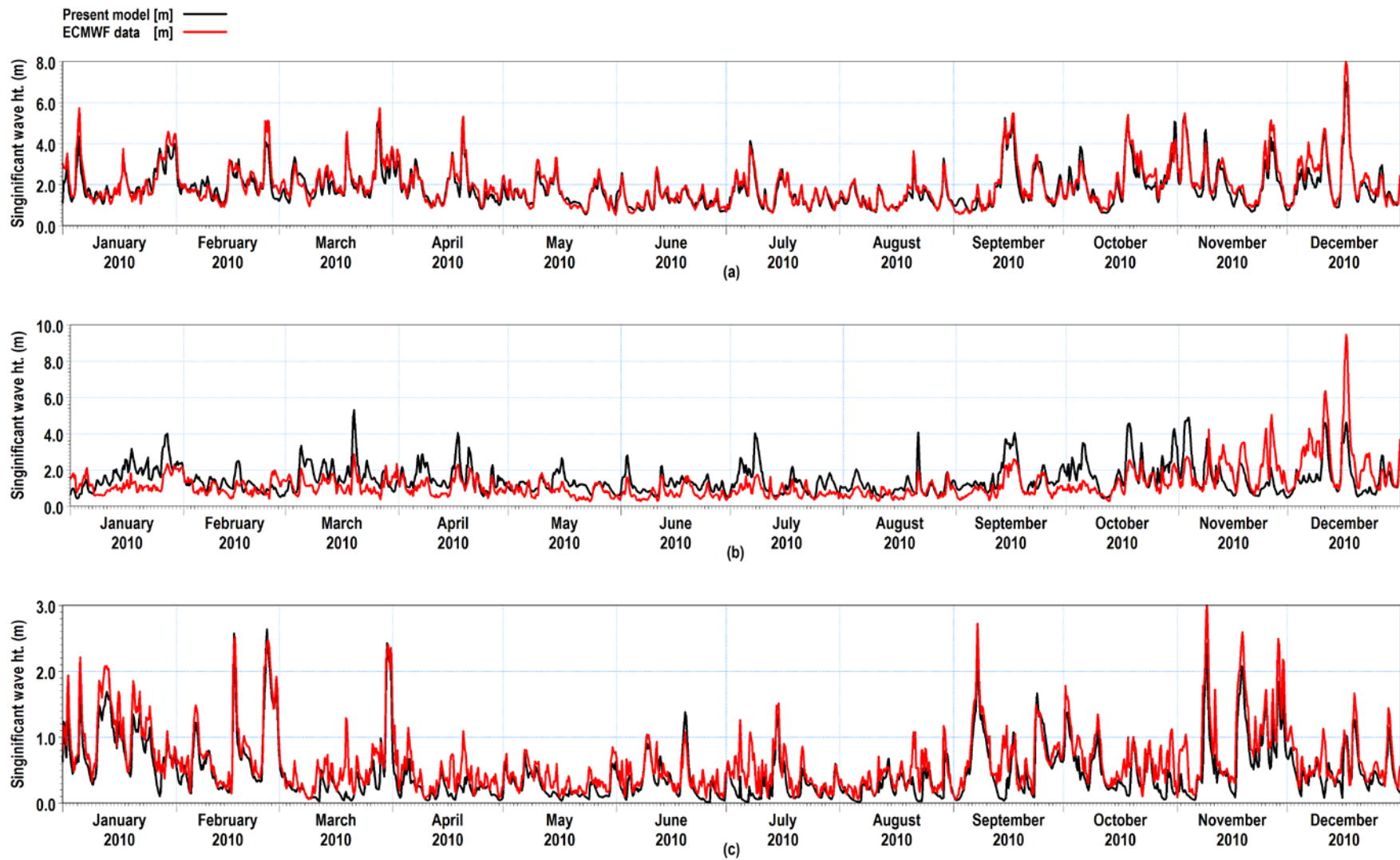


Figure 21. Comparison of significant wave height from MIKE21 and WAM models for shallow water locations: (a) Isle of Lewis, (b) Westray and (c) Dornoch for January-December 2010.

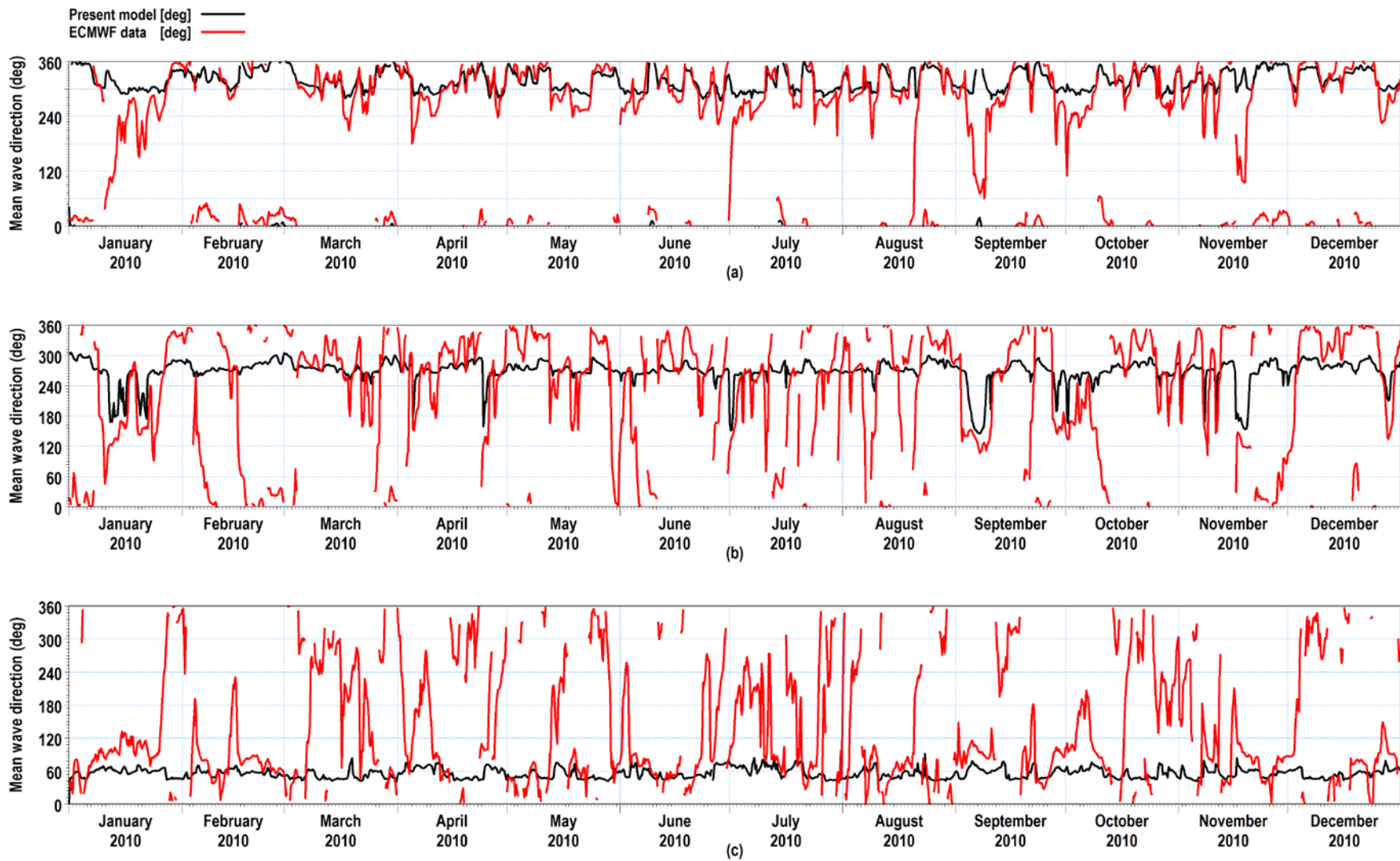


Figure 22. Comparison of mean wave direction from MIKE21 and WAM models for shallow water locations: (a) Isle of Lewis, (b) Westray and (c) Dornoch for January-December 2010.

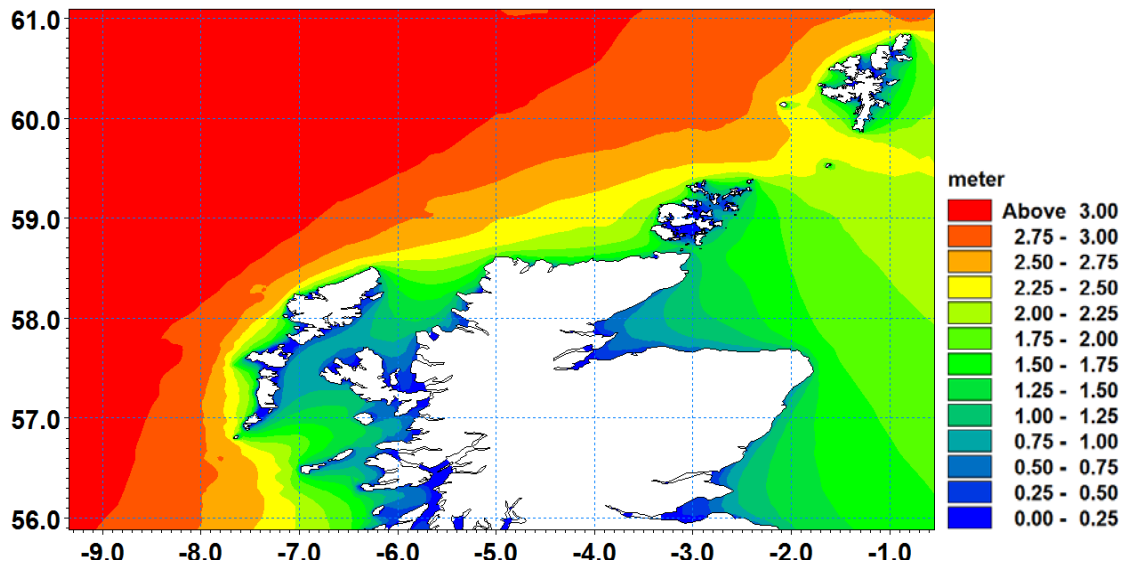


Figure 23. Mean significant wave height for January-December 2010

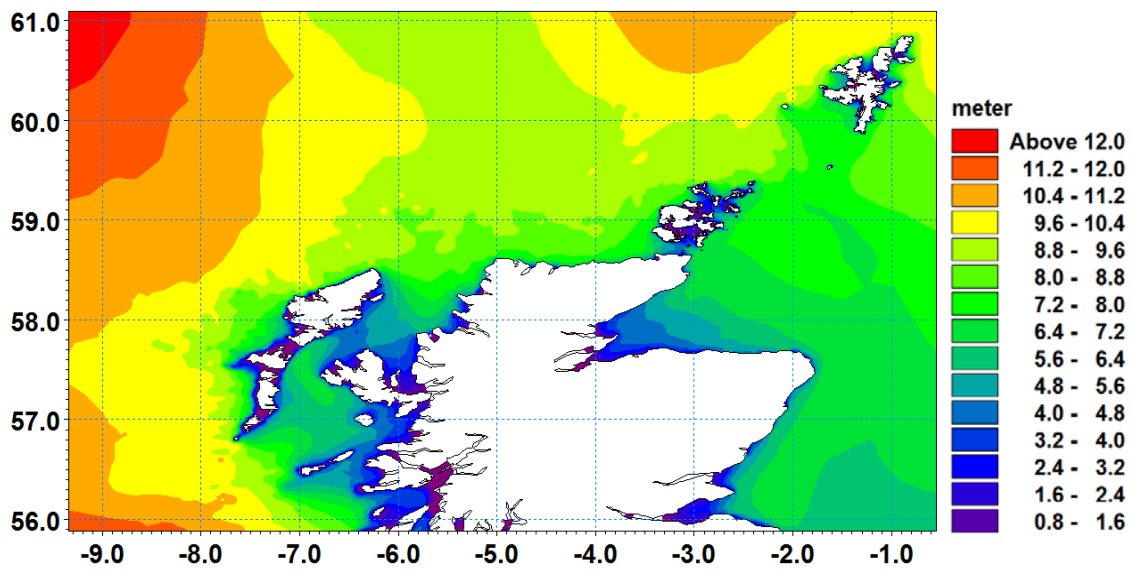


Figure 24. Maximum significant wave height for January-December 2010

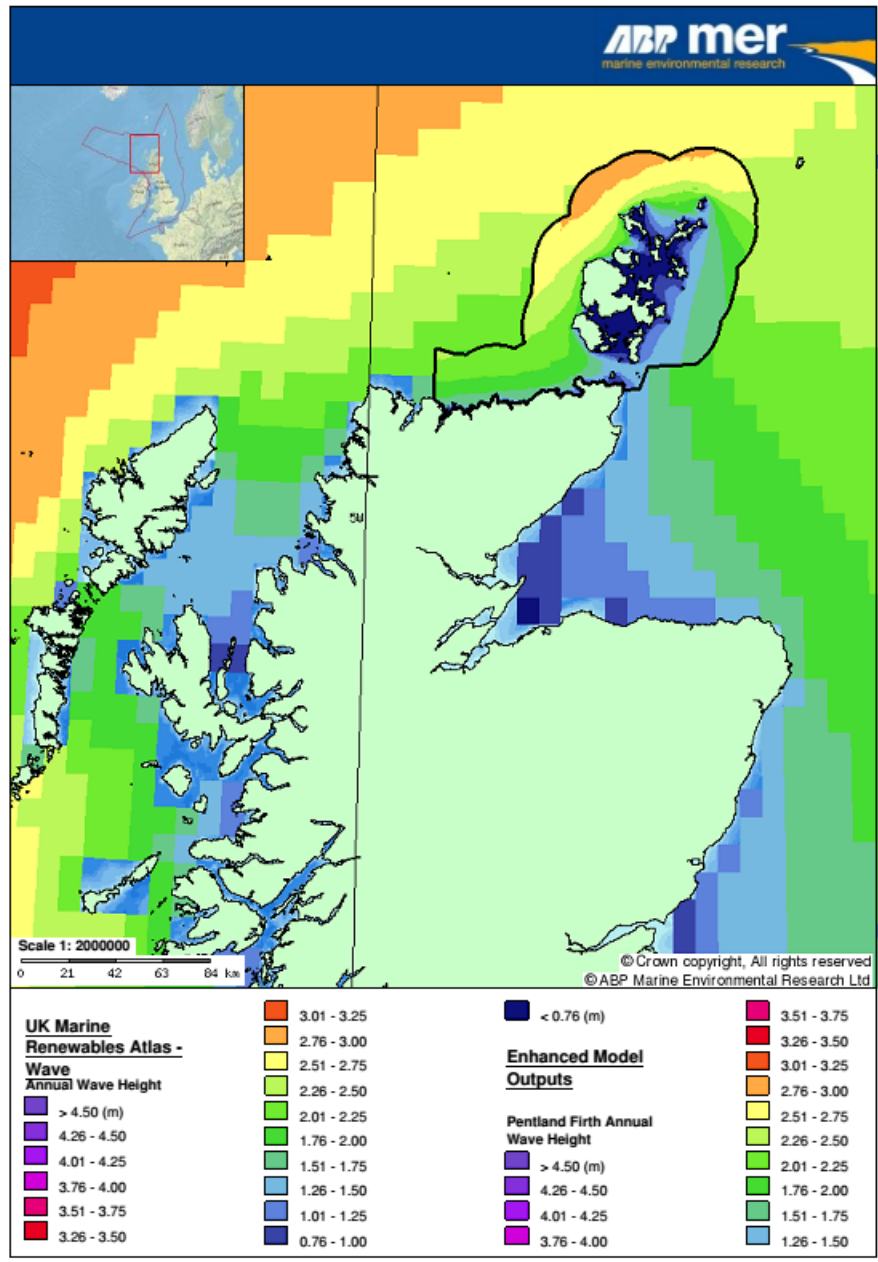


Figure 25. Annual significant wave height extracted from Atlas of UK Marine Renewable Energy Resources [ABPmer, 2008], Reproduced from <http://www.renewables-atlas.info/> © Crown Copyright.

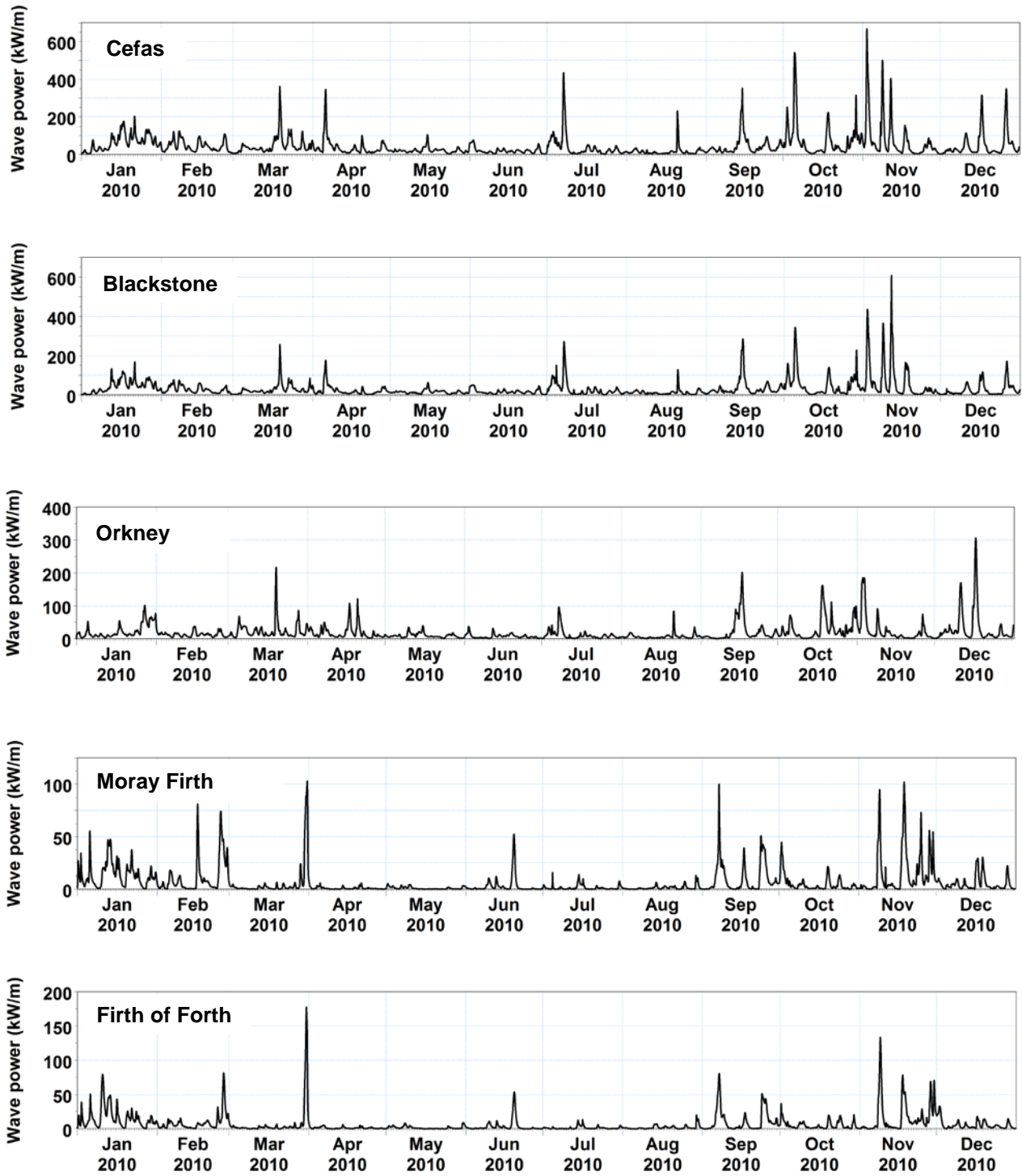


Figure 26(a). Wave power computed January-December 2010, from top to bottom, for Cefas, Blackstone, Orkney, Moray Firth and Firth of Forth.

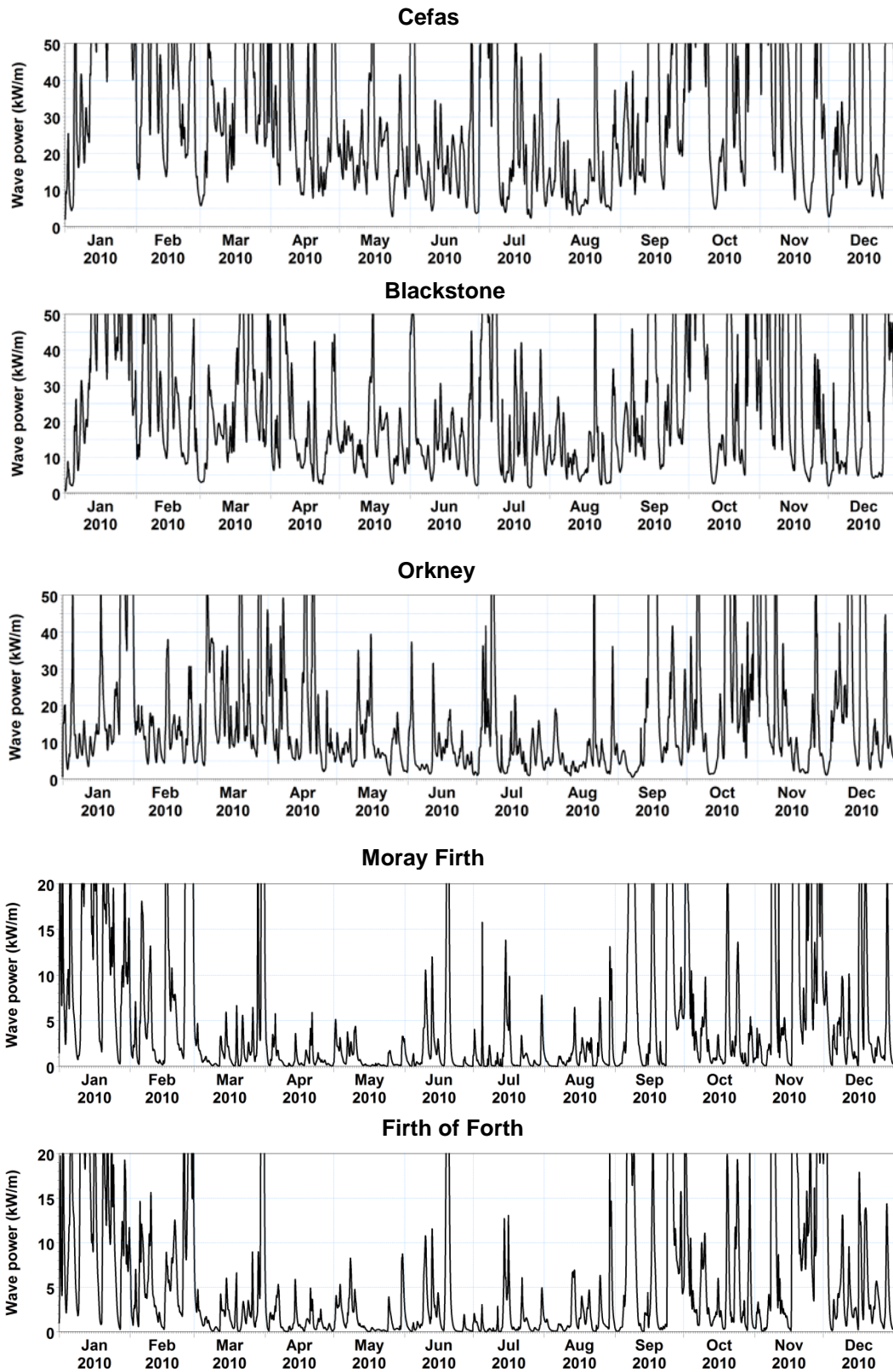


Figure 26(b). Wave power computed January-December 2010. Same as in Figure 26 (a), but with enlarged vertical scale: Cefas, Blackstone, Orkney, Moray Firth and Firth of Forth.

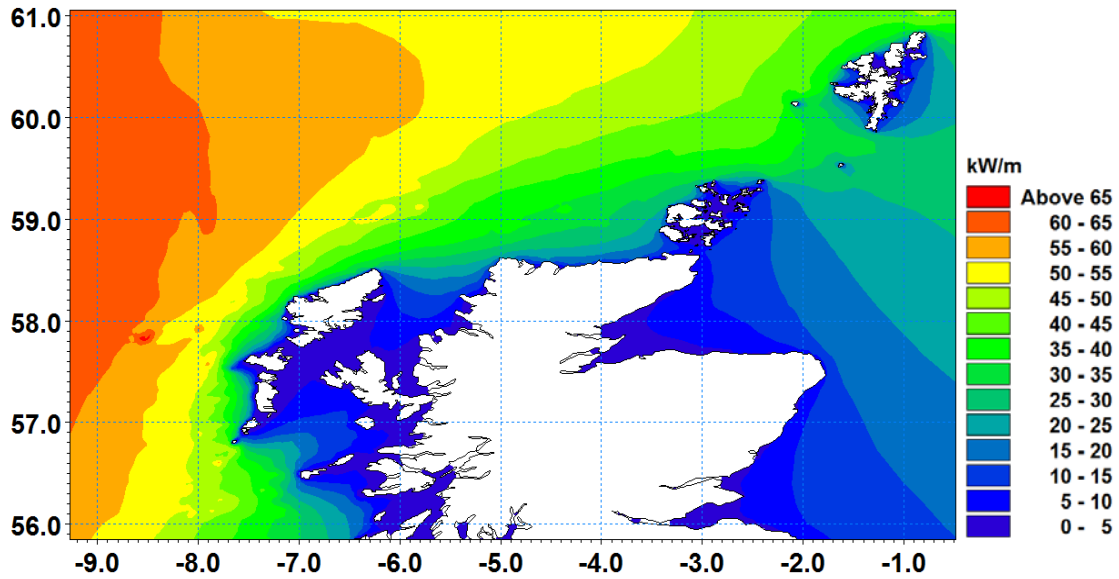


Figure 27. Mean wave power for January-December 2010

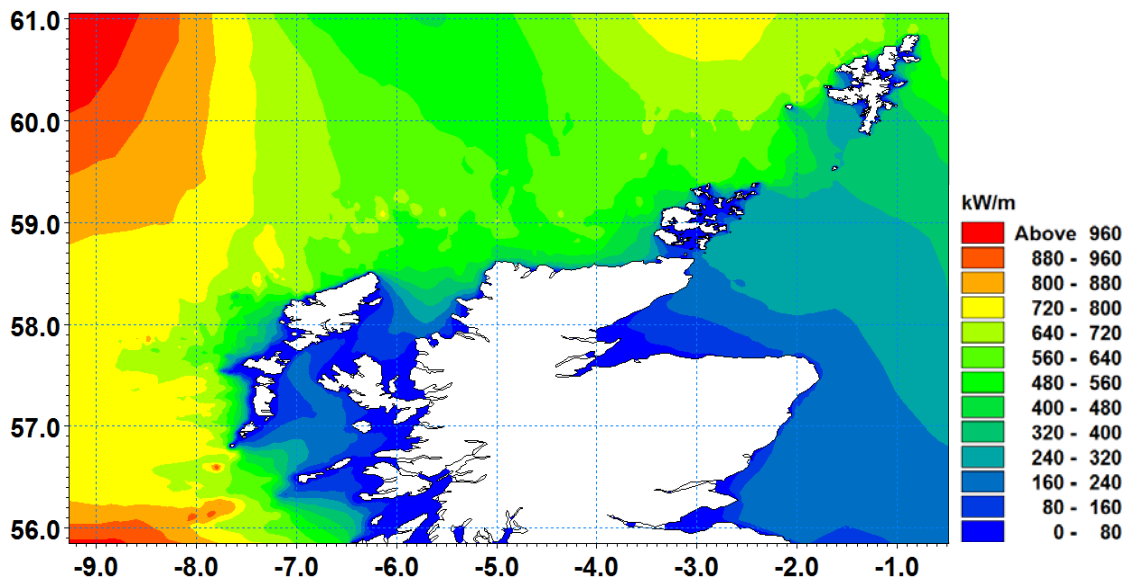


Figure 28. Maximum wave power for January-December 2010

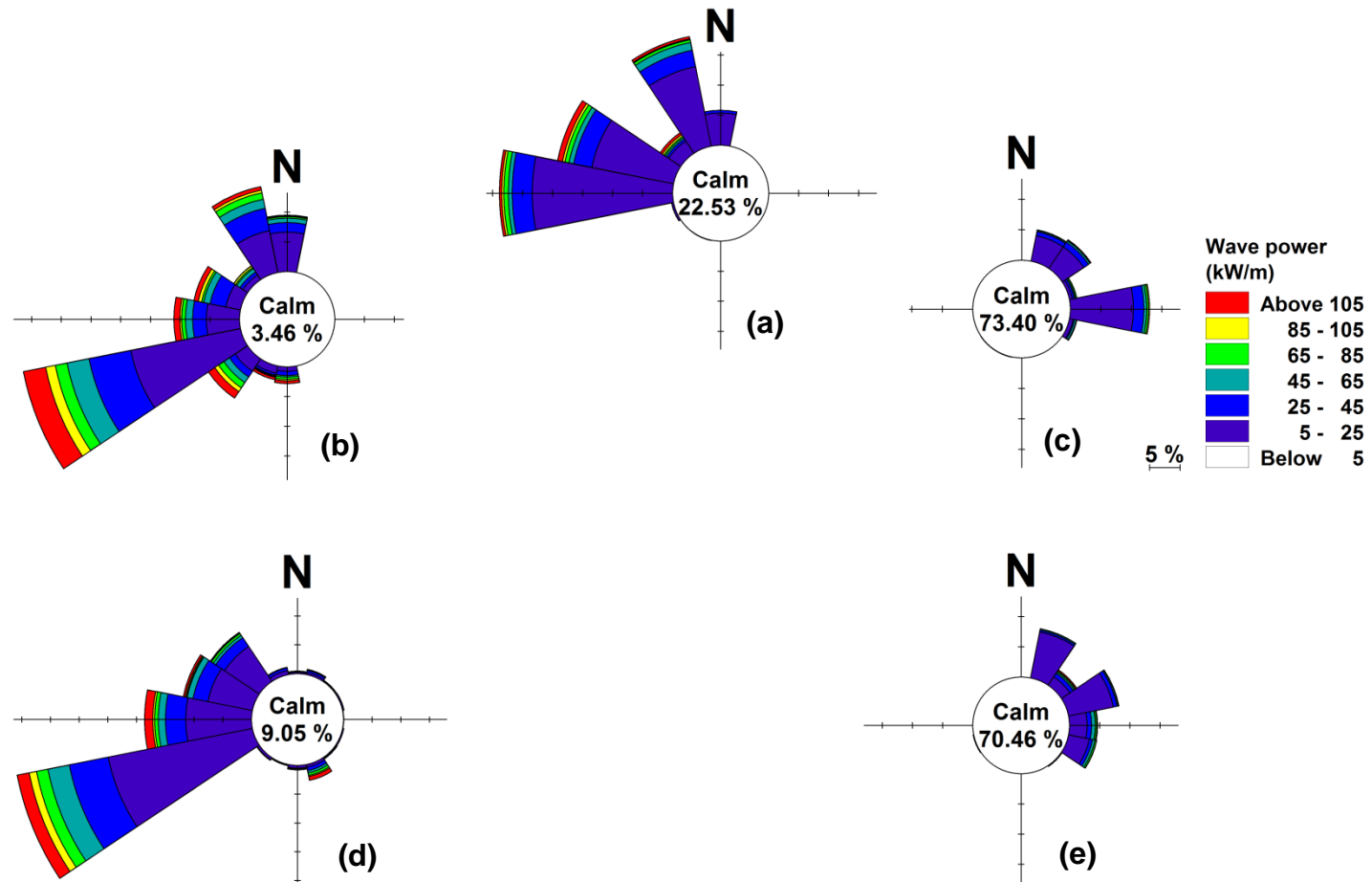


Figure 29. Rose plots for wave power with peak wave direction for different locations: (a) Orkney, (b) Cefas, (c) Moray Firth, (d) Blackstone and (e) Firth of Forth. Calculated from model results for January-December 2010.